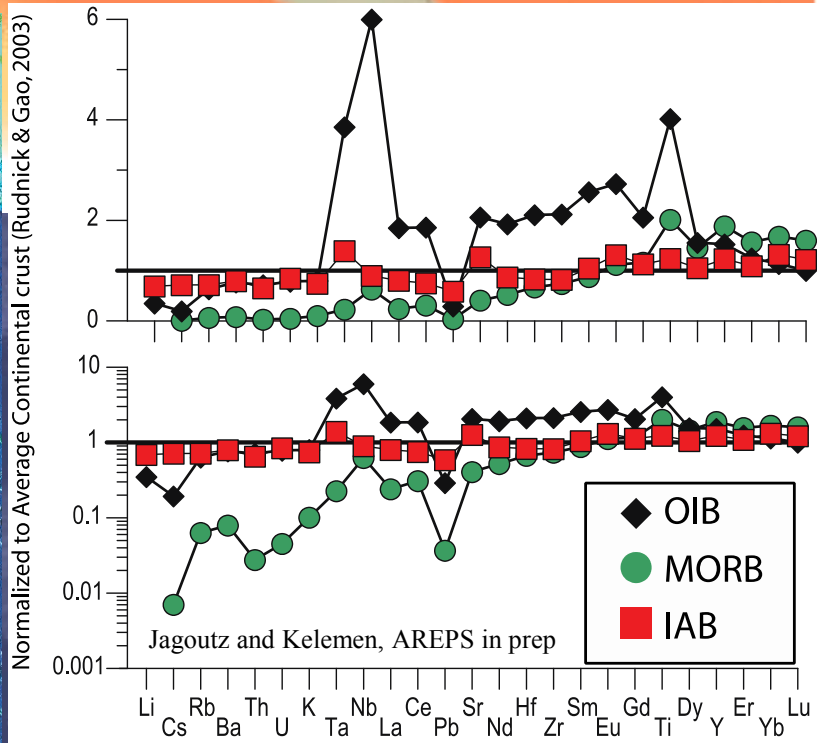
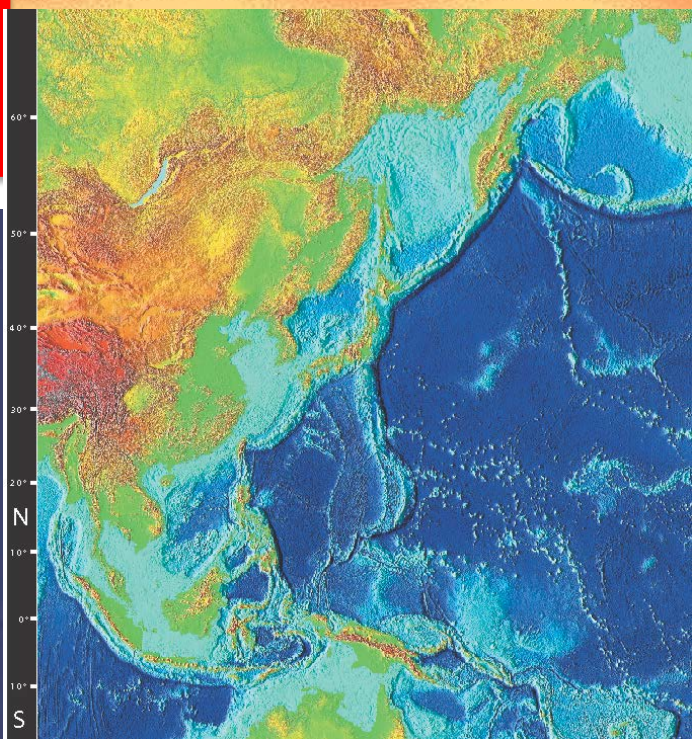
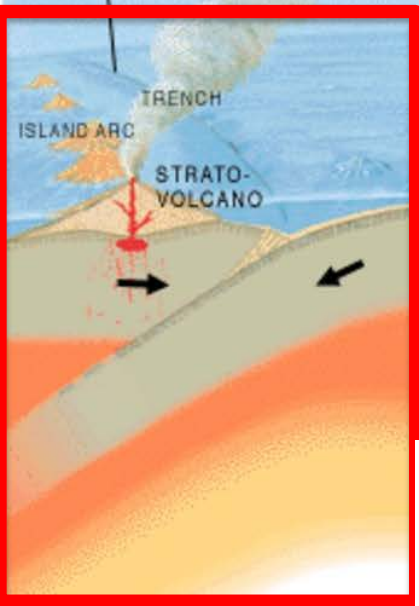
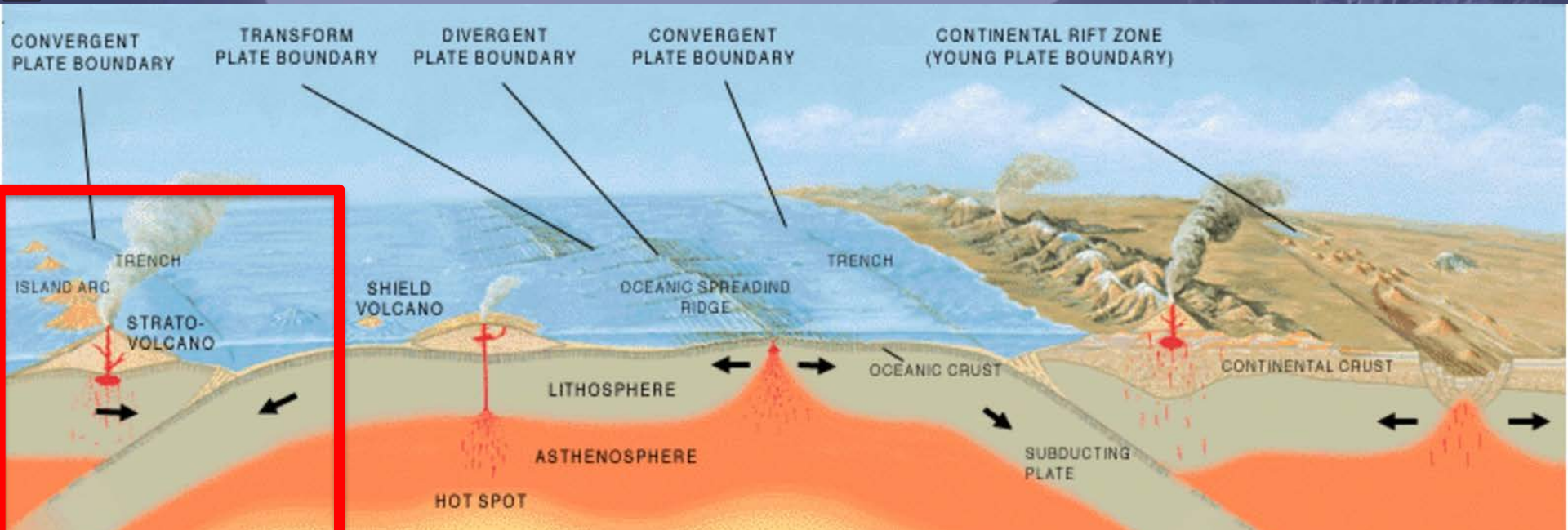
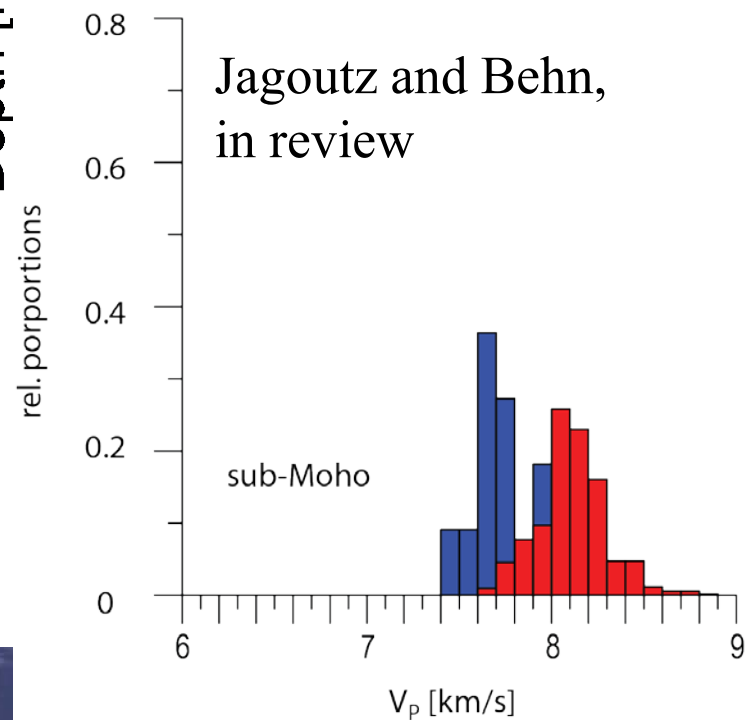
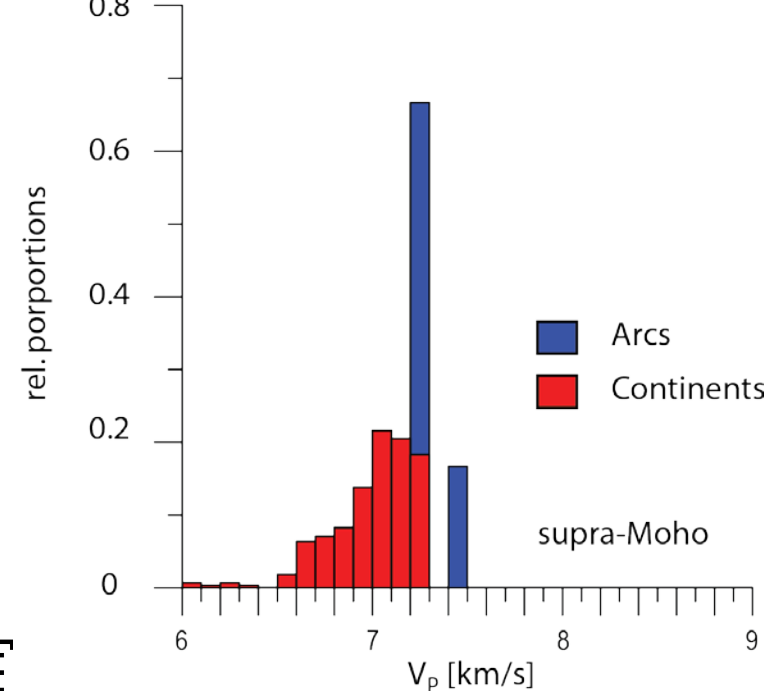
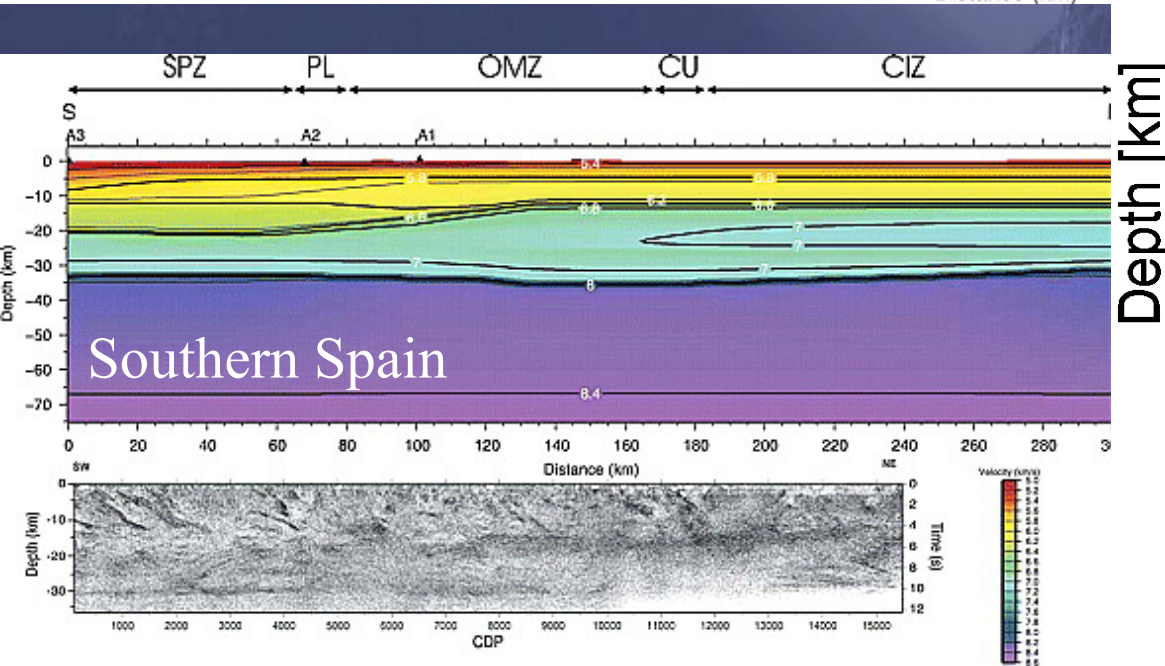
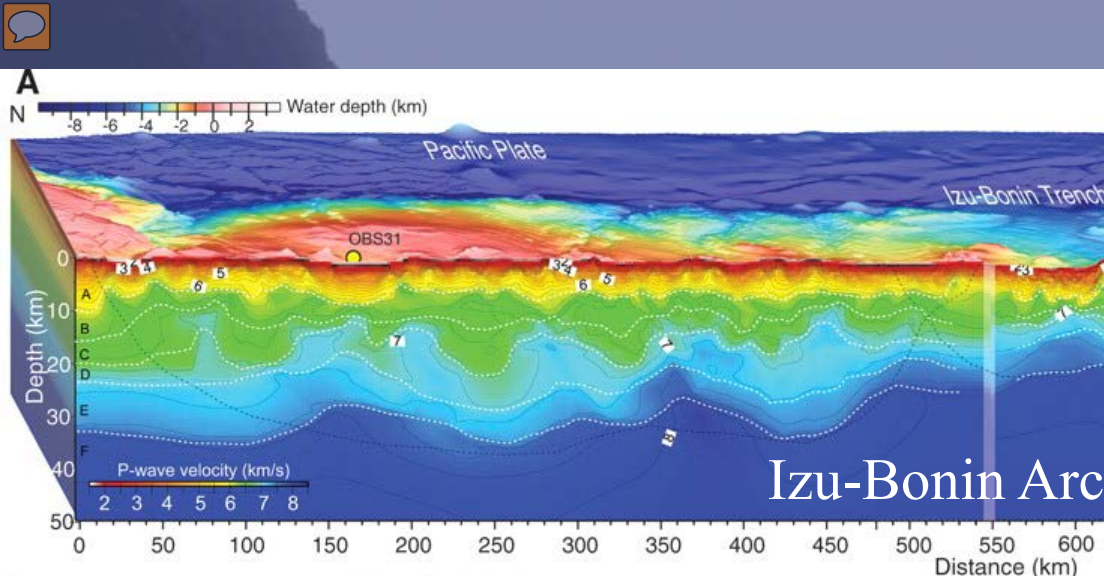


A world map showing seismic activity, with colors ranging from blue (low activity) to red (high activity). The map is overlaid with a semi-transparent blue rectangle containing text and a list of names. The map shows high seismic activity along plate boundaries, particularly in the Pacific and Indian Oceans, and around the Atlantic and African continents. A color scale bar is located at the bottom left of the map.

