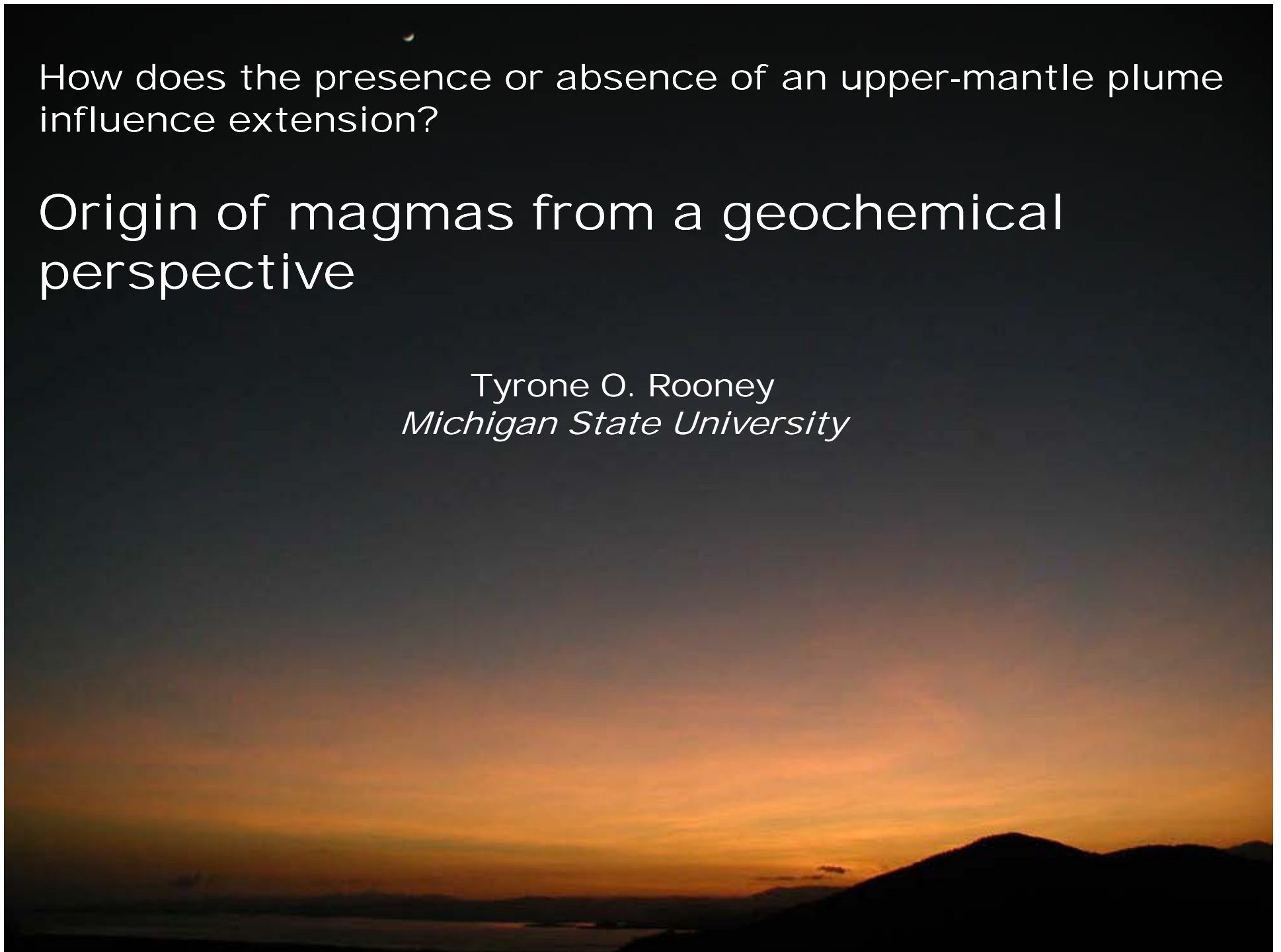


How does the presence or absence of an upper-mantle plume influence extension?

Origin of magmas from a geochemical perspective

Tyrone O. Rooney
Michigan State University



cenozoic volcanism



Mantle plumes and the origin of magmas from a geochemical perspective

why magma ?

Magmas may be necessary to weaken the lithosphere to initiate rifting

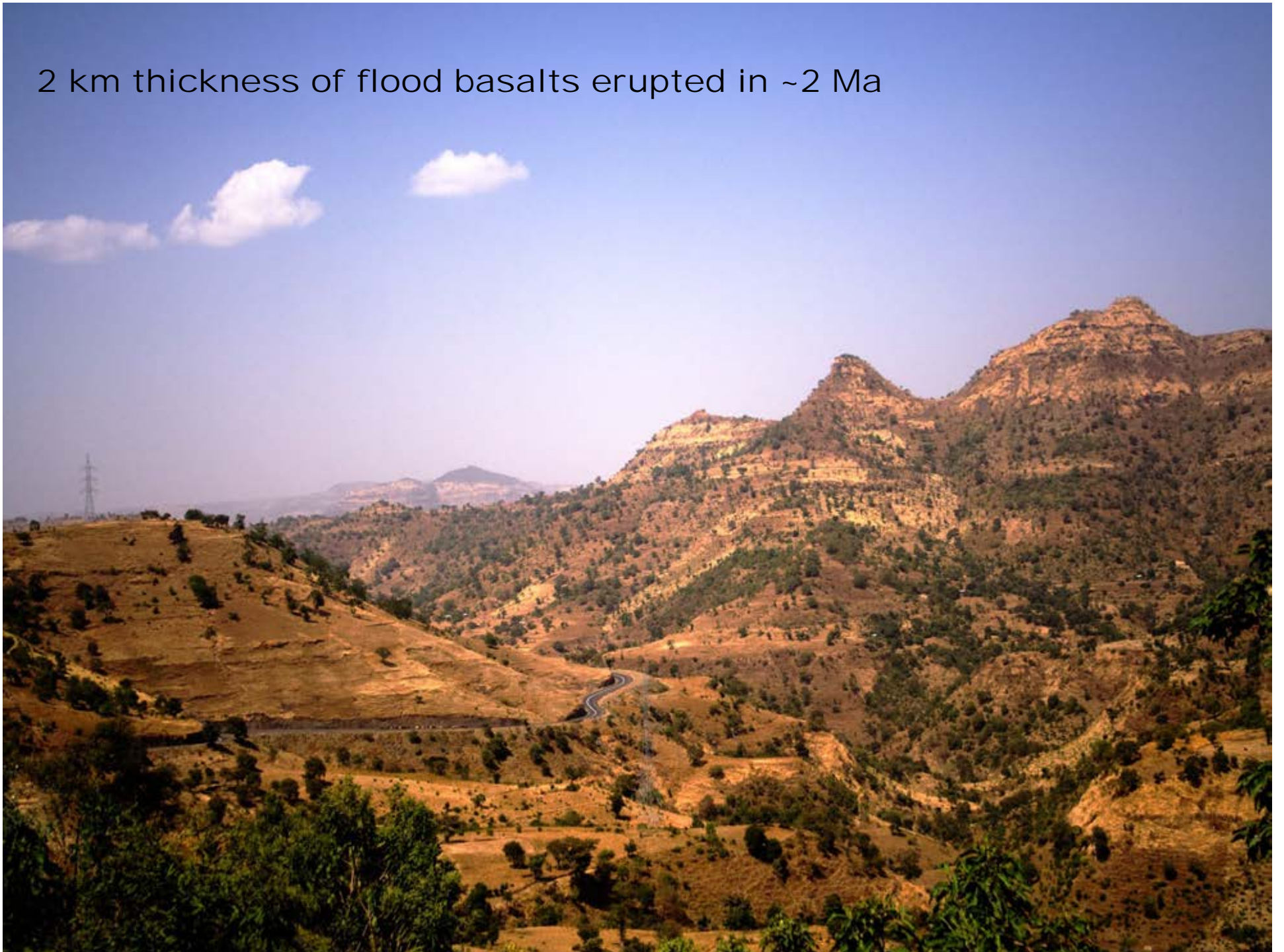
Magmas directly record the thermo-chemical state of the upper mantle during rift evolution