“The formation of continental crust in arcs: The seismological perspective”

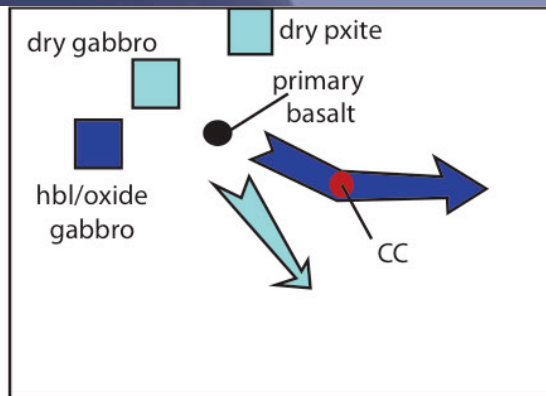
- Oliver Jagoutz (MIT)
- Mark Behn (WHOI)
- Peter Kelemen (LDEO)





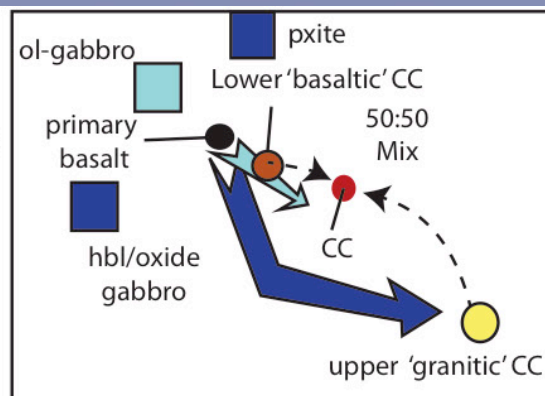


molar Mg#



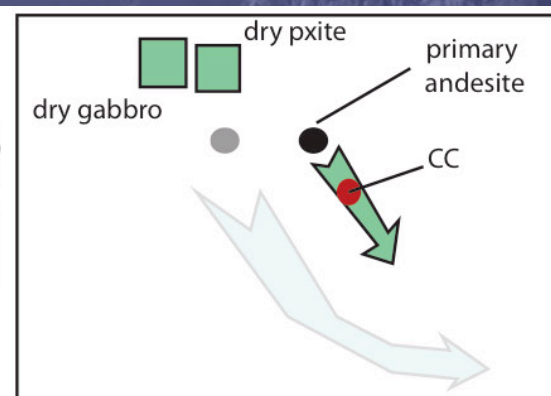
wt% SiO₂

molar Mg#

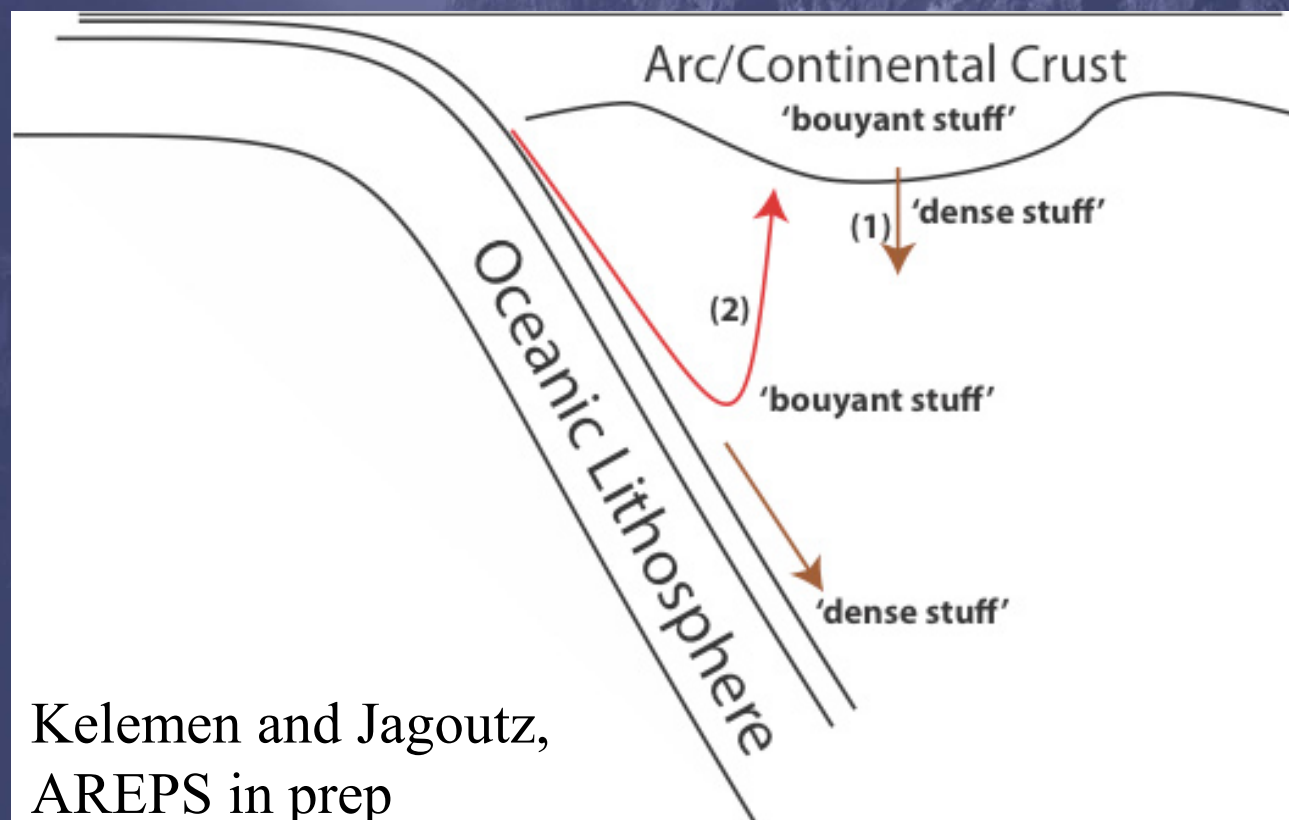


wt% SiO₂

molar Mg#

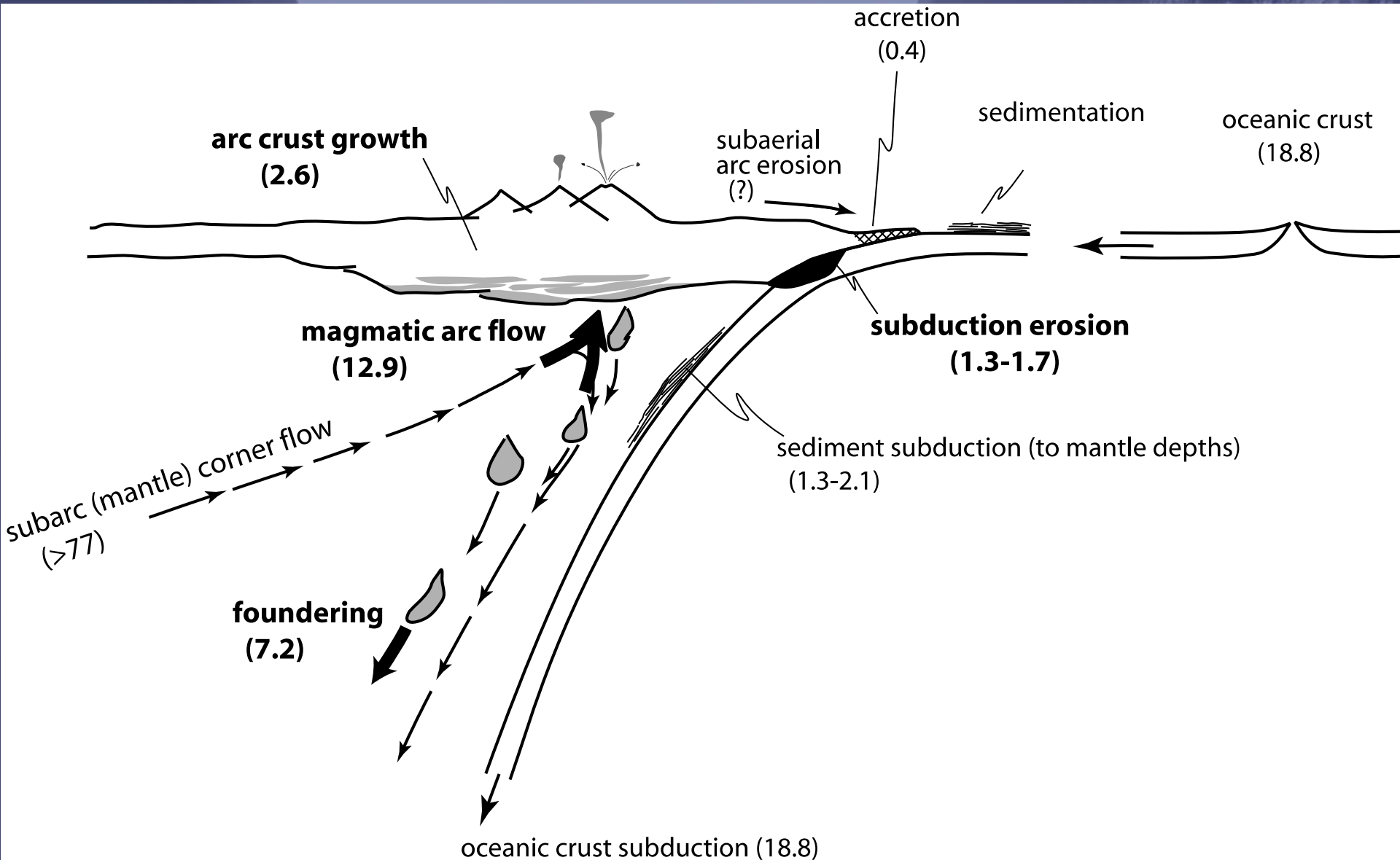


wt% SiO₂

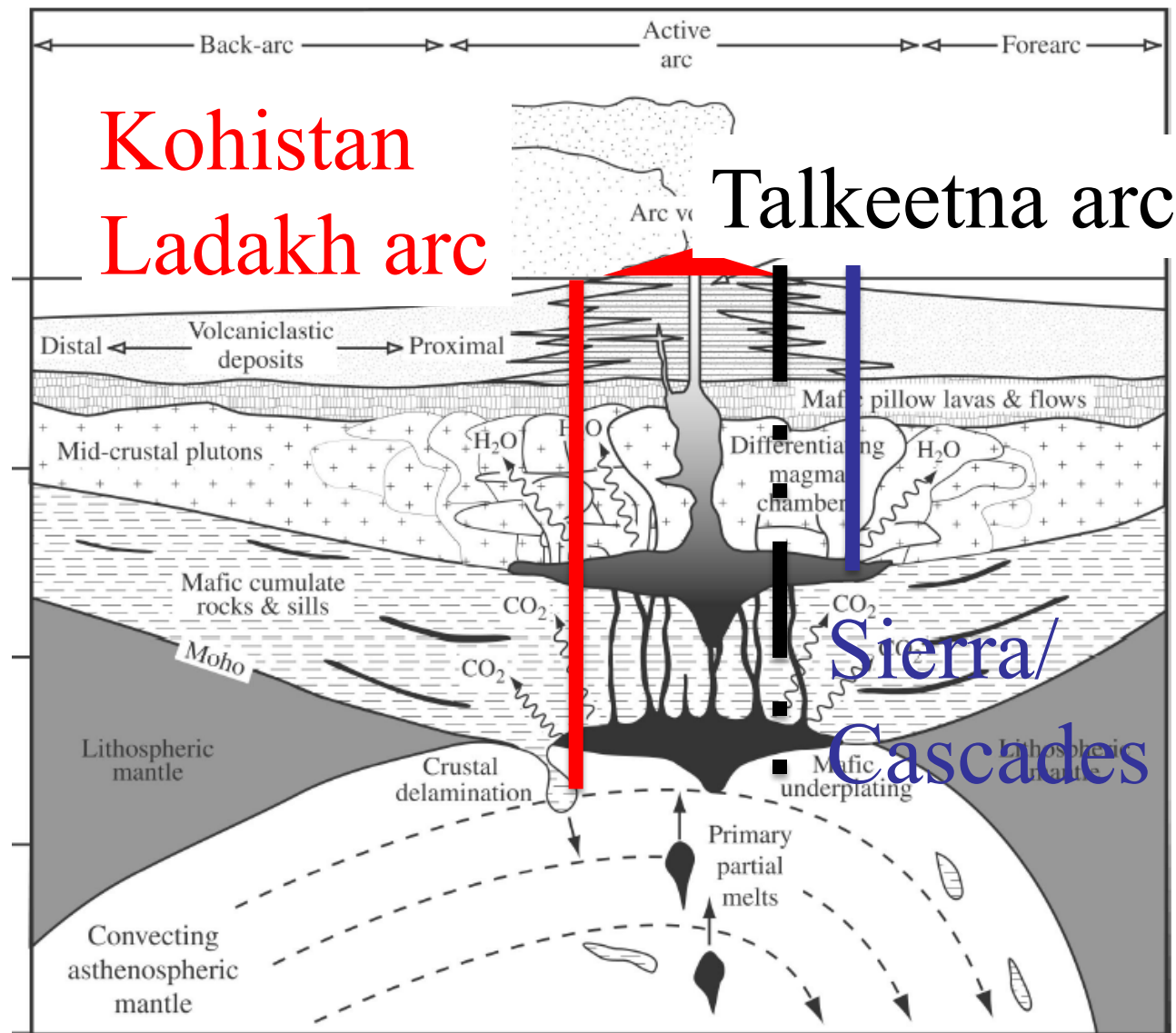


Kelemen and Jagoutz,
AREPS in prep

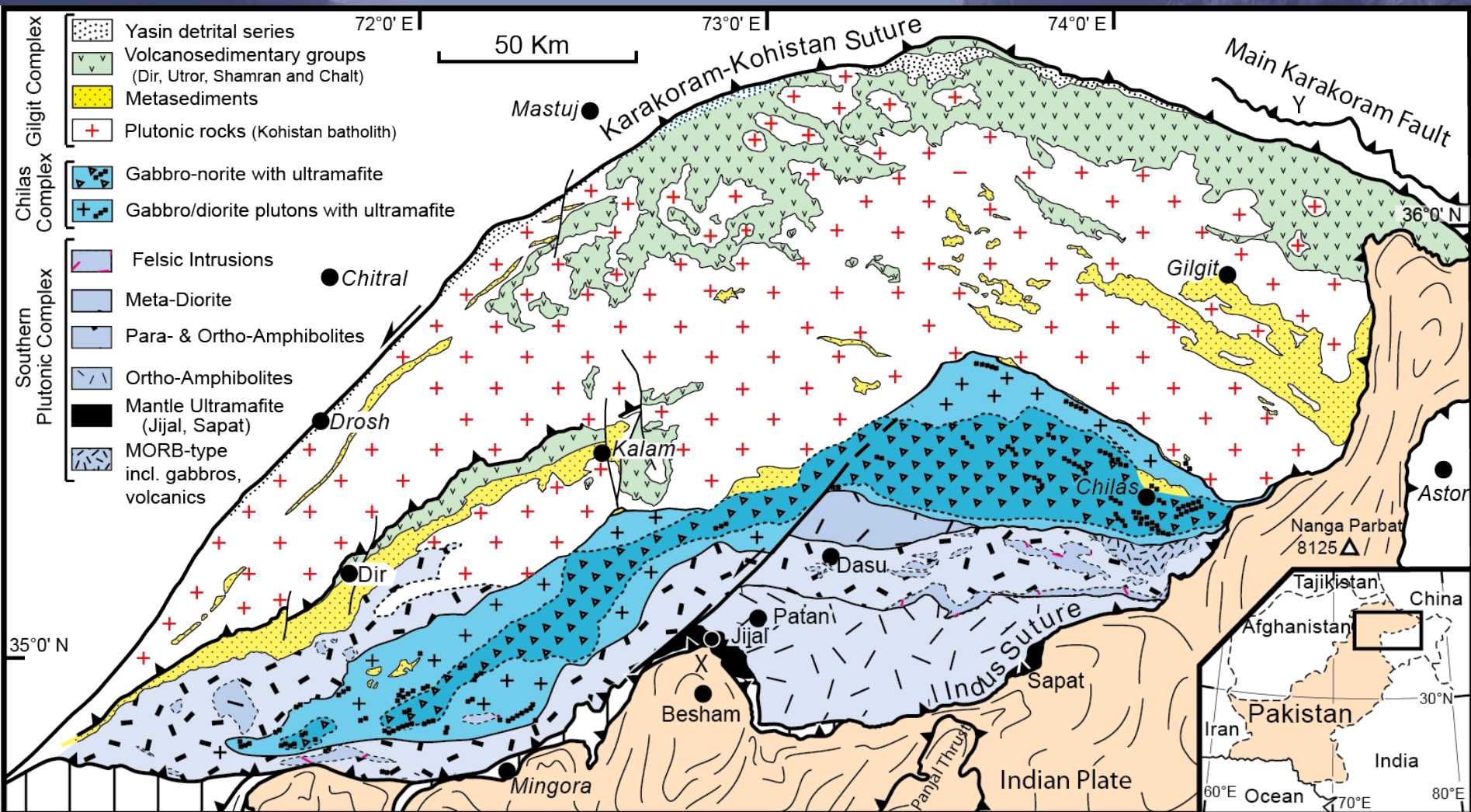
Subduction zone mass balance [km^3/a]



Exposed arc sections



Stern, 2002

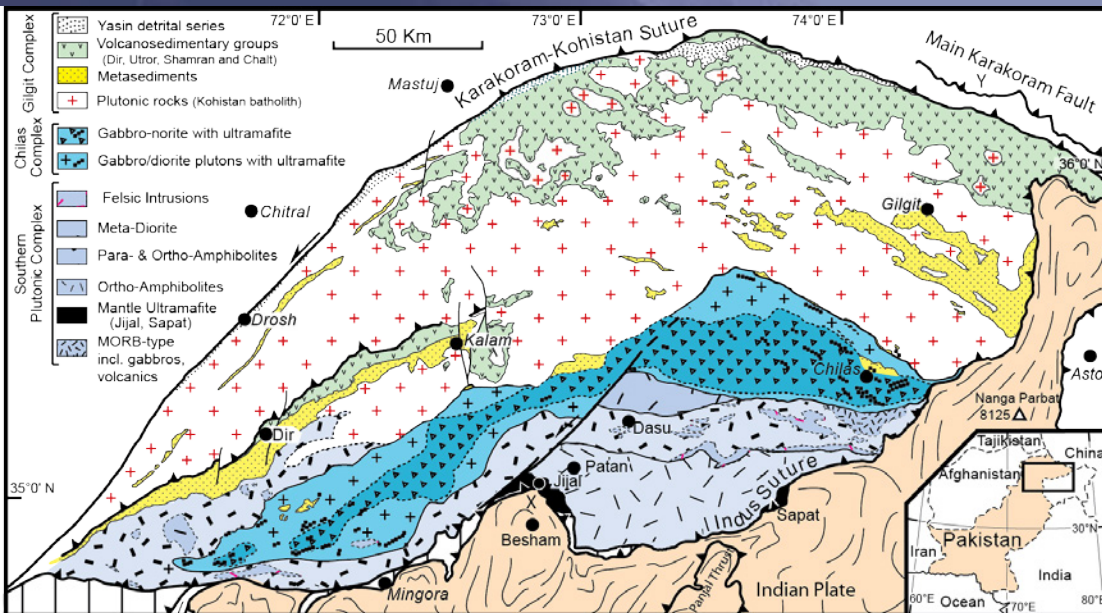


Jagoutz et al., 2011, EPSL

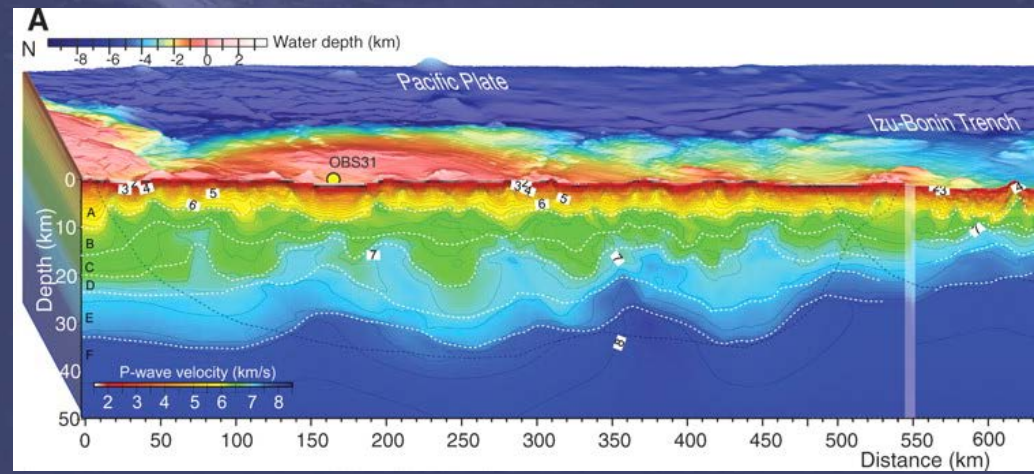
Depth (km)



How to 'translate' an exposed arc section

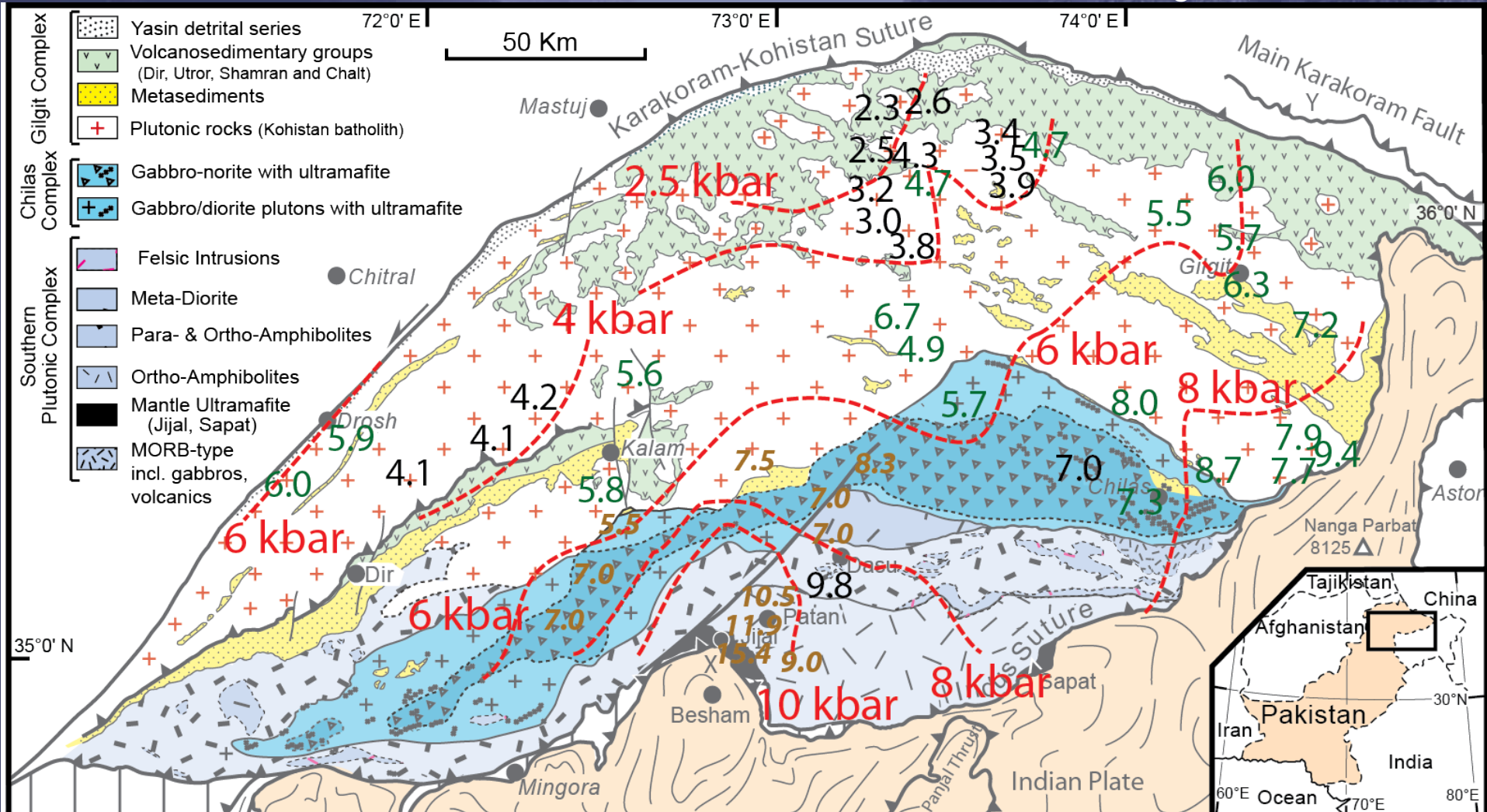


???