Magmas provide a time window into upper mantle processes

The presence of magma profoundly influences the interpretation of geophysical images of the asthenosphere and lithosphere

Magmas eventually accommodate lithospheric extension

2 km thickness of flood basalts erupted in ~2 Ma

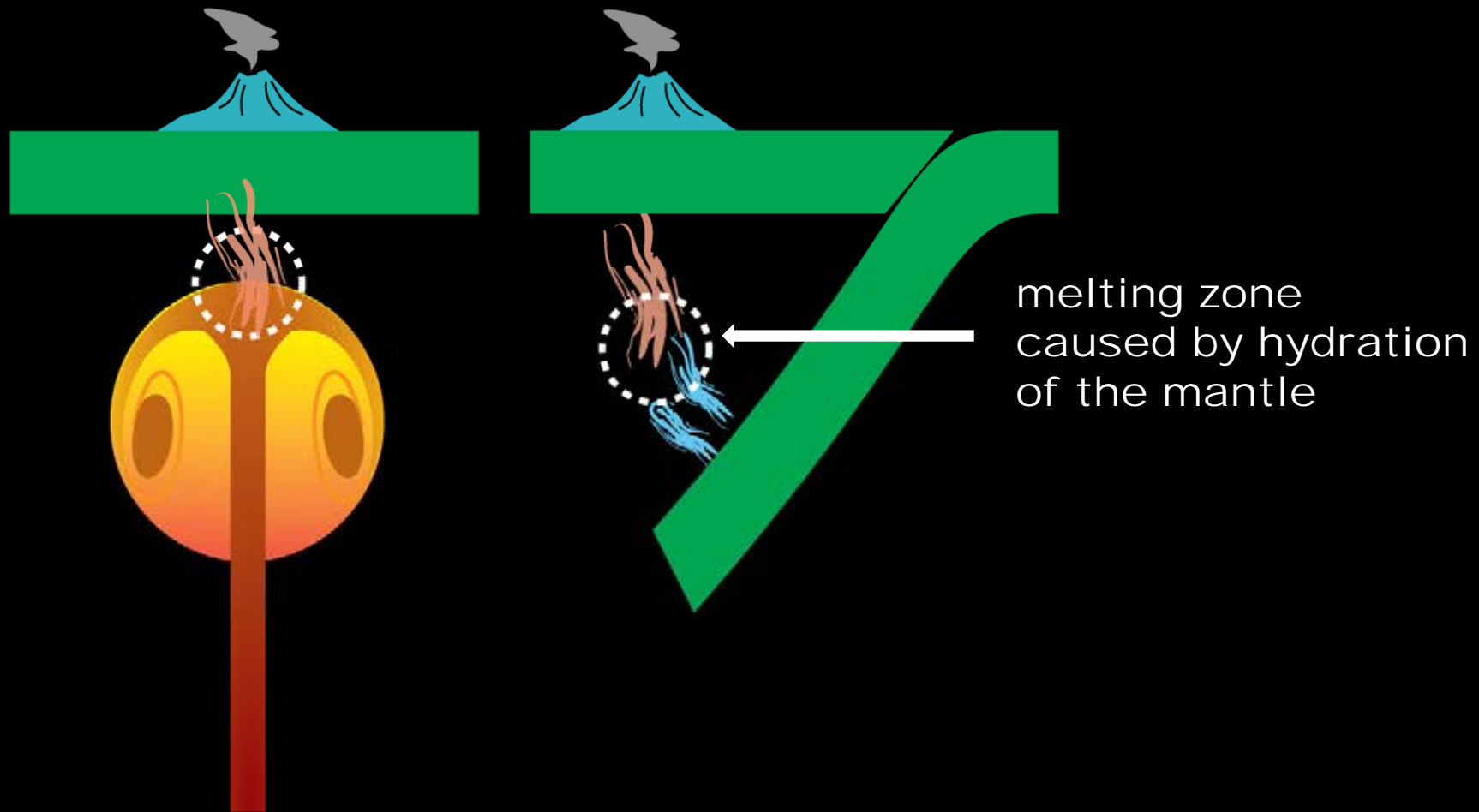


creating melts

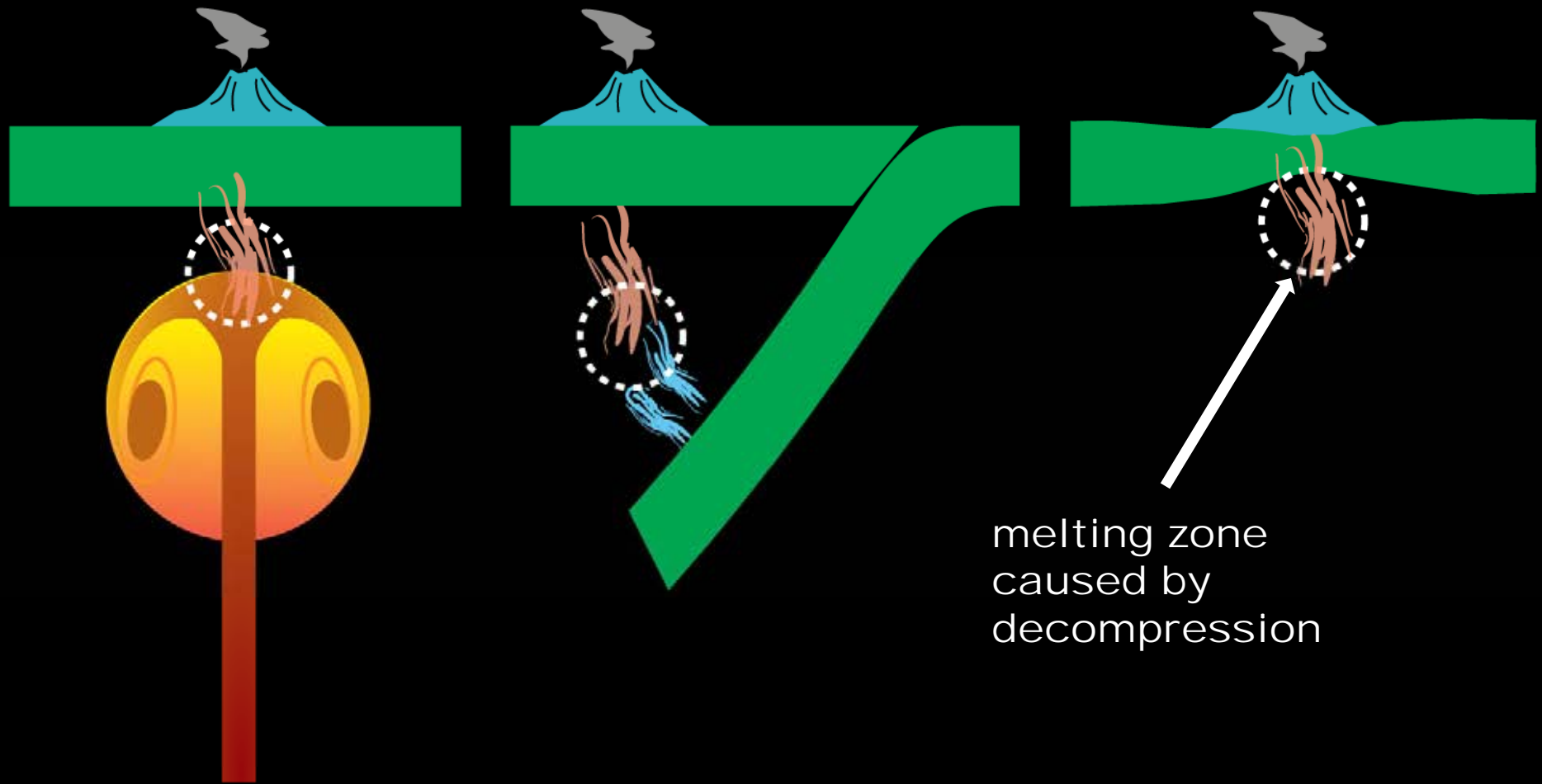


melting zone
caused by increased