How to 'translate' an exposed arc section

Jagoutz, EPSL in review



Composition vs Depth

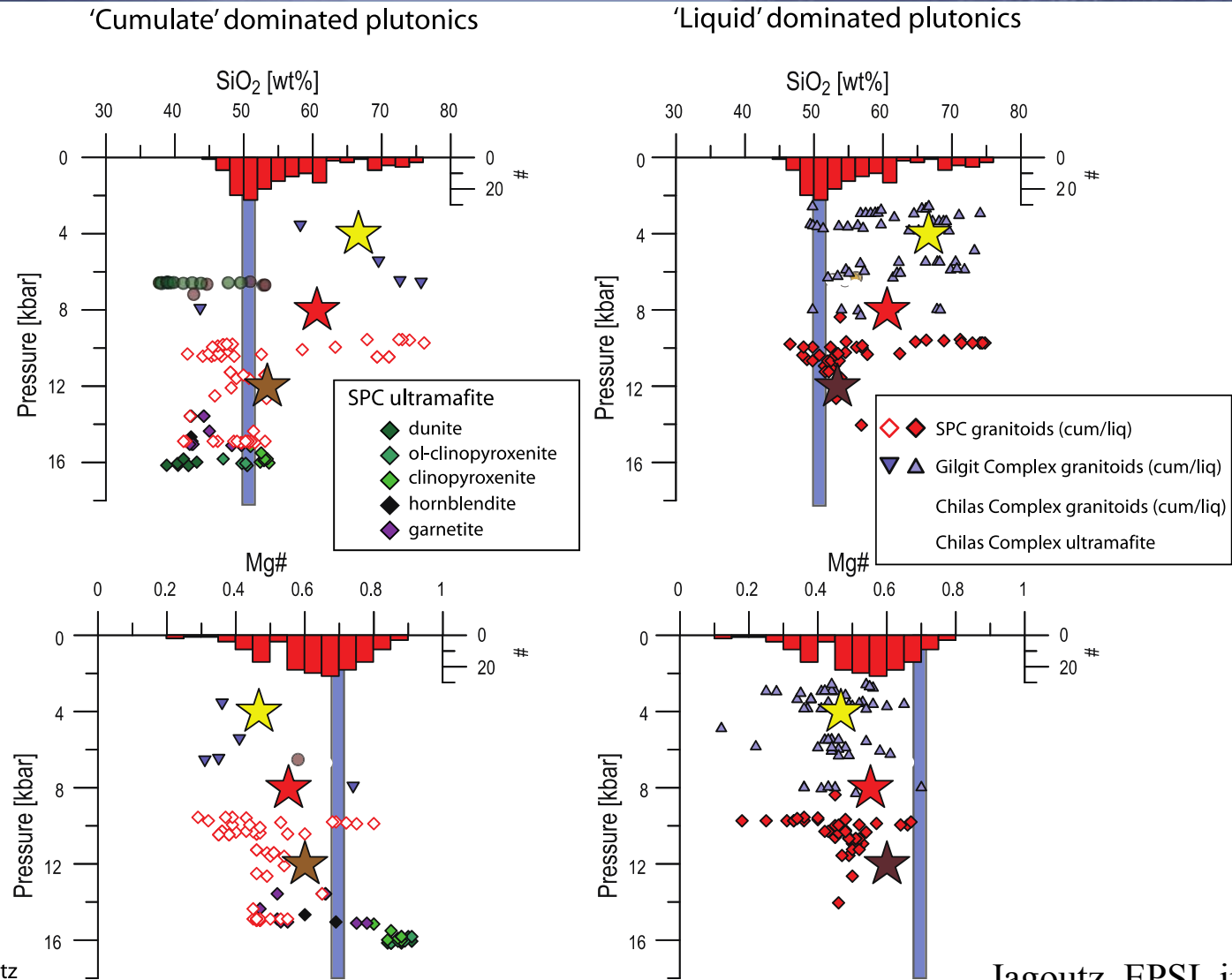
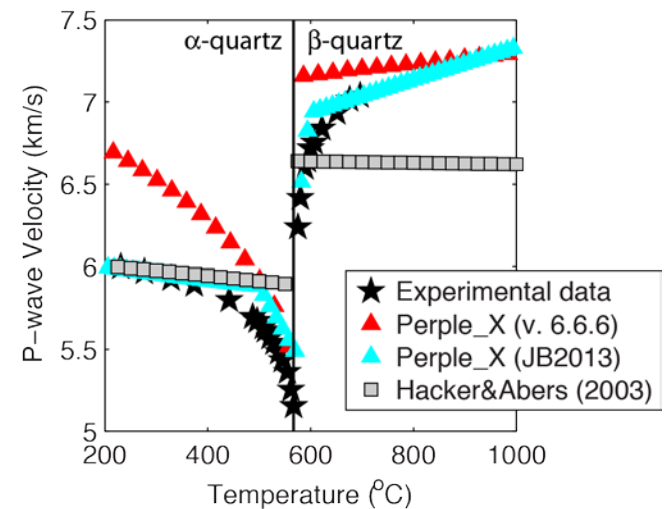
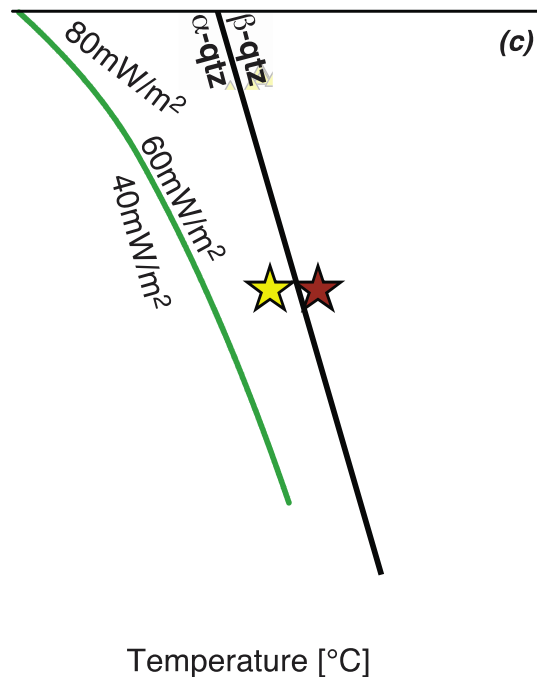
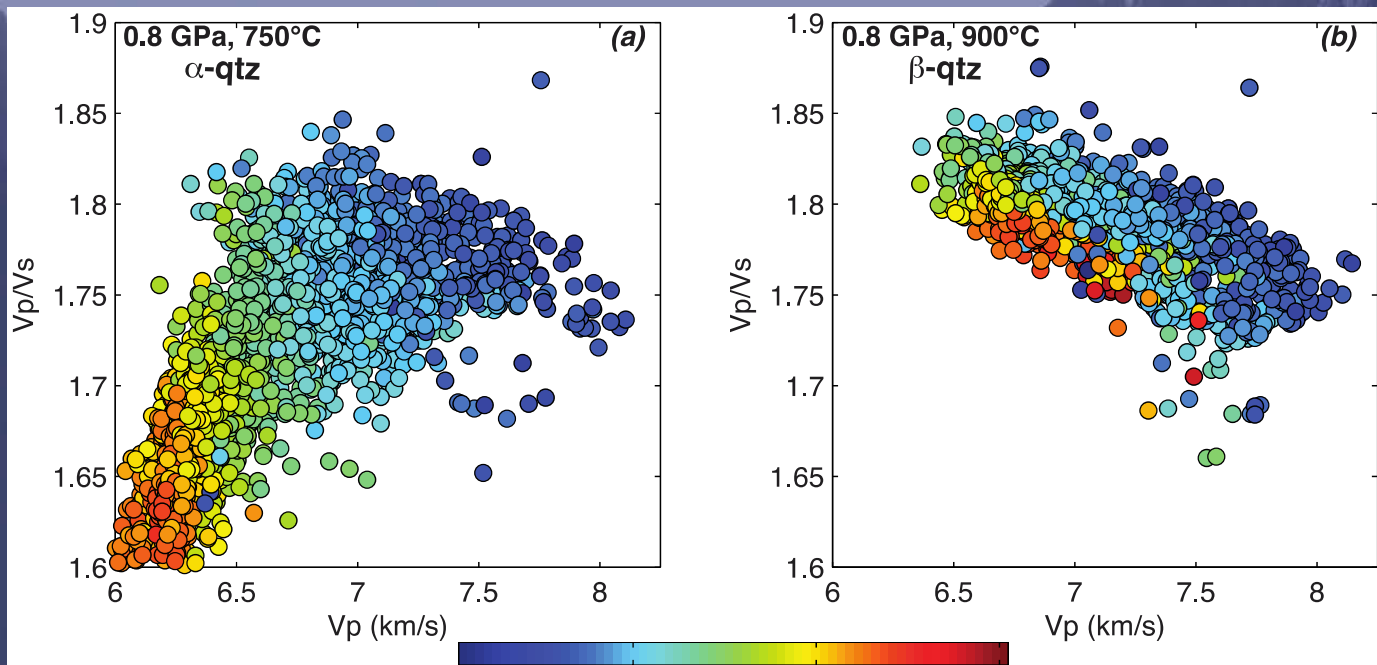
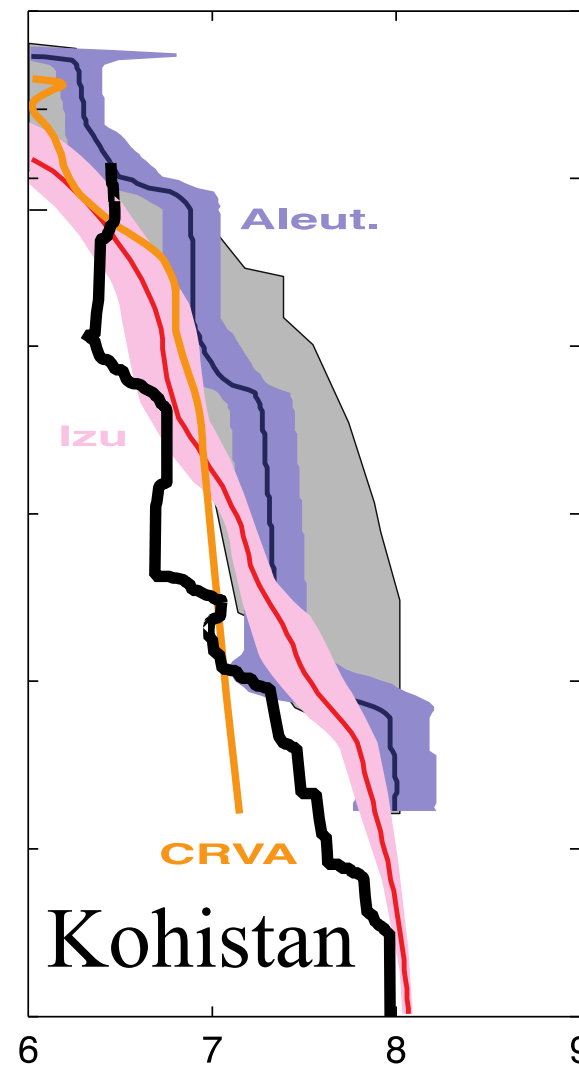
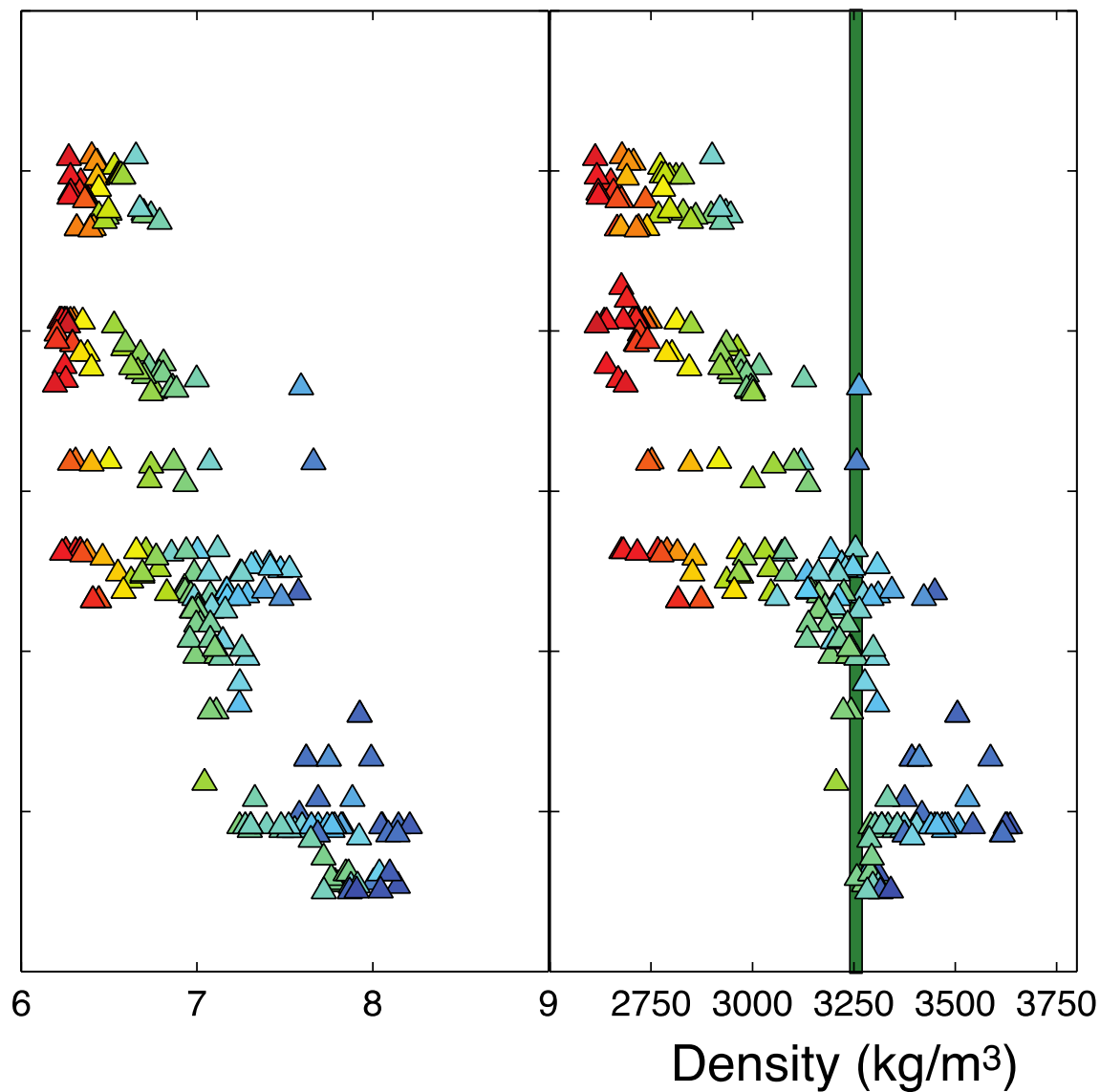
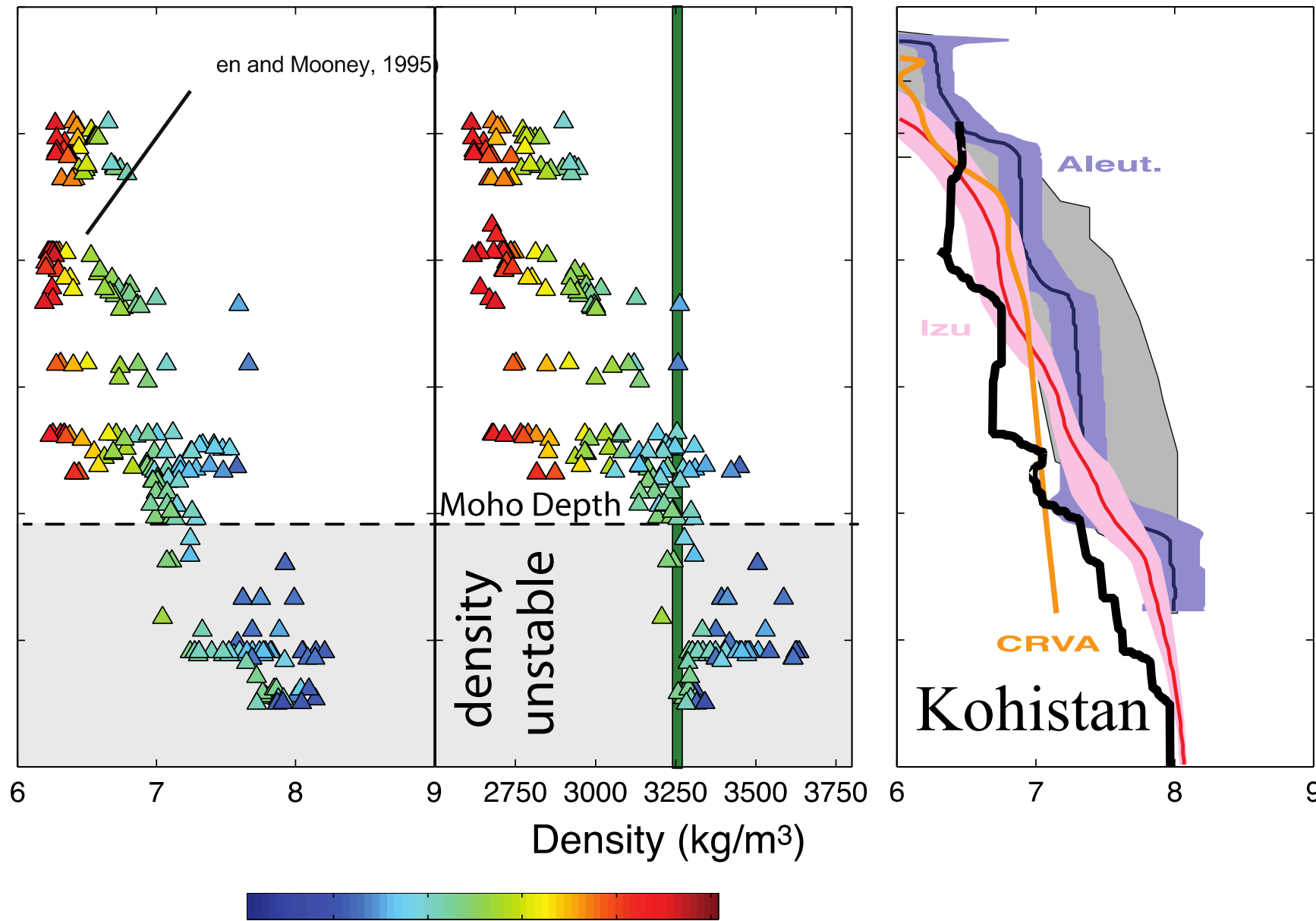


Fig 3 Jagoutz



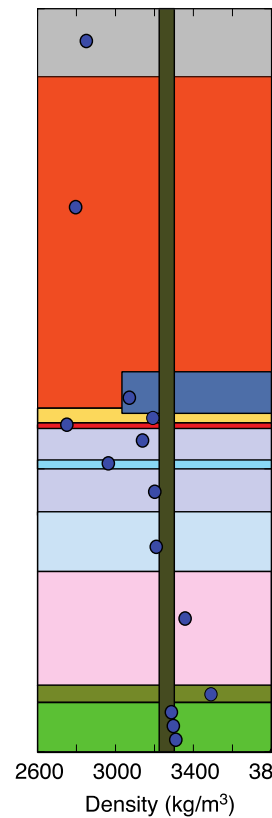
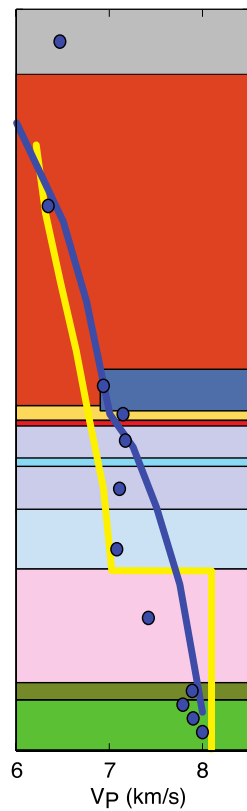
Jagoutz and Behn, under review



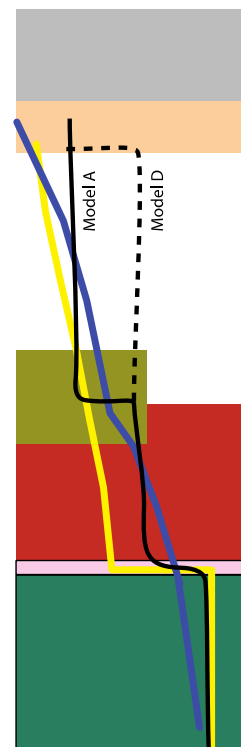
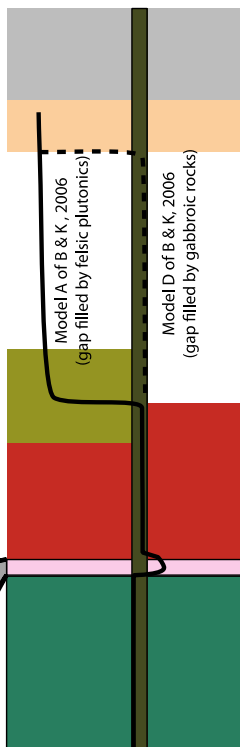




Depth (km)