mantle temperature

creating melts

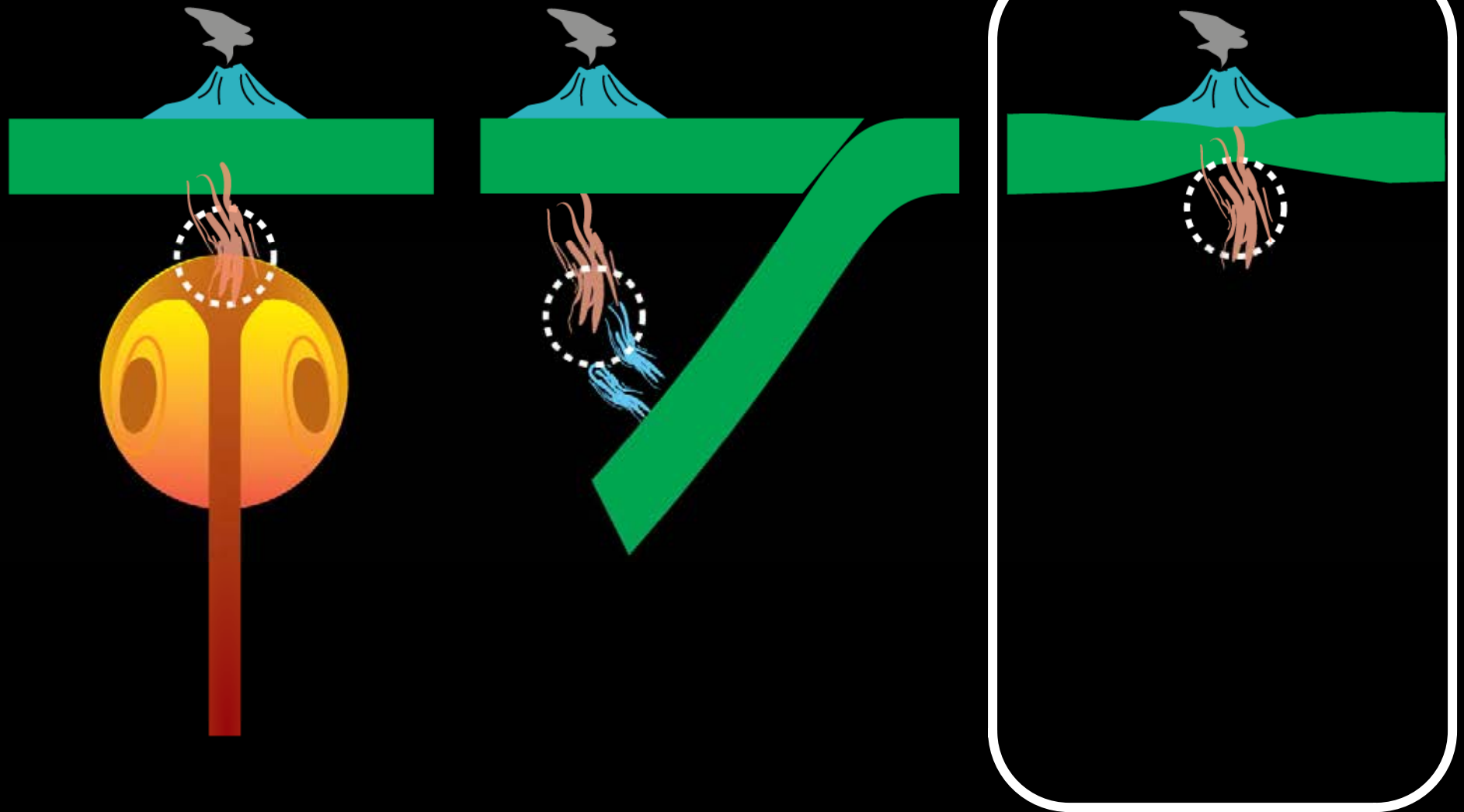


creating melts



melting zone
caused by
decompression

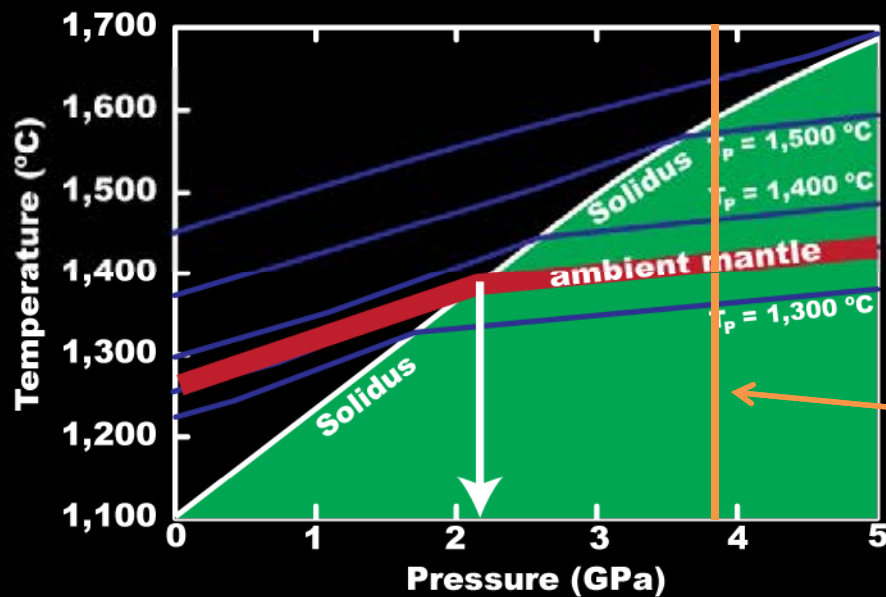
creating melts



Mantle plumes and the origin of magmas from a geochemical perspective

MICHIGAN STATE
UNIVERSITY

creating melts

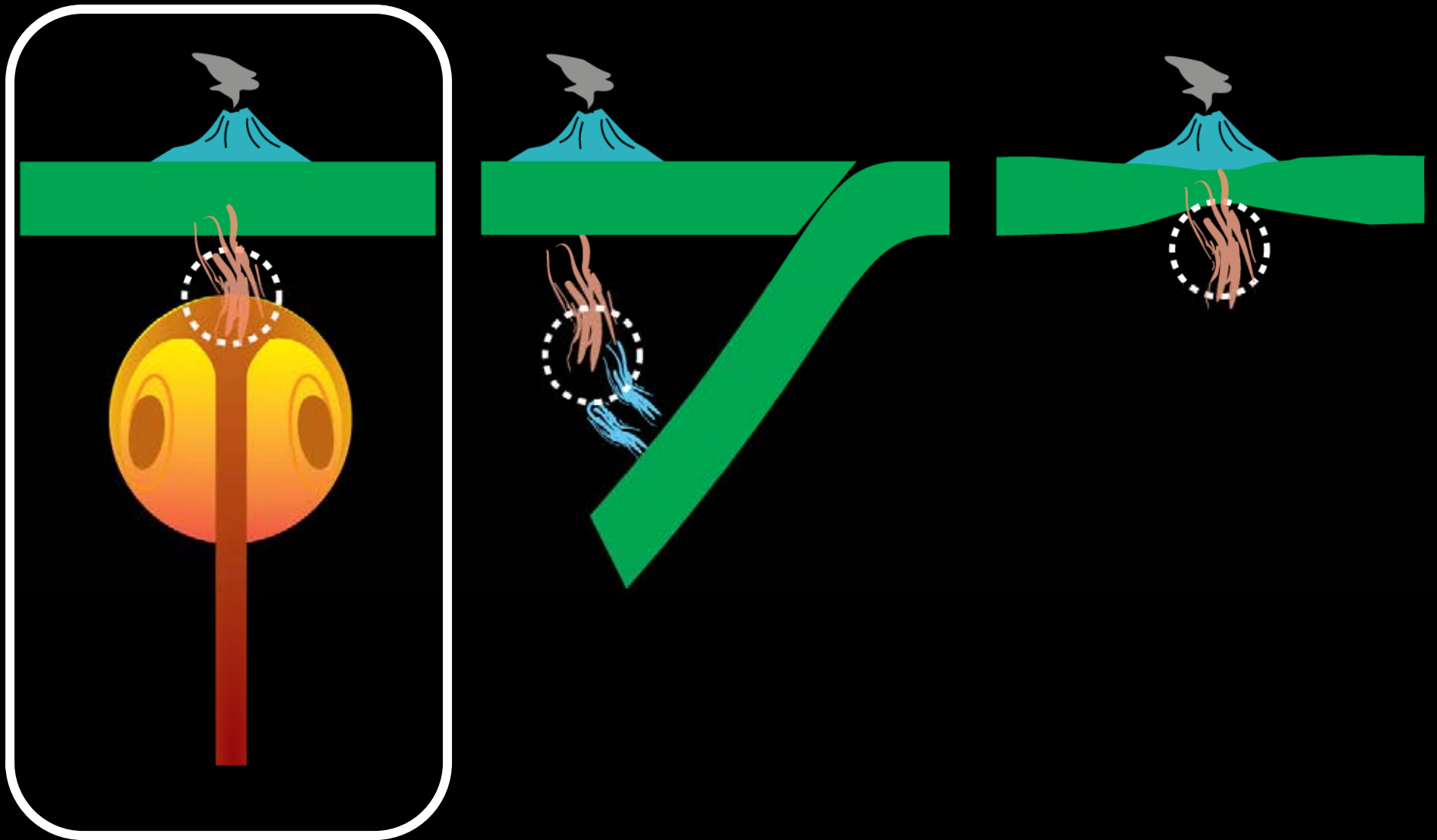


Lithospheric thickness

lithosphere must thin by 50% before melt may be generated

(figure derived from Herzberg & O'Hara, 2002; lithospheric thickness from Dugda et al., 2007)

creating melts

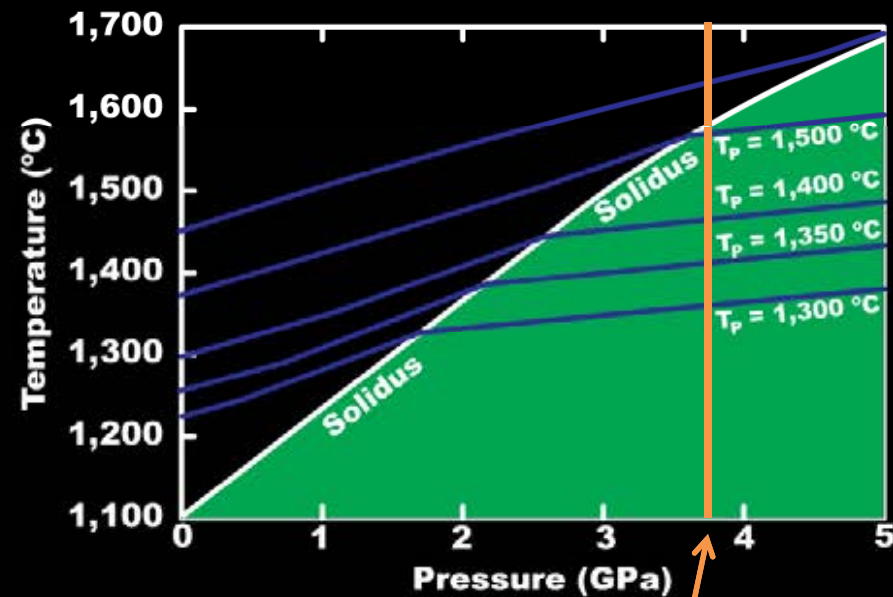


Mantle plumes and the origin of magmas from a geochemical perspective

creating melts



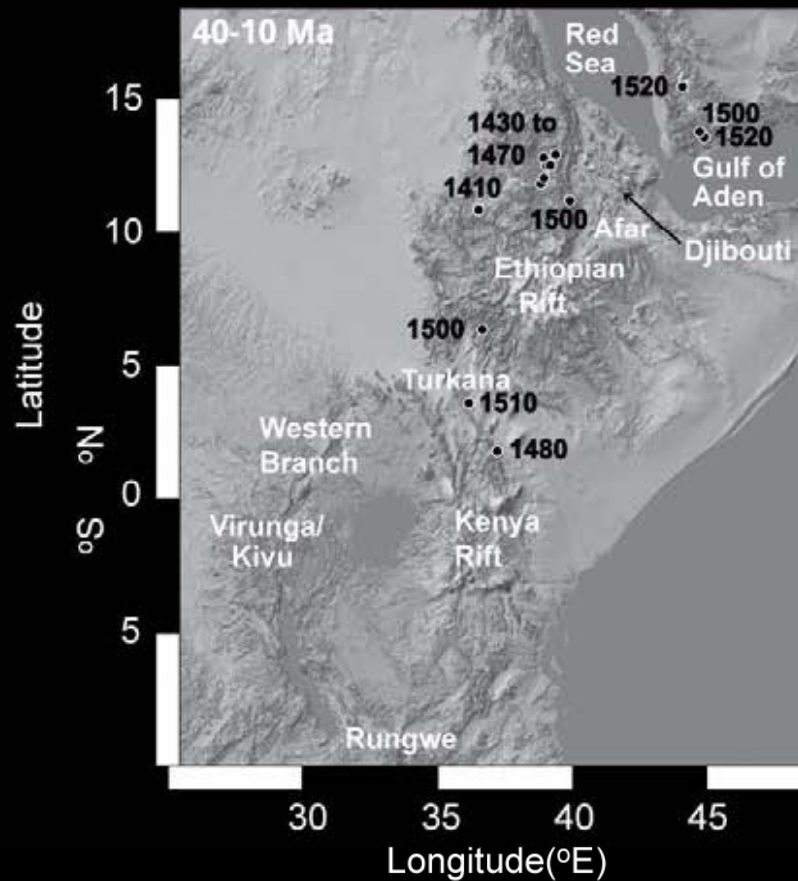
(figure derived from Herzberg & O'Hara, 2002;
lithospheric thickness from Dugda et al., 2007)