Density unstable



eechina
qtzdiorite & tonalite,
both ± garnet

Gap

Jagoutz and Behn, under review

Magmatic growth of a cumulate layer

$$h = q\alpha t$$

h = Cumulate layer thickness

t = time

q = magma flux

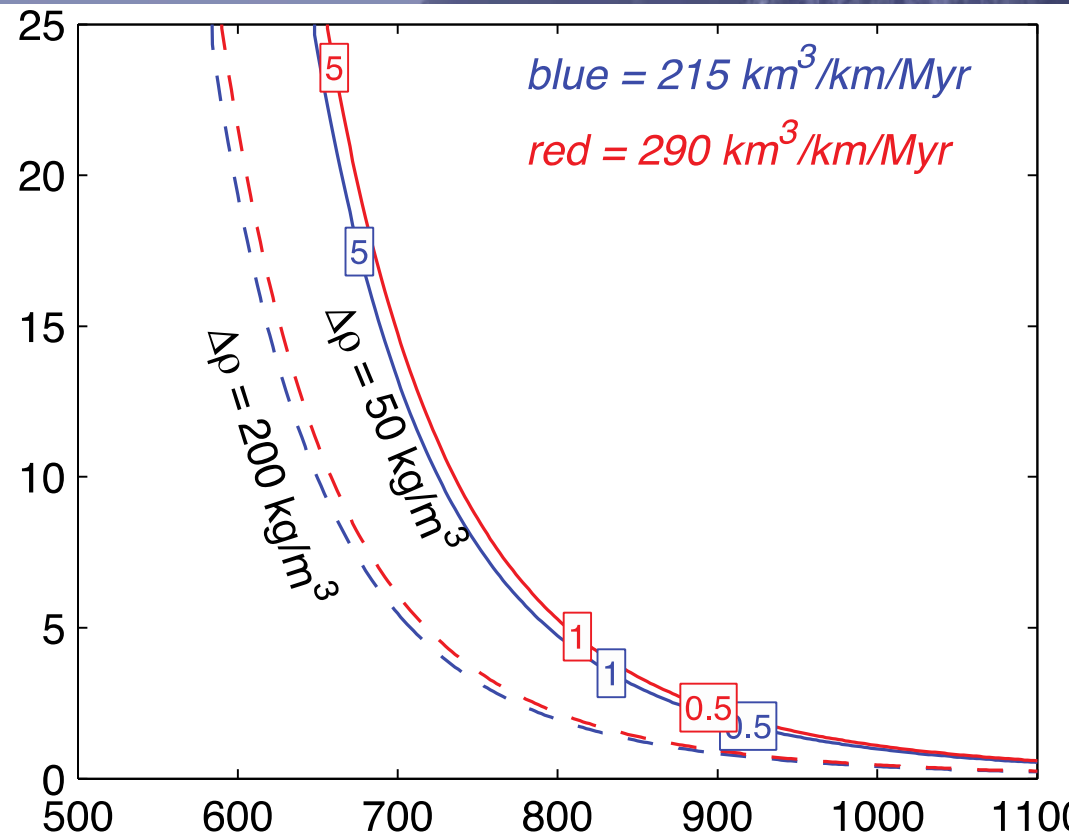
α = fractionation factor

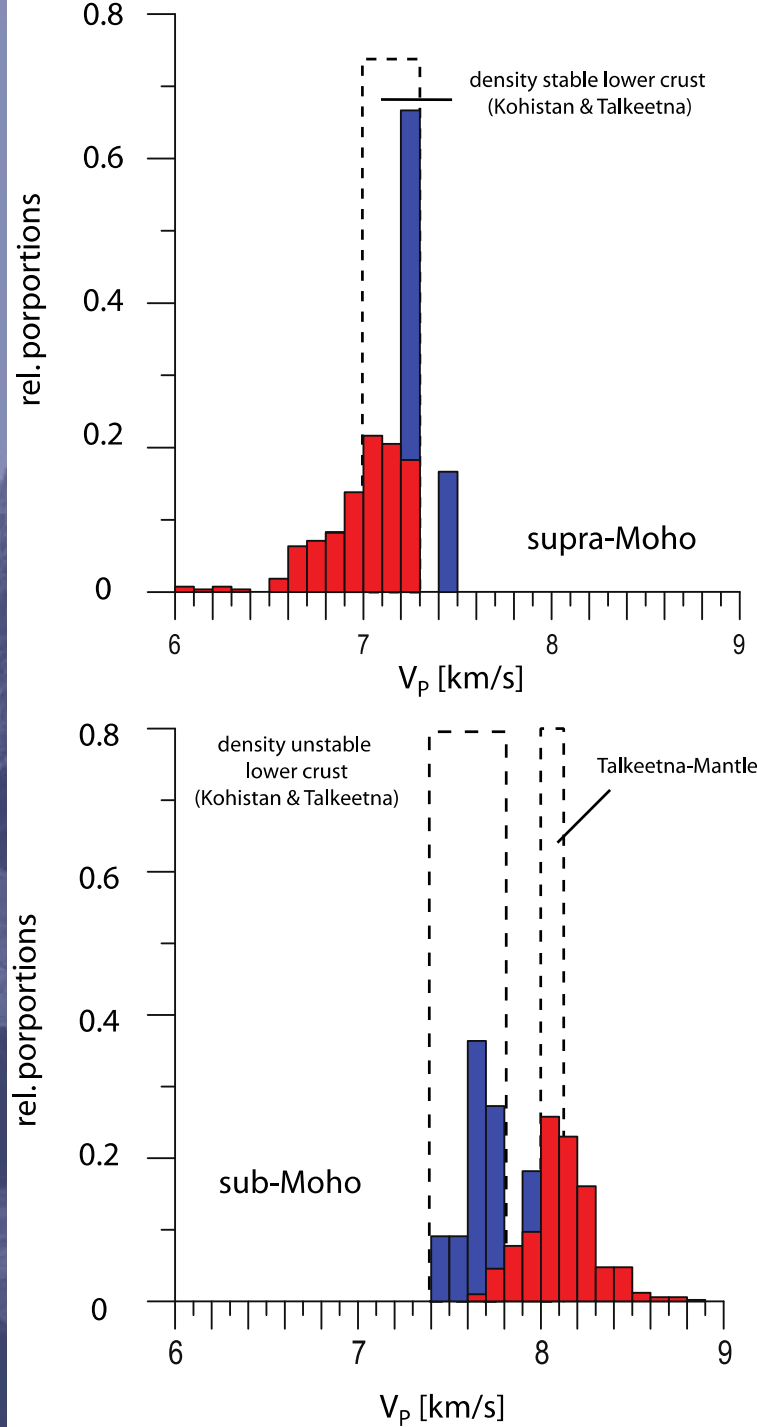
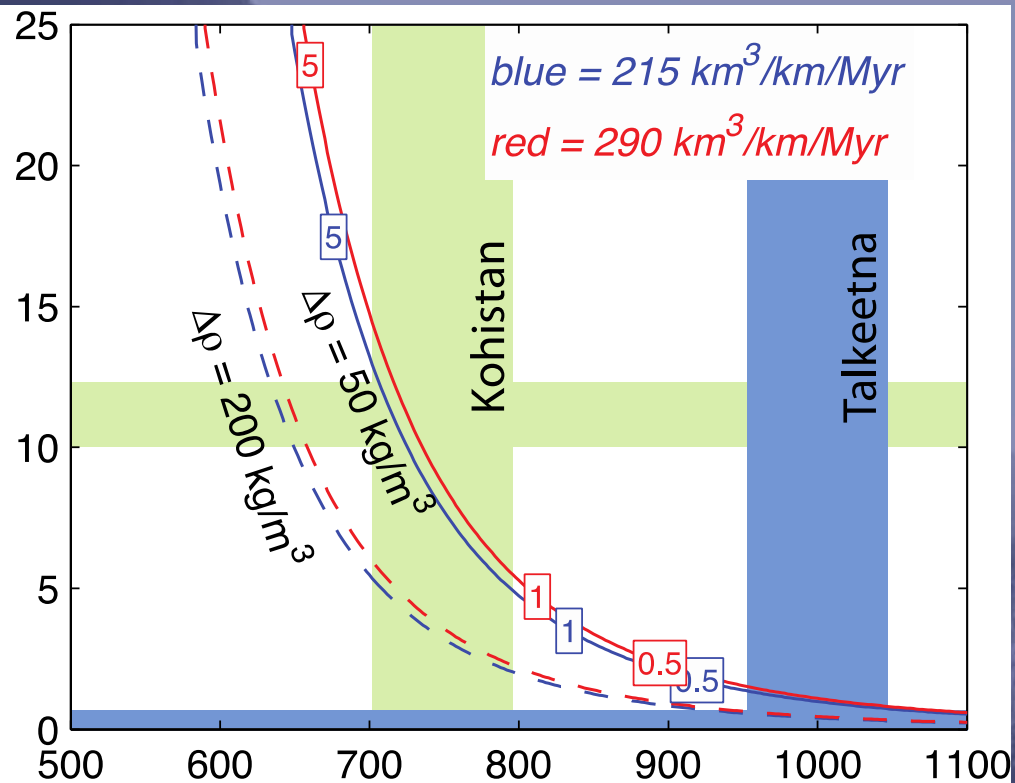
Instability growth rates (after Jull and Kelemen, 2001)

$$t_b = \left(\frac{n}{C'} \right)^n \frac{Z_0^{(1-n)}}{n-1} S$$

$$S = \left(\frac{B}{2\Delta\rho gh} \right)^n$$

$$B = A^{-1/n} \exp\left(\frac{Q}{nRT}\right)$$





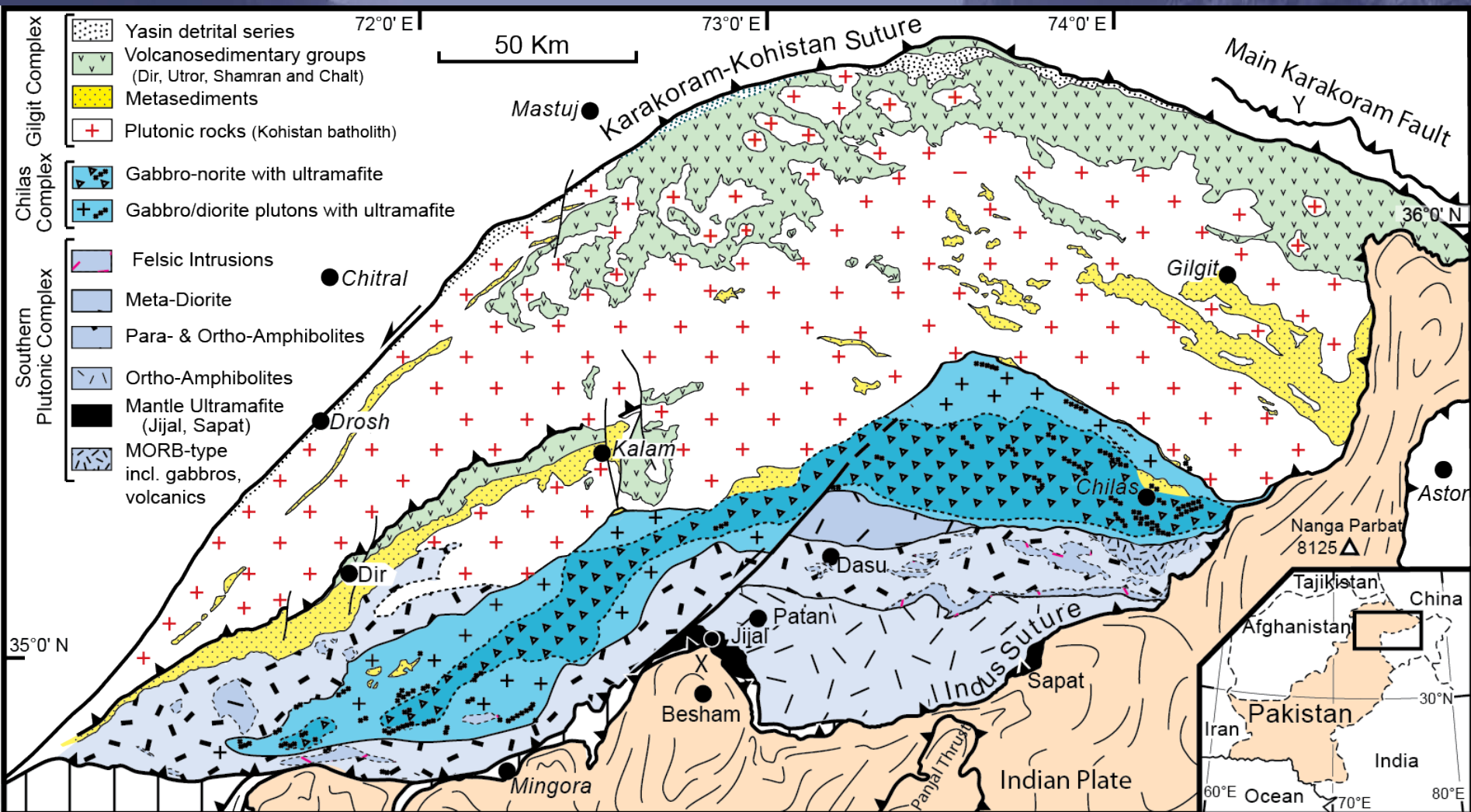
Jagoutz and Behn,
under review

Conclusions

- Foundering of the density unstable lower arc crust can explain the location and the characteristics of the continental Moho

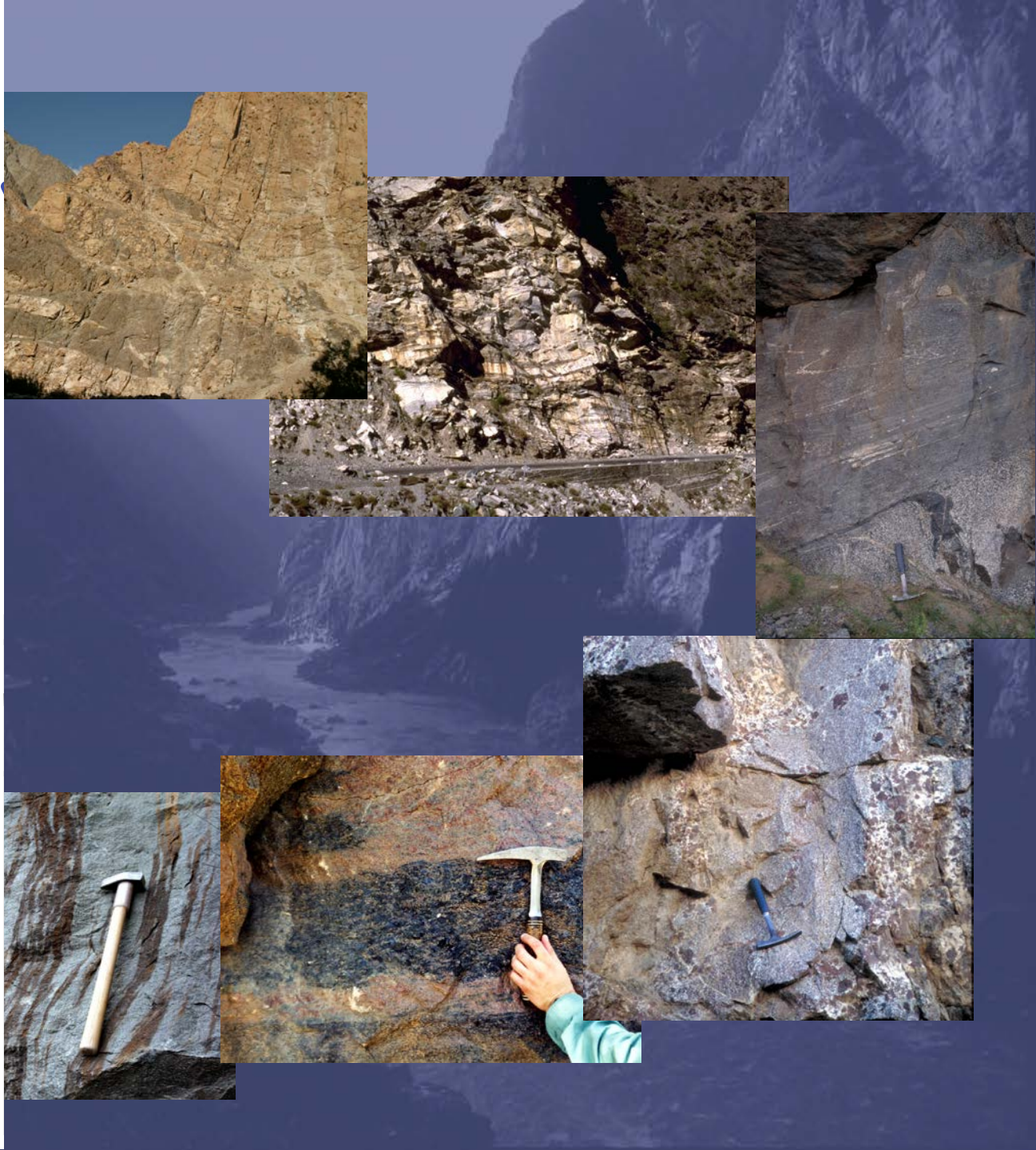
Conclusions

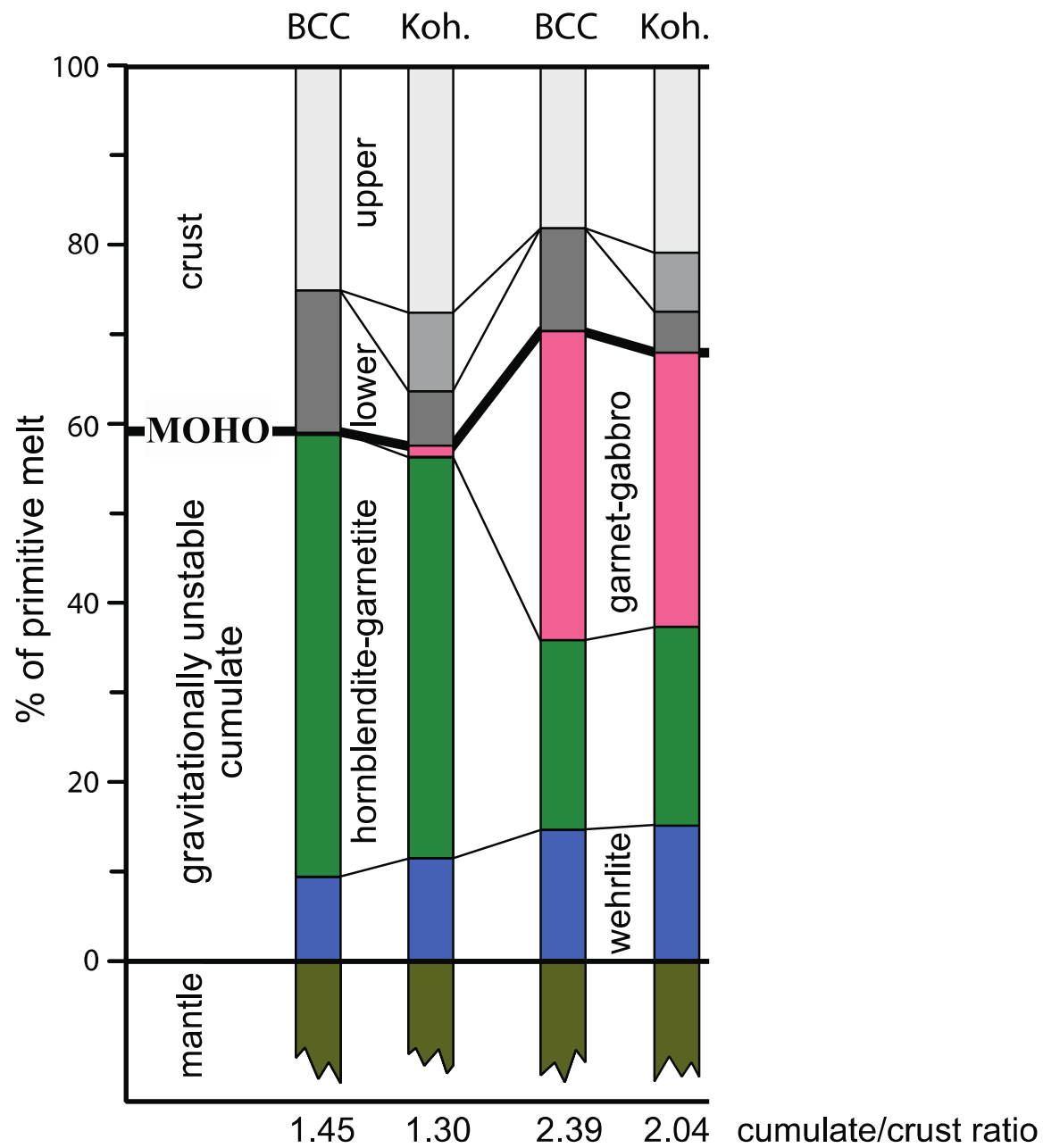
- Continental crust can be formed in oceanic arcs
- Foundering of density unstable lower crust is an essential mechanism :
 - The composition of the CC
 - The location and characteristics of the CC Moho
 - Influence on the global ‘mantle zoo’



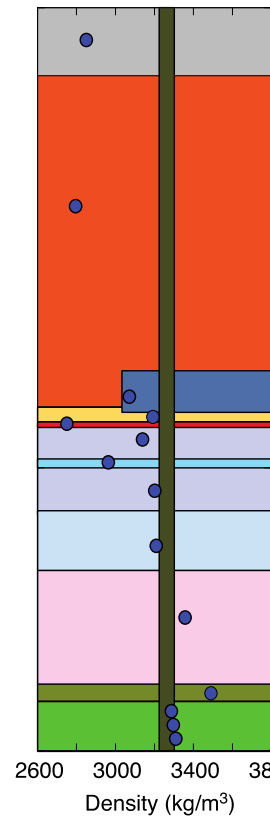
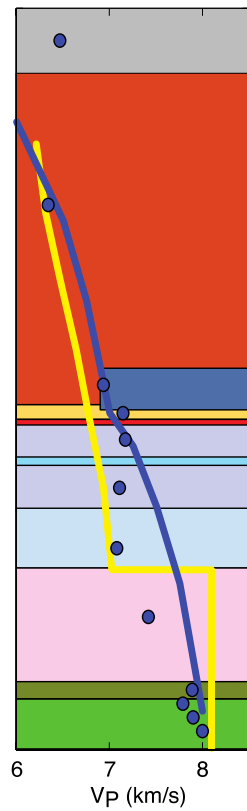
Jagoutz et al., 2011, EPSL

A collage of seven photographs showing various geological features. The top left shows a steep, light-colored rock face. The top center shows a road winding through a rocky, mountainous landscape. The top right is a close-up of layered rock. The middle left shows a wide river valley. The bottom left, center, and right are close-ups of rock samples with hammers for scale, showing different textures and colors like grey, brown, and dark blue.