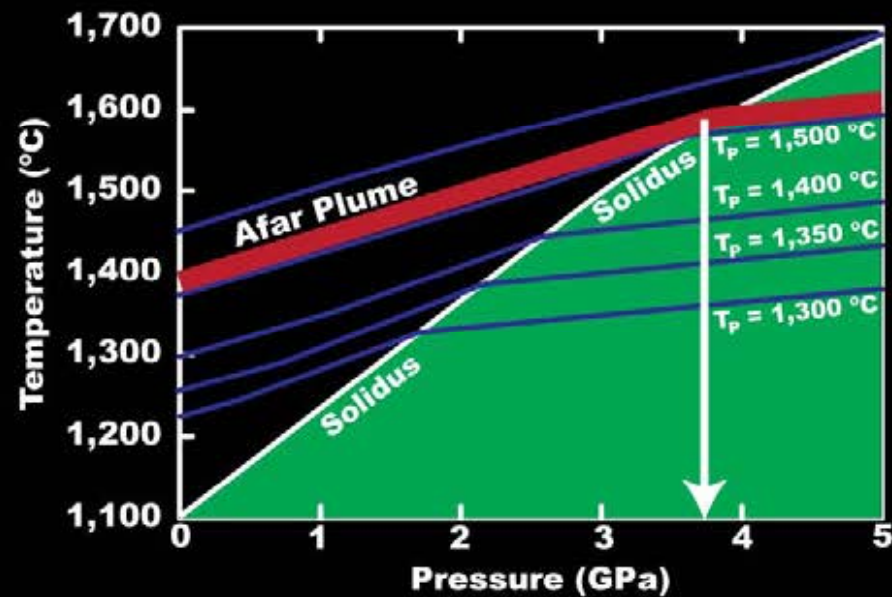
Lithospheric
thickness

creating melts

ambient mantle – 1350°C

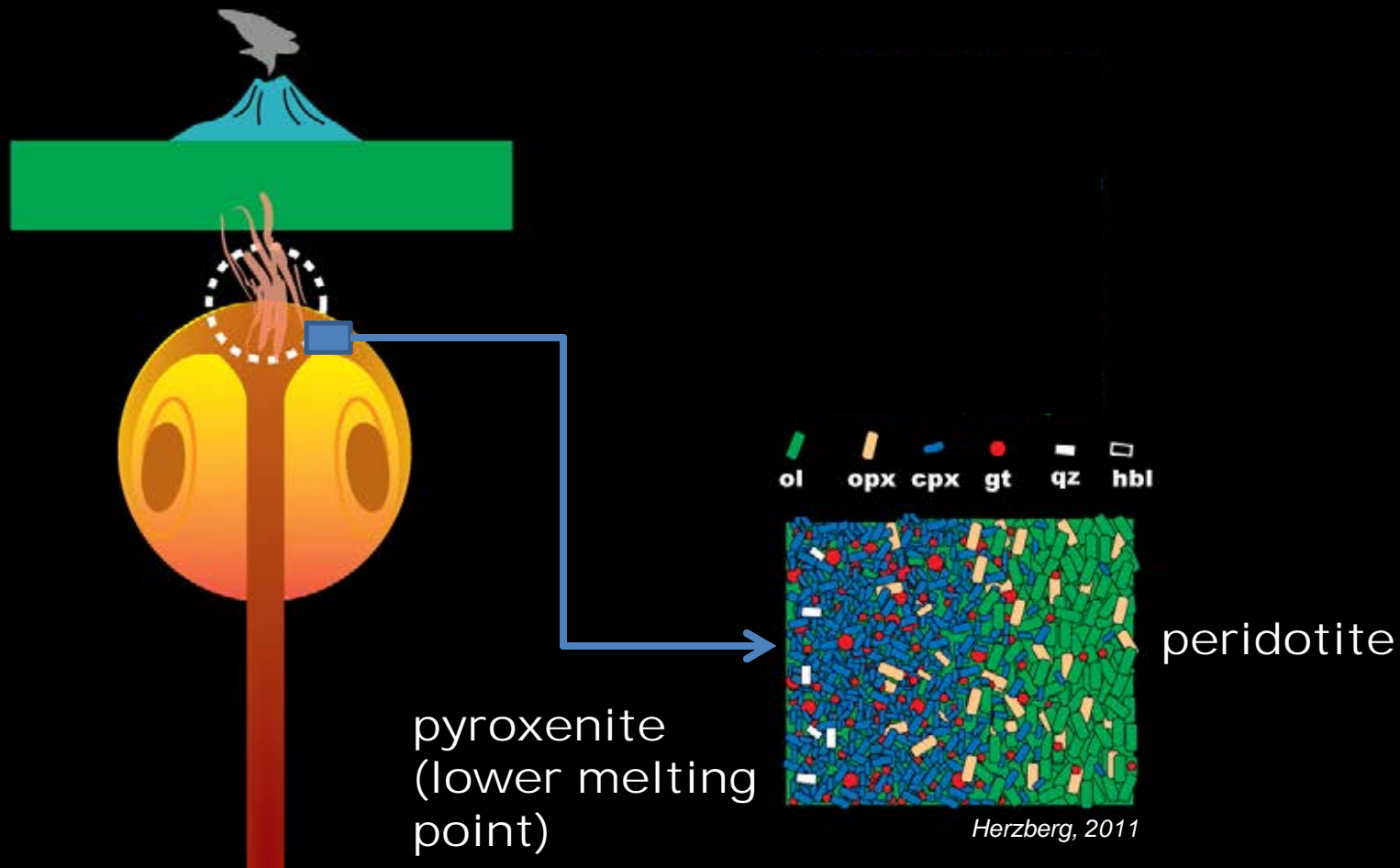


Rooney et al., (2012)



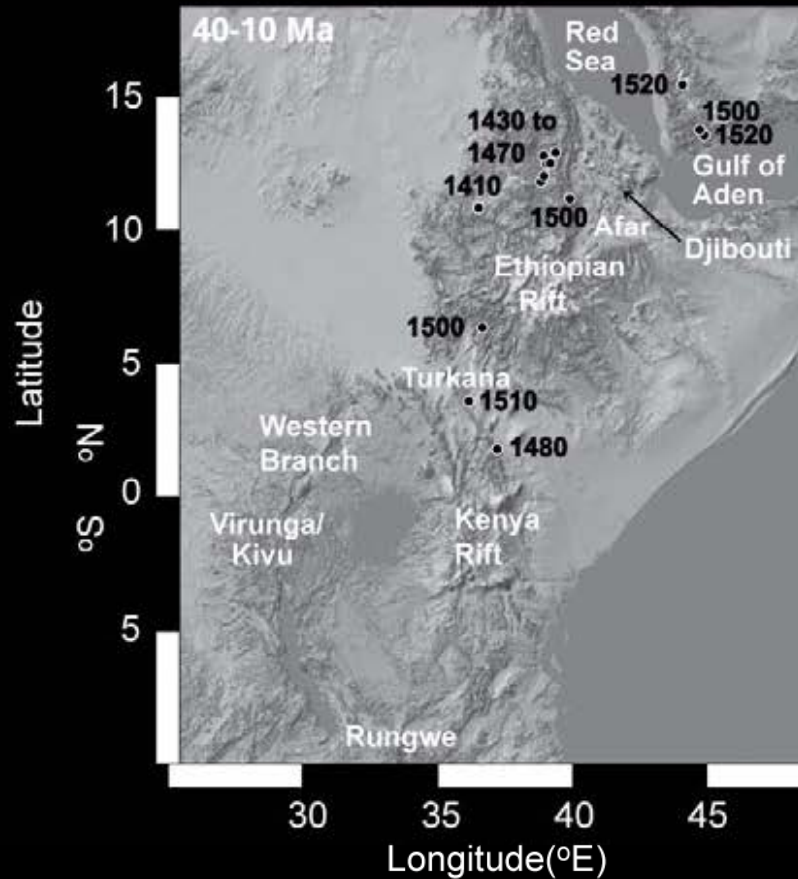
(figure derived from Herzberg & O'Hara, 2002; lithospheric thickness from Dugda et al., 2007; T_p from Rooney et al., 2012)