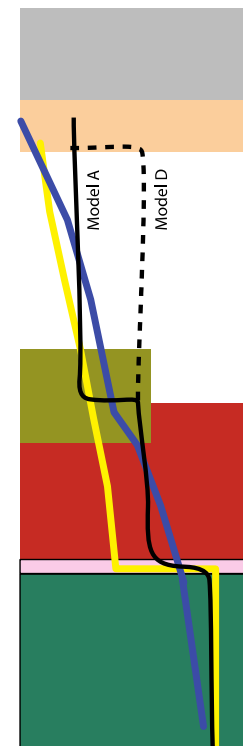
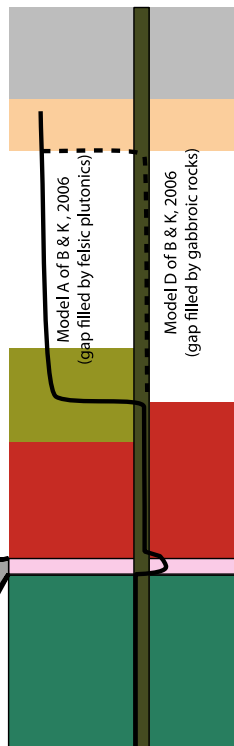




Depth (km)



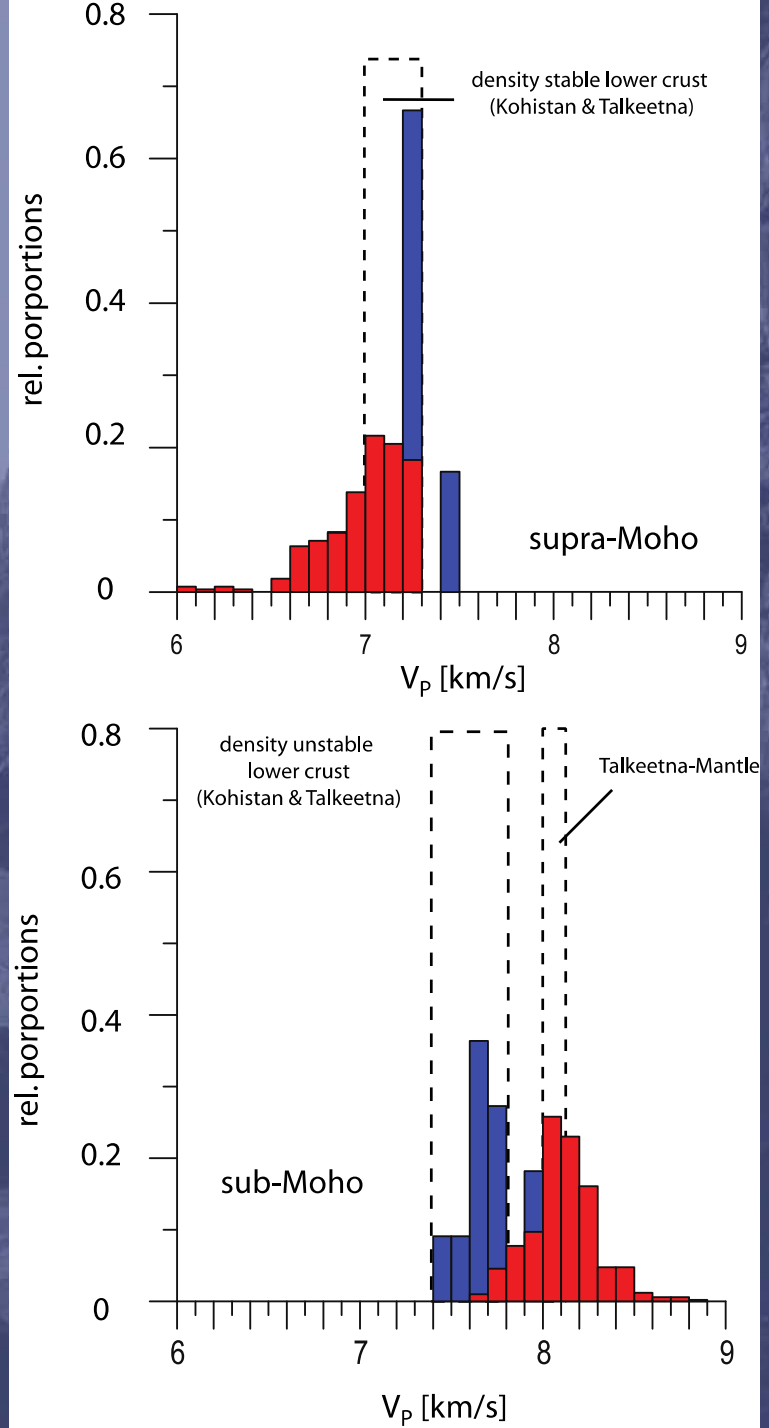
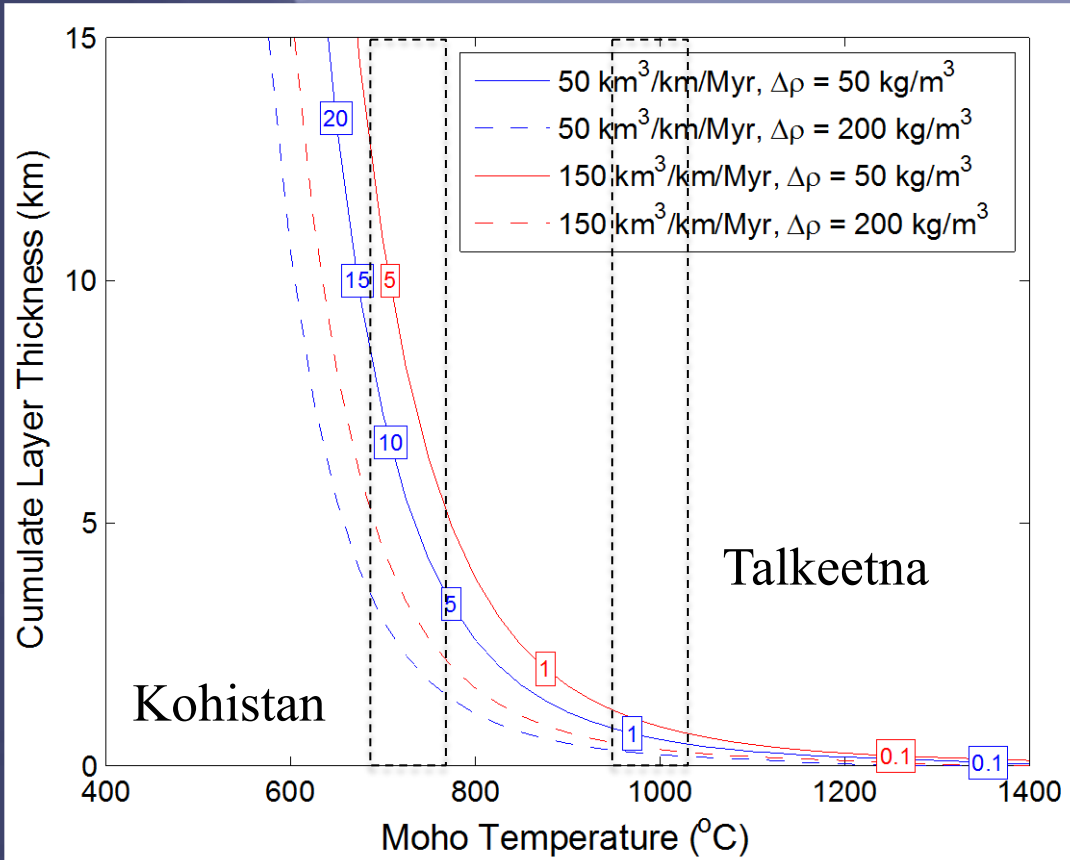
Density
unstable



eechina
qtzdiorite & tonalite,
both ± garnet

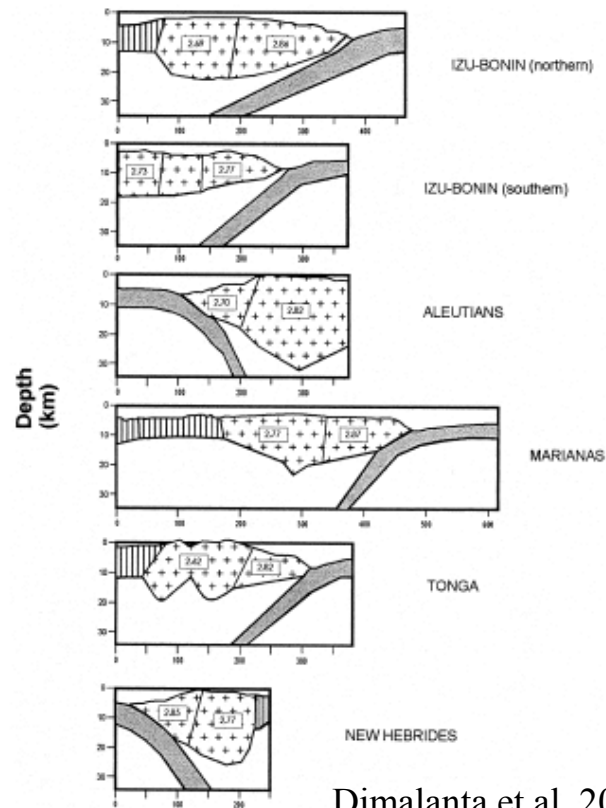
Gap

Jagoutz and Behn, in prep



Jagoutz and Behn,
in prep

The problem



If $C_{\text{arc}} = C_{\text{parcmelt}}$

\Rightarrow Arc production rate = Arc magma flux

Else

Arc production rate =
Arc magma flux - delamination flux
and

$$yC_{\text{arc}} = xC_{\text{parcmelt}} - (1-x)C_{\text{delaminat}}$$

We need:

C_{arc} , $C_{\text{delaminat}}$, C_{parcmelt} , x

Arc production rate
 $\sim 1.5-7 \text{ km}^3/\text{a}$

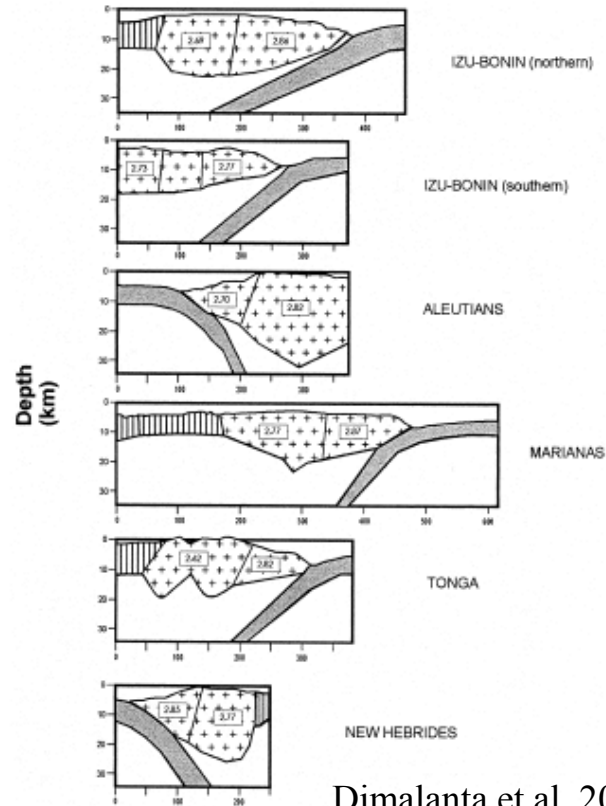
Take home messages

- A 'hidden' mass flux at the mantle-crust interface in arc exist that has a similar magnitude as the subduction of oceanic crust
- Foundering of the density unstable lower arc crust can explain the location and the characteristics of the continental Moho

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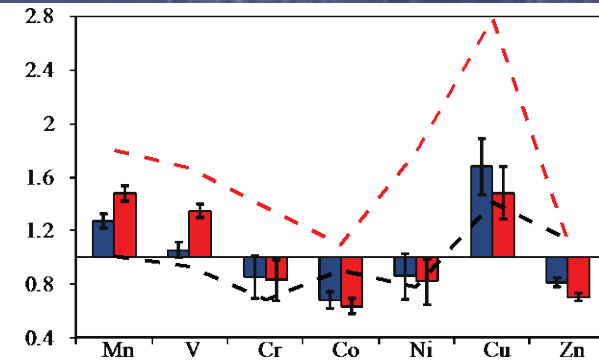
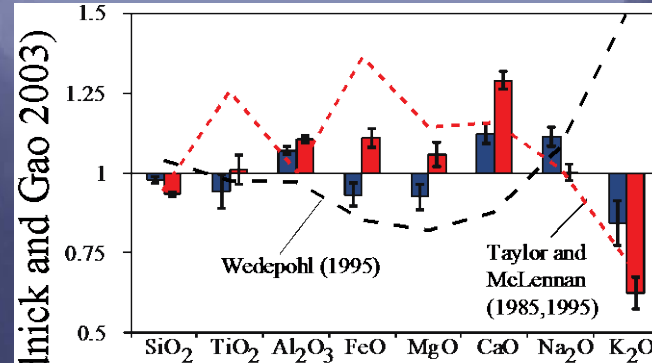
Arc production rate
 $\sim 1.5-7 \text{ km}^3/\text{a}$

Step 1: modern mature island arc crust (= the Kohistan) is similar to continental crust

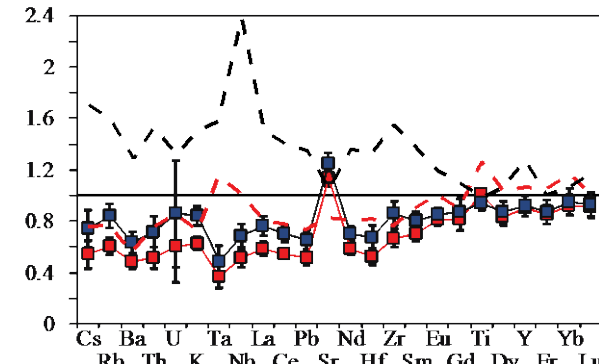
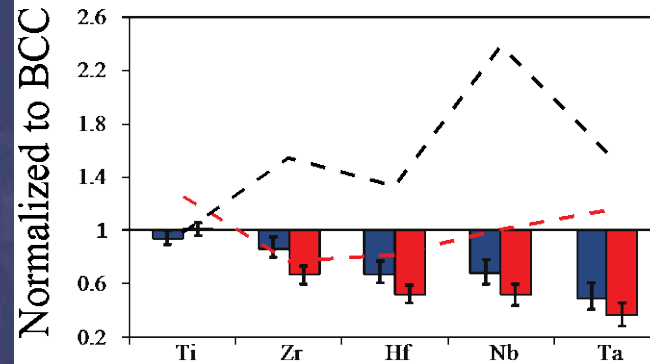
	bulk crust	
	Kohistan arc	continental
	Koh-m#1	RG
SiO ₂ (wt%)	59.3	60.6
TiO ₂	0.68	0.72
Al ₂ O ₃	17.0	15.9
FeO ^T	6.26	6.71
MnO	0.13	0.10
MgO	4.32	4.66
CaO	7.21	6.41
Na ₂ O	3.42	3.07
K ₂ O	1.52	1.81
P ₂ O ₅	0.19	0.13
X _{Mg} ^b	0.552	0.553