composition



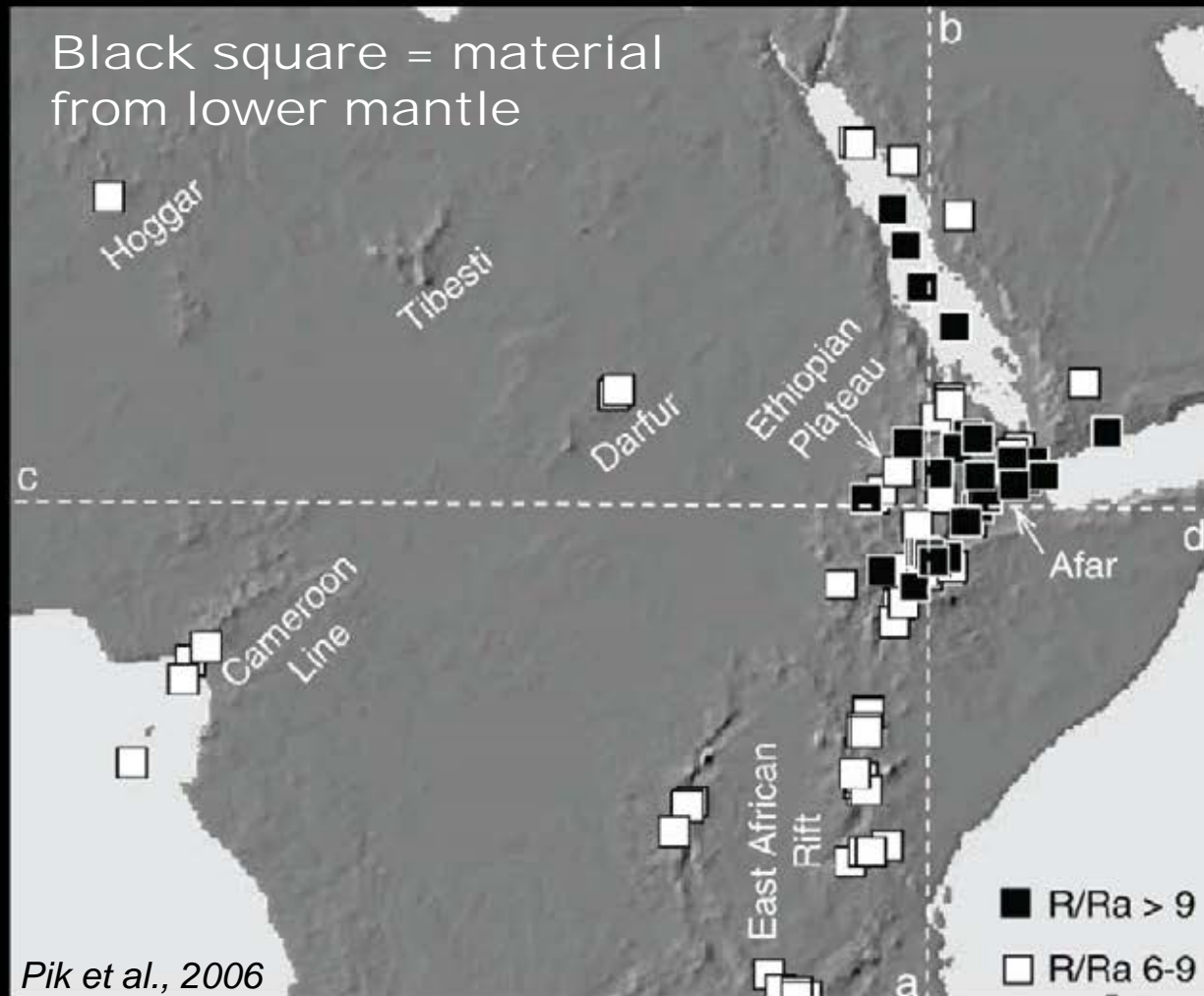
mantle temperature

So what is the role of a plume today?

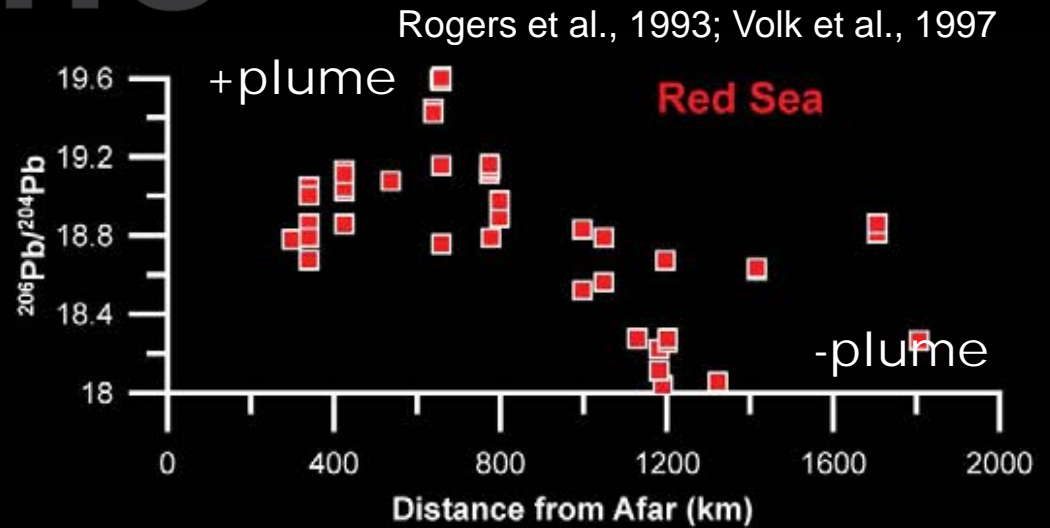


ambient mantle – 1350°C

helium isotopes

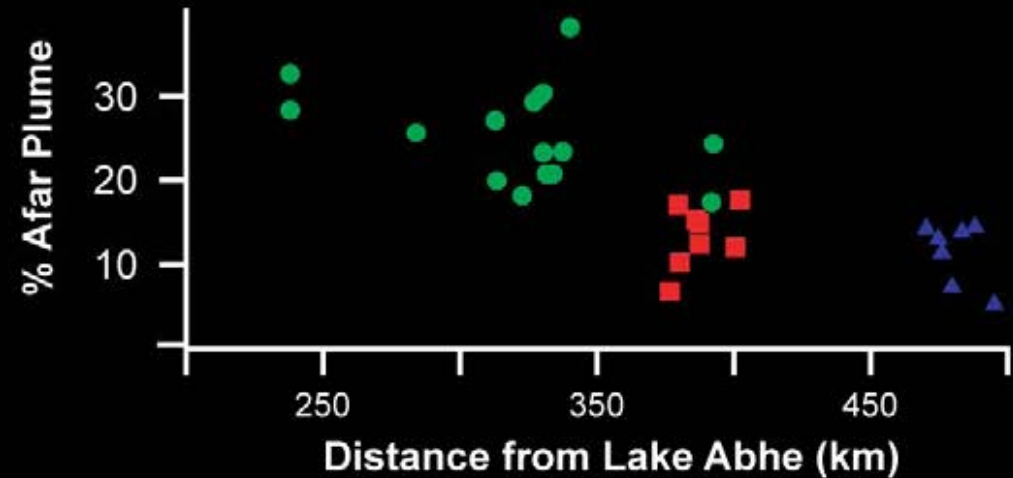
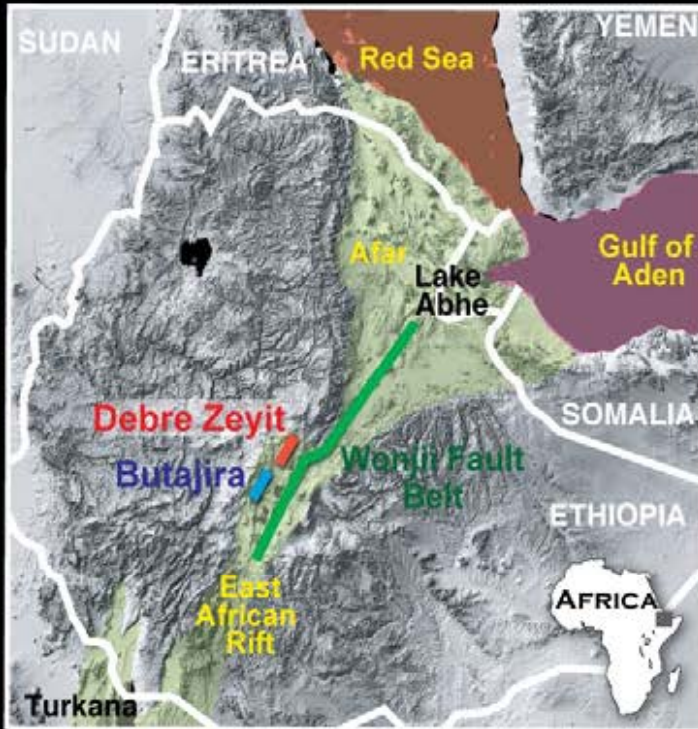


afar plume



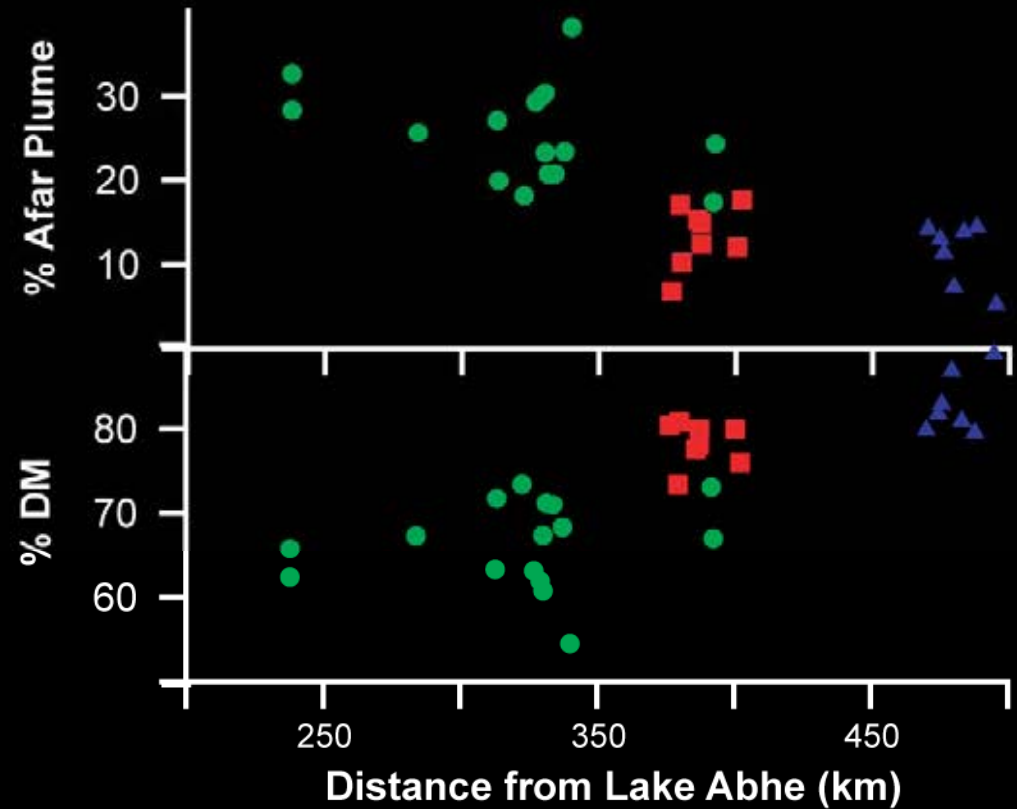
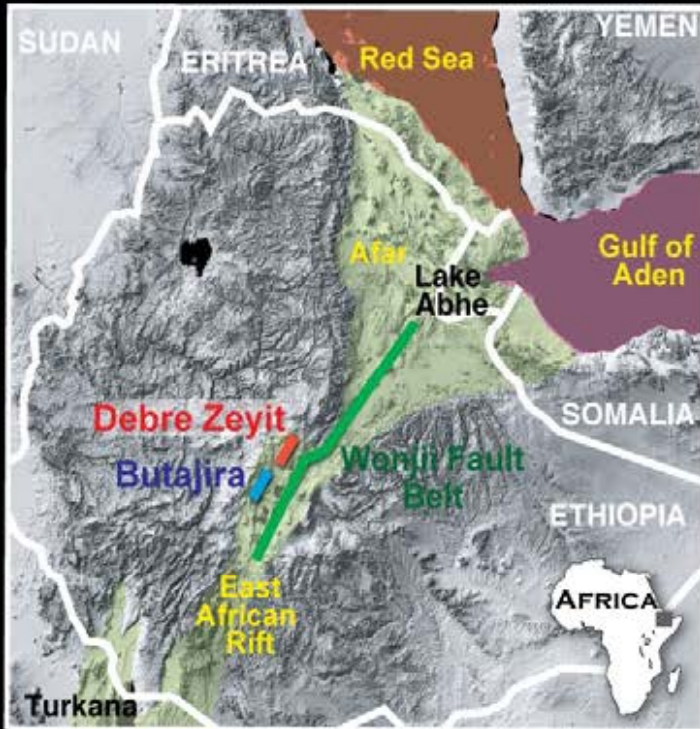
Data from PETDB