Rudnick+Gao (2003)

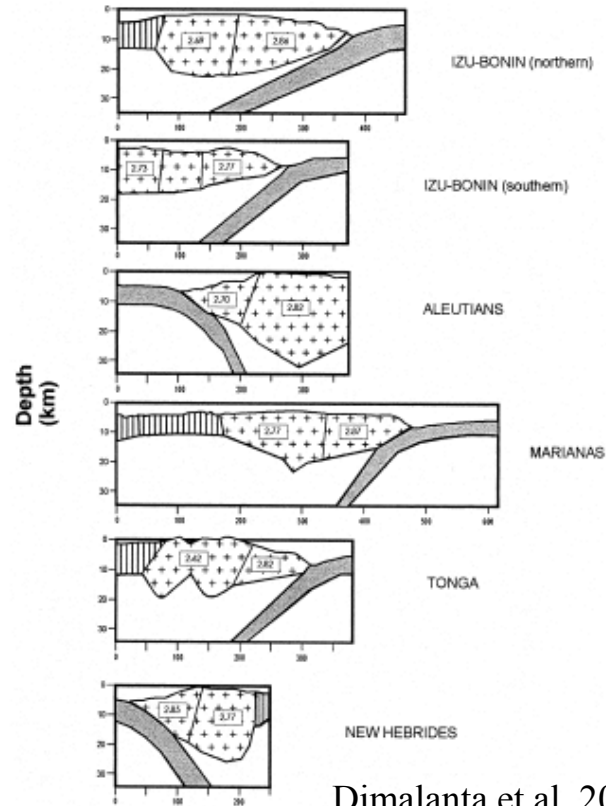
Jagoutz+Schmidt (2012)



Model #1 mapped thickness
Model #2 barometric thickness



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We need:

$$C_{\text{arc}}, C_{\text{delaminat}}, C_{\text{parcmelt}}, X$$

Arc production rate
 $\sim 1.5-7 \text{ km}^3/\text{a}$

Step 2: primitive (basaltic calc-alkaline/tholeiitic) mantle melts in arcs

Compilation of primitive volcanics from GEOROC data base (Mainz)

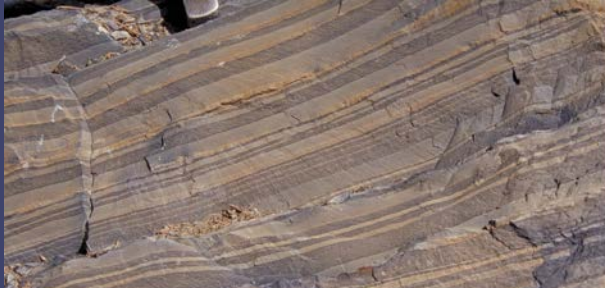
	calc-alkaline / tholeiitic									
	Aleutians	Bismark	Kuriles	Marianas	Tonga	Vanuatu	Yap	Palau	Kohistan	average
<i>n</i>	12	9	14	20	13	31	4	13	10	
SiO₂ (wt%)	49.8	50.5	51.5	50.2	49.3	49.4	49.5	51.2	50.7	50.2
TiO₂	0.75	0.99	0.83	0.91	0.60	0.69	0.59	0.60	0.80	0.8
Al₂O₃	15.7	14.9	14.7	15.7	12.0	12.9	15.2	16.5	15.6	14.8
FeO^{tot}	8.8	9.0	8.8	8.1	9.3	9.6	8.6	7.9	8.6	8.8
MnO	0.16	0.17	0.16	0.13	0.18	0.19	0.17	0.16	0.14	0.2
MgO	11.1	10.9	10.9	10.9	13.4	12.7	11.1	10.7	10.9	11.4
CaO	10.5	11.1	9.9	11.3	10.5	11.1	12.9	9.9	9.9	10.8
Na₂O	2.4	2.2	2.3	2.3	2.1	2.0	1.7	2.6	2.4	2.2
K₂O	0.70	0.31	0.83	0.37	2.12	0.98	0.18	0.28	0.61	0.7
P₂O₅	0.13	0.11	0.14	0.12	0.46	0.20	0.05	0.07	0.19	0.2
X_{Mg}^b	0.691	0.684	0.687	0.705	0.718	0.702	0.697	0.707	0.693	0.7
Cr (ppm)	611	545	647	456	697	666	507	689	512	592
Ni	208	213	196	240	300	242	153	207	167	214

X_{Mg} = 0.65-0.74

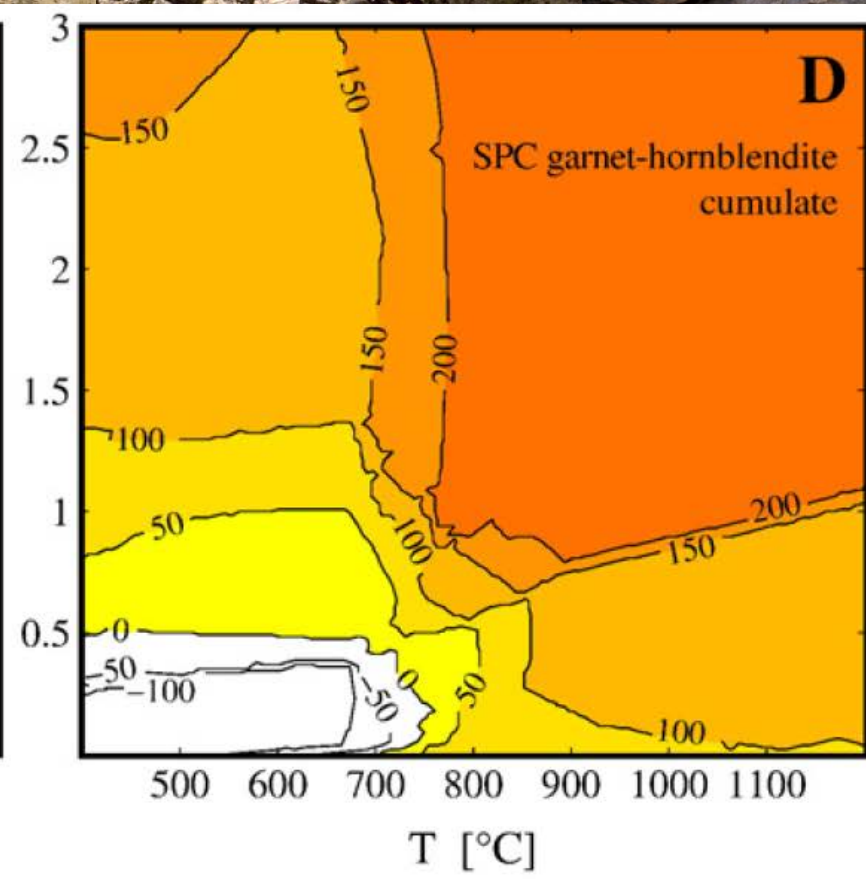
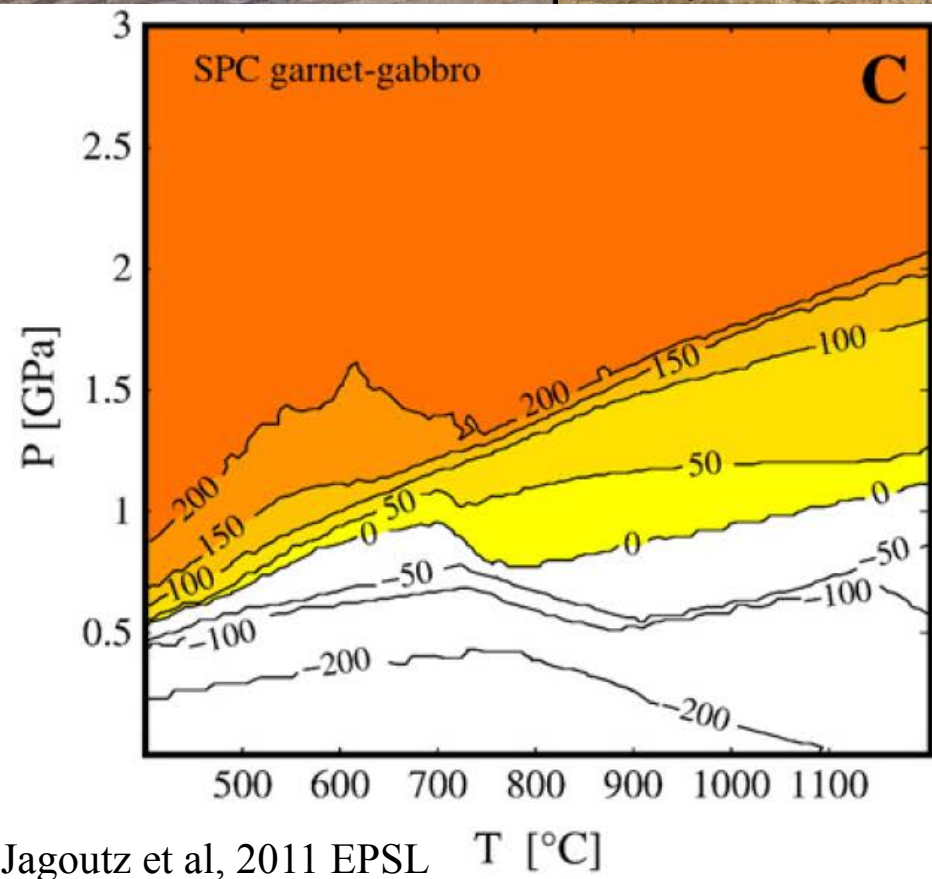
Ni = 150-500

Cr = 300-1100

**Tonga, Marianas and Aleutians also have
primitive high-Mg andesites**

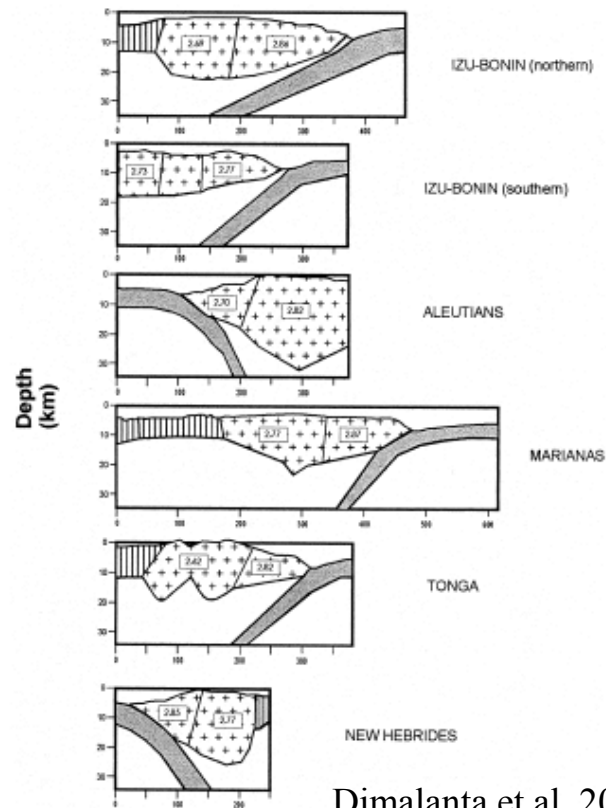


$\Delta\rho$ cumulate – primitive mantle (kg/m^3)



Jagoutz et al, 2011 EPSL

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We need:

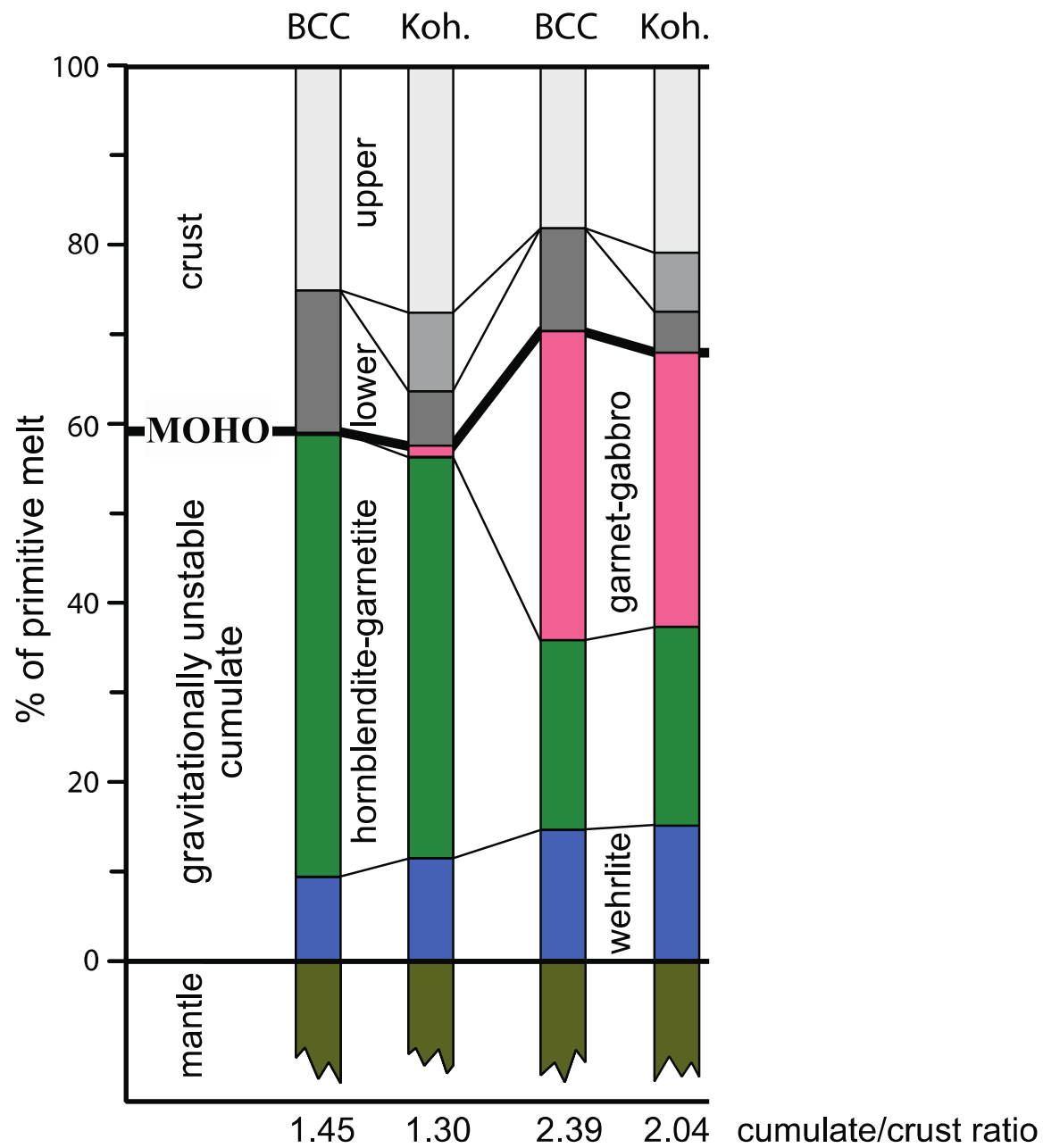
$$C_{\text{arc}}, C_{\text{delaminat}}, C_{\text{parcmelt}}, x$$

Arc production rate
 $\sim 1.5-7 \text{ km}^3/\text{a}$

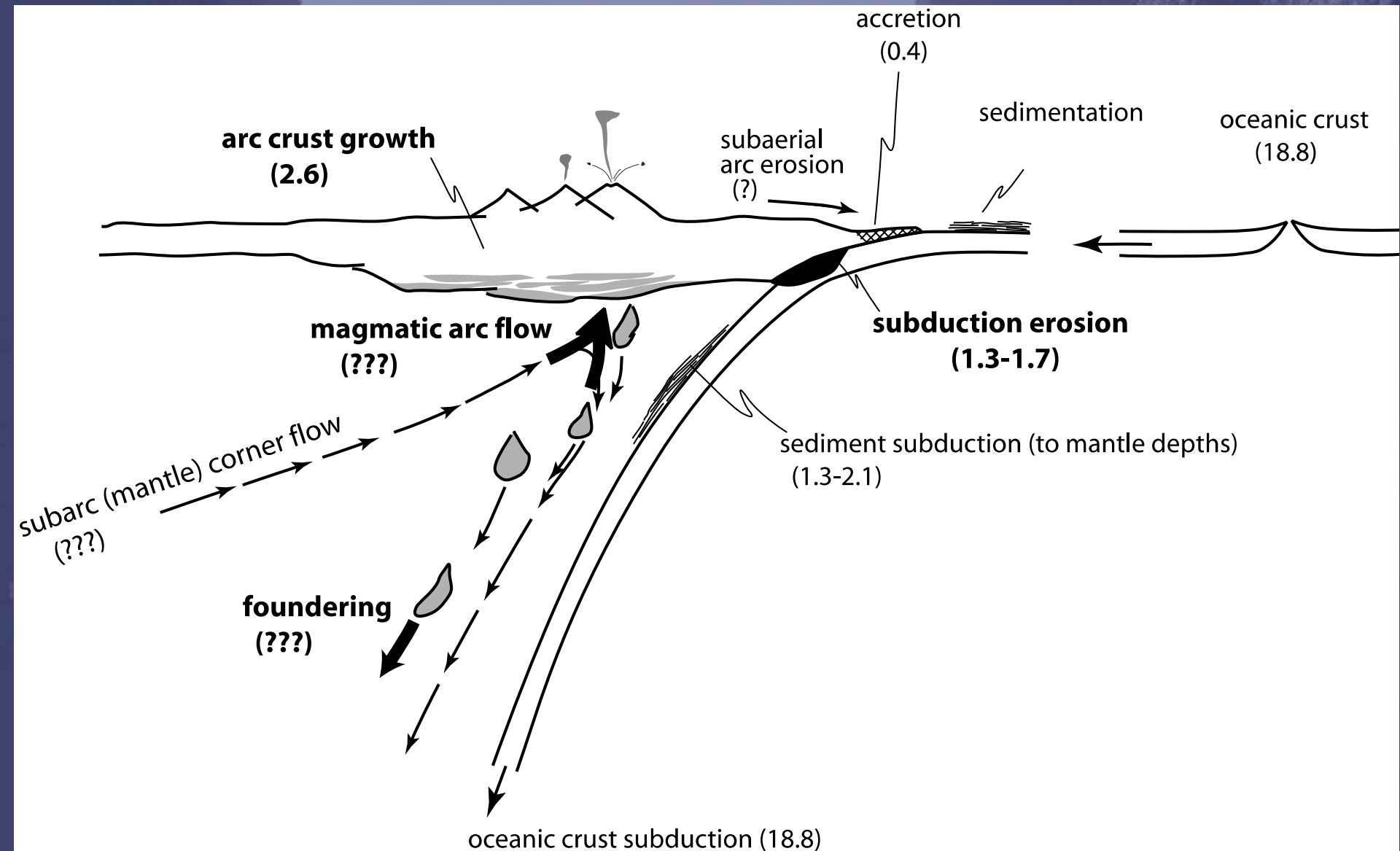
Simple global mass balance



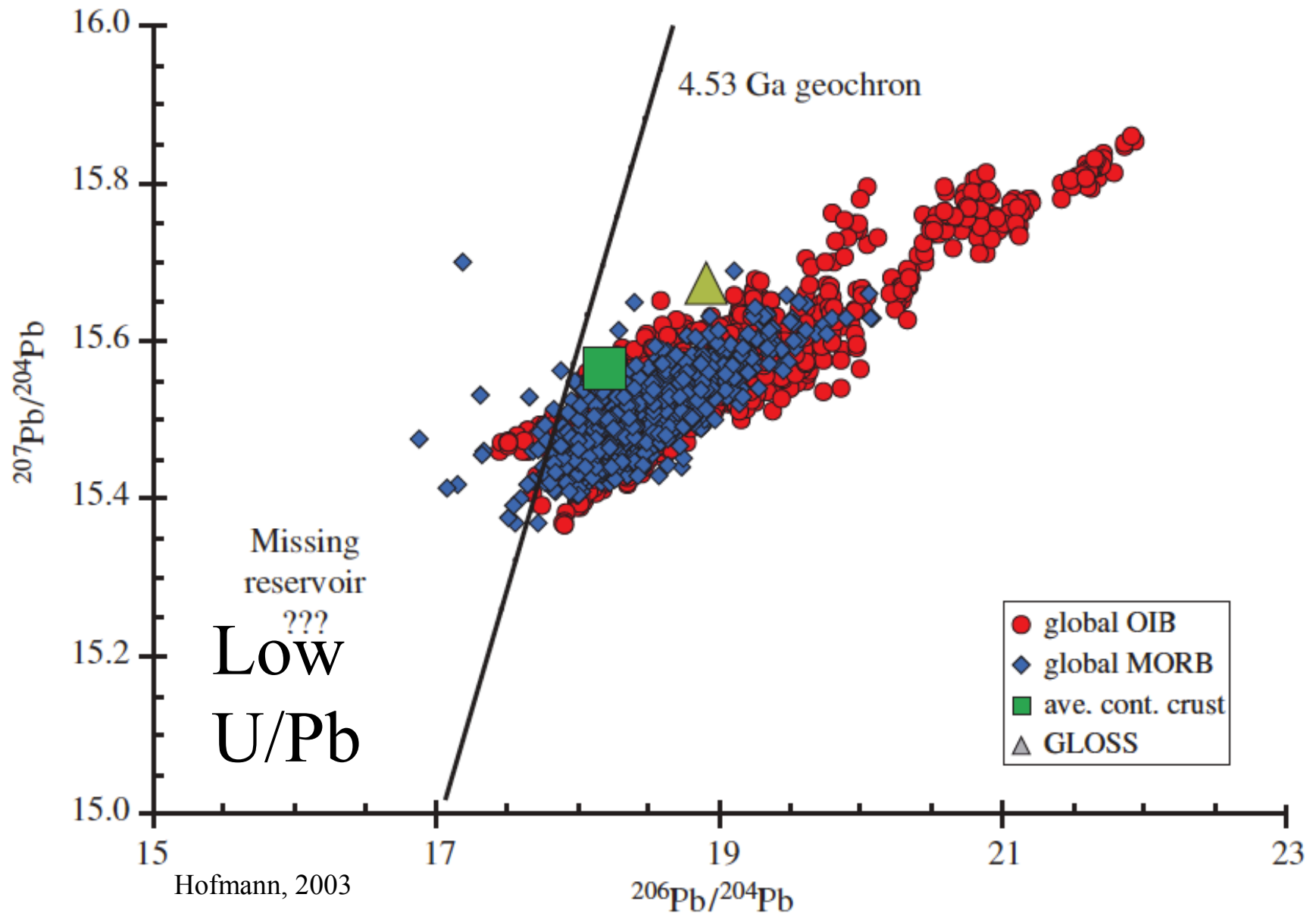
	continental crust		primitive arc melts			
	Kohistan	Rudnick+Gao	Kuriles calc-alkaline/tholeiitic	Kohistan	Aleutians alkaline	Mariannas boninite
Major elements [wt%]						
SiO ₂	56.94	60.60	51.50	50.72	45.41	56.68
TiO ₂	0.73	0.72	0.82	0.80	2.42	0.34
Cr ₂ O ₃	0.00	0.00	0.10	0.06	0.06	0.09
Al ₂ O ₃	17.17	15.90	14.74	15.67	13.17	14.08
DFeO	7.58	6.71	8.82	8.60	11.06	7.63
MnO	0.15	0.10	0.16	0.14	0.18	0.15
MgO	4.96	4.66	10.93	10.87	12.62	10.42
CaO	8.02	6.41	9.85	9.88	9.88	7.94
Na ₂ O	3.09	3.07	2.25	2.43	3.49	2.00
K ₂ O	1.19	1.81	0.78	0.61	1.30	0.67
P ₂ O	0.17	0.13	0.14	0.19	0.49	0.06
Total	100.0					
Mg#	0.54	0.55	0.69	0.69	0.67	0.71
	Kohistan					
		wherlite	15.5	17.7	-4.1	12.5
		hblgrtt	14.5	7.0	215	74.0
		gabbro	28.9	45.4	239	-59.1
		Koh.crust	41.0	29.9	129	72.6
		r ²	0.84	2.6	6.5	50
	RG					
		wherlite	16.5	19.3	-3.4	13.0
		hblgrtt	1.0	-8.7	190	58.7
		gabbro	59.2	75.2	-166	-16.6
		RG-crust	23.2	14.3	80.3	44.9
		r ²	1.2	3.1	9.2	49

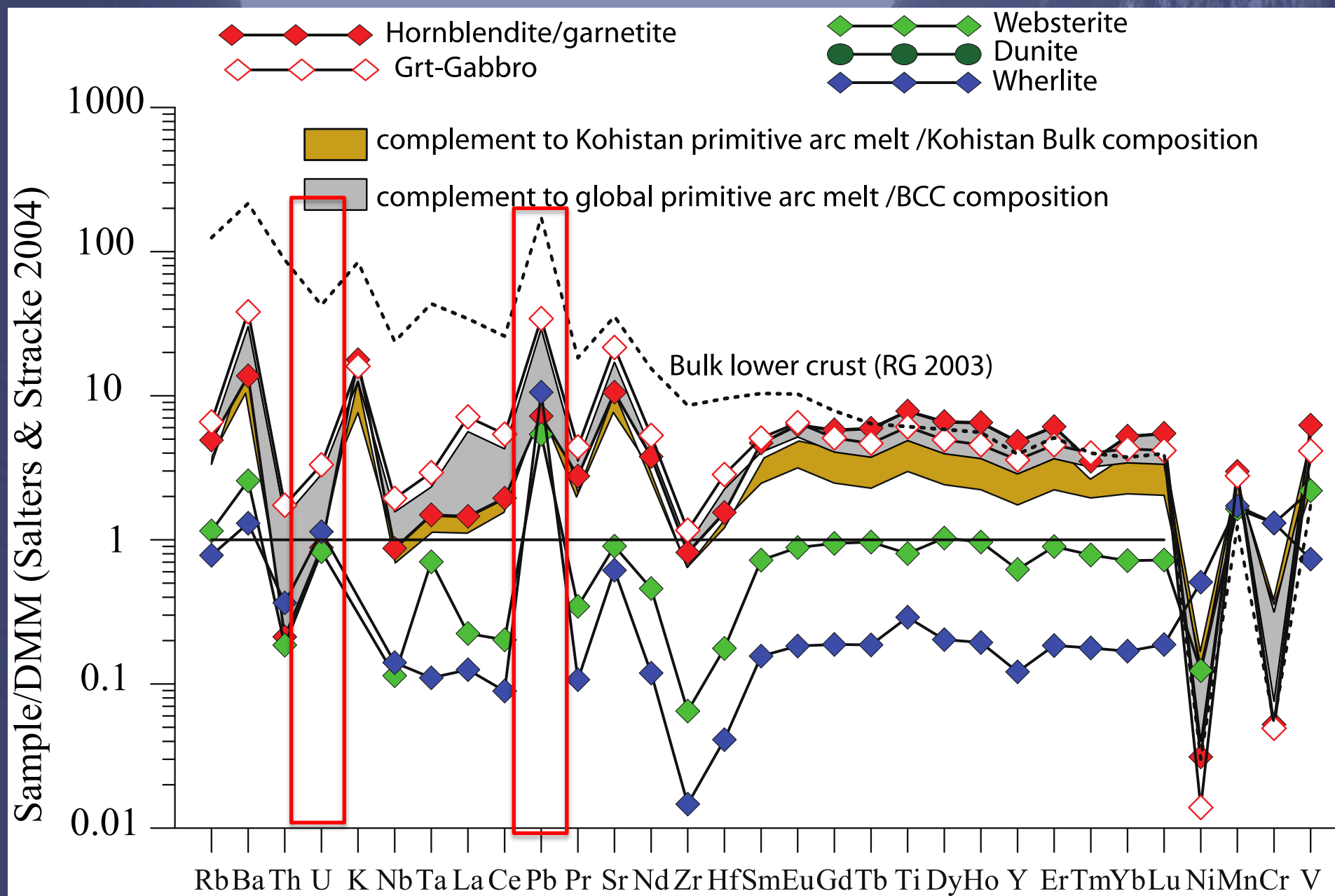


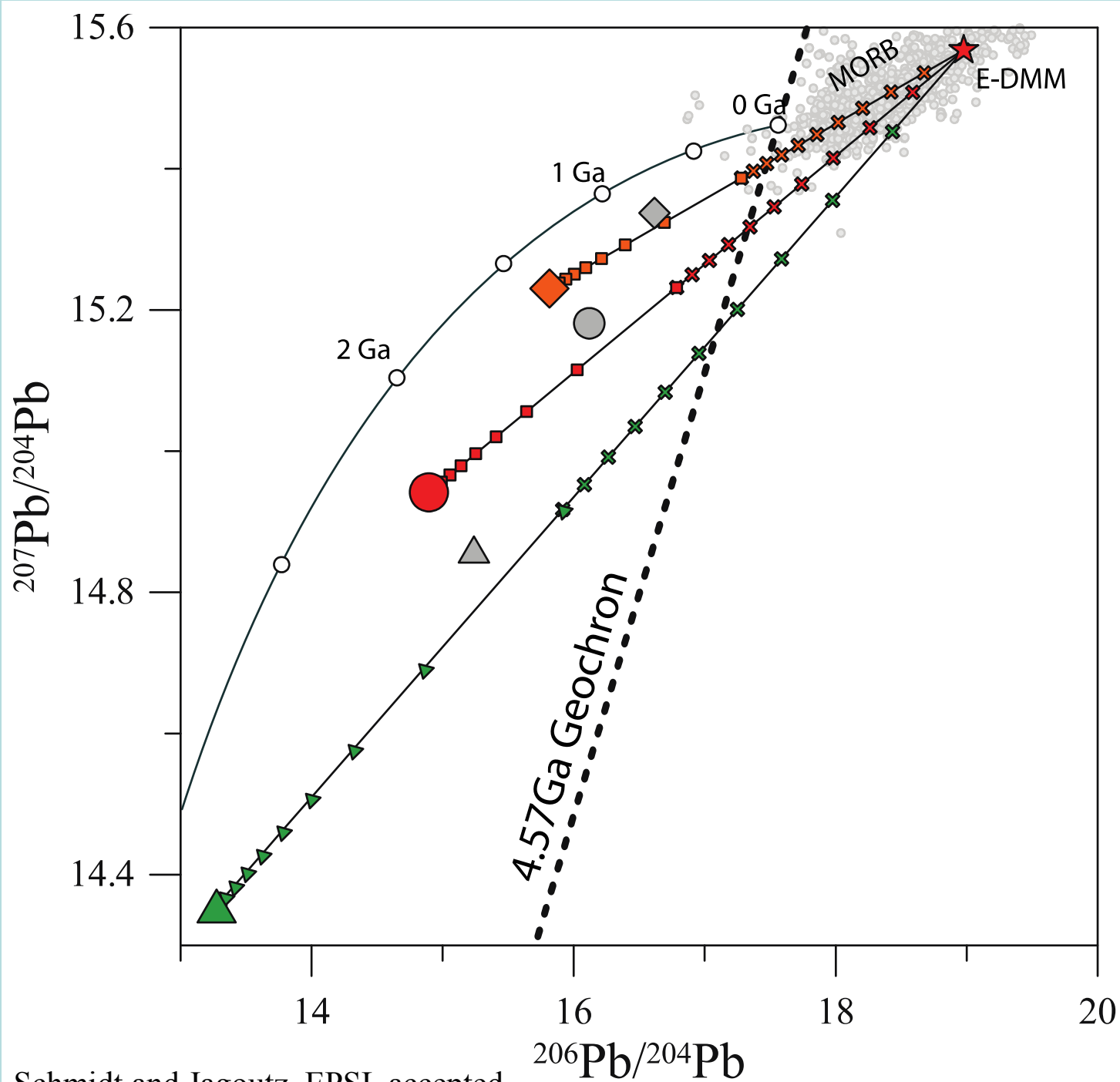
Subduction zone mass balance [km^3/a]



Important reservoir for the Pb paradox ?







Take home messages

- A 'hidden' mass flux at the mantle crust interface in arc exist that has a similar magnitude as the subduction of oceanic crust
- **Foundering of the density unstable lower arc crust can explain the location and the characteristics of the continental Moho**