plume contribution



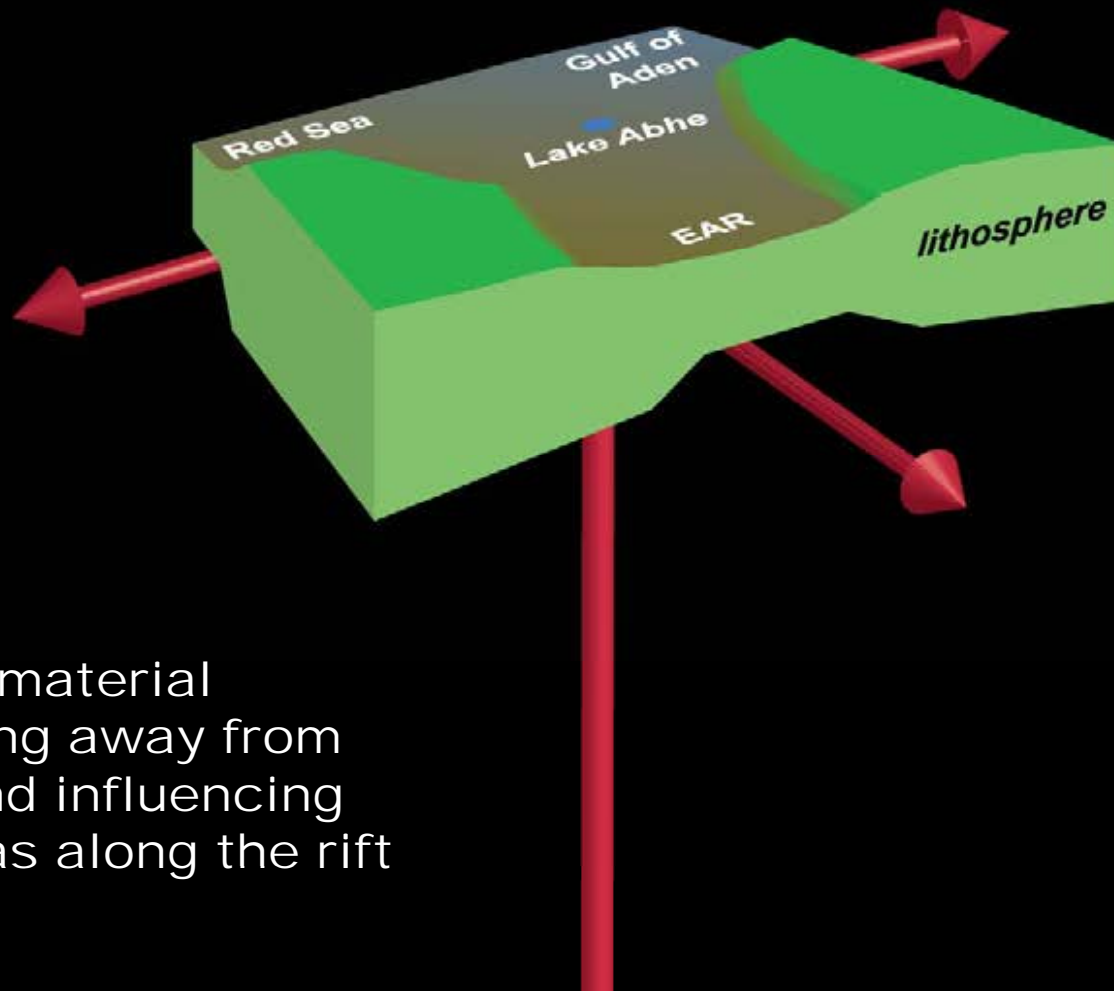
Data from Rooney et al., 2012 *J. PET*

plume contribution



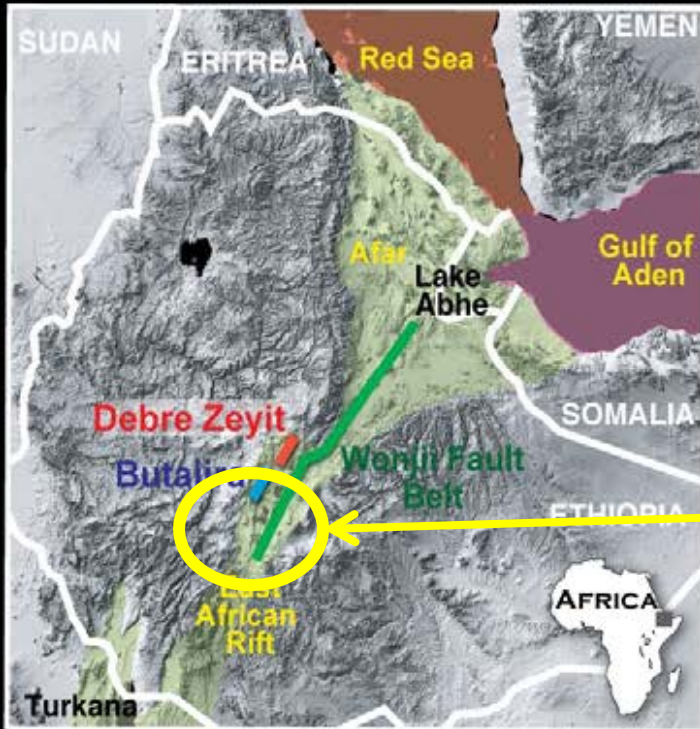
Data from Rooney et al., 2012 - J. PET

radial dispersion

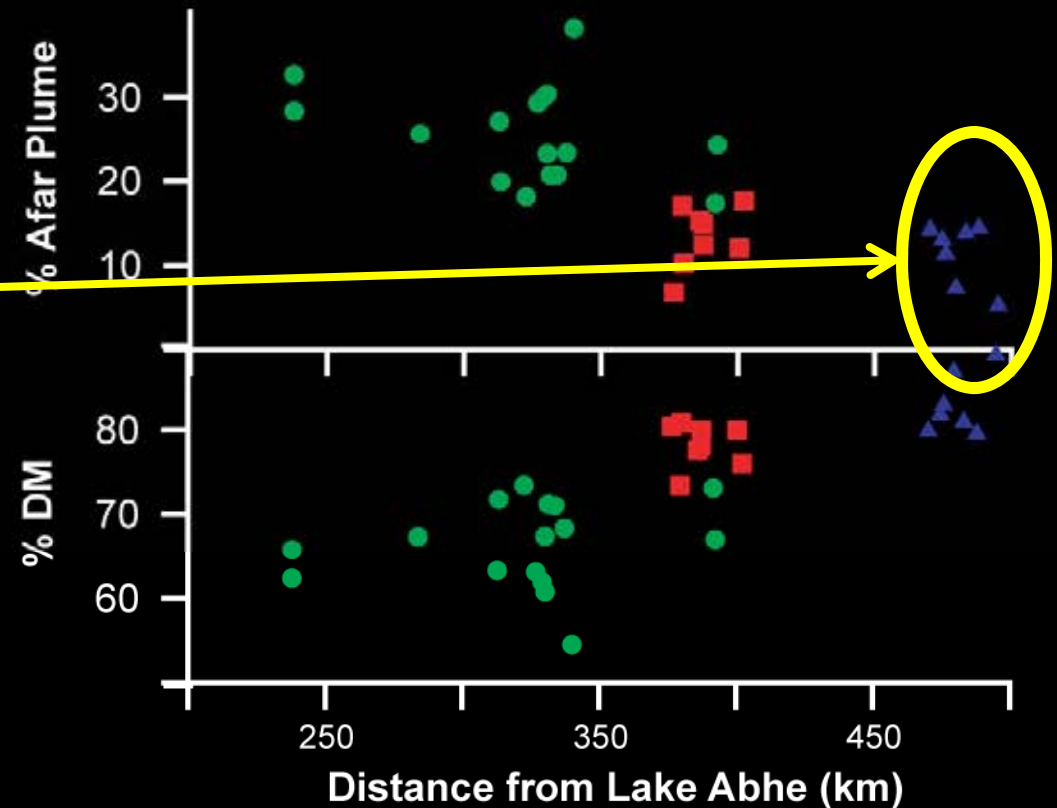


Plume material
radiating away from
Afar and influencing
magmas along the rift

plume contribution



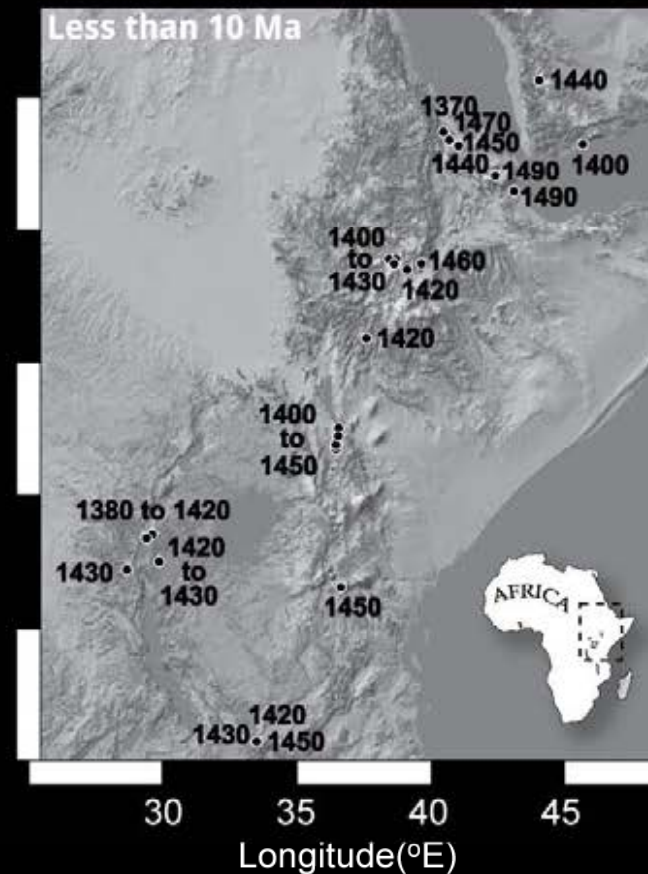
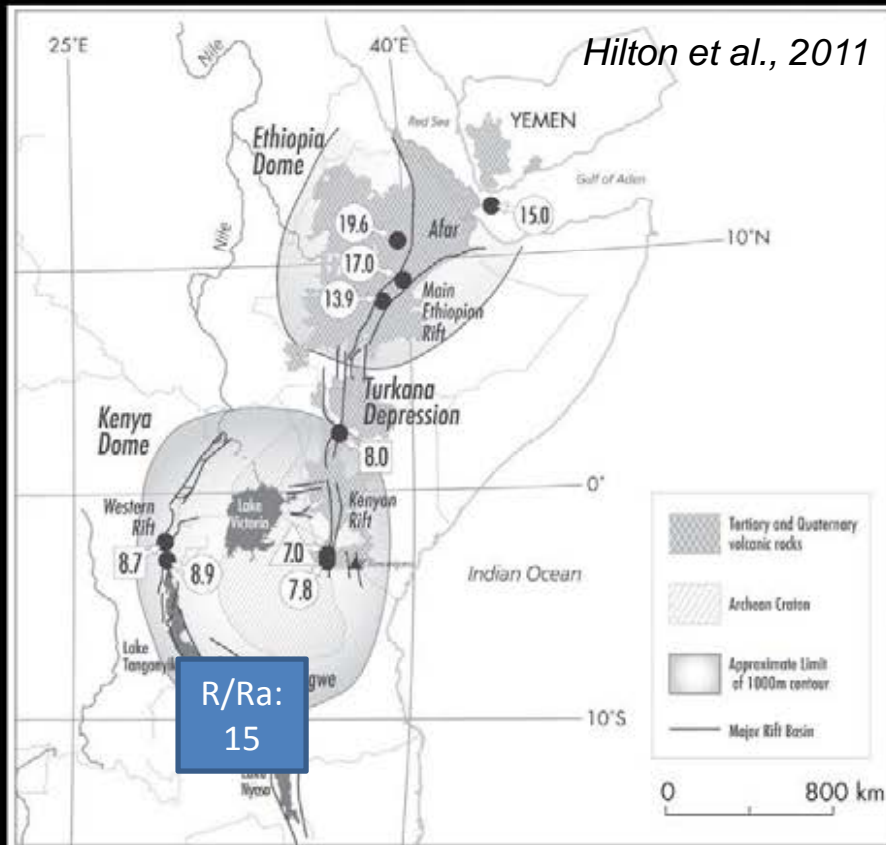
The problem is that we reach zero plume contribution in southern Ethiopia



Data from Rooney et al., 2012 - J. PET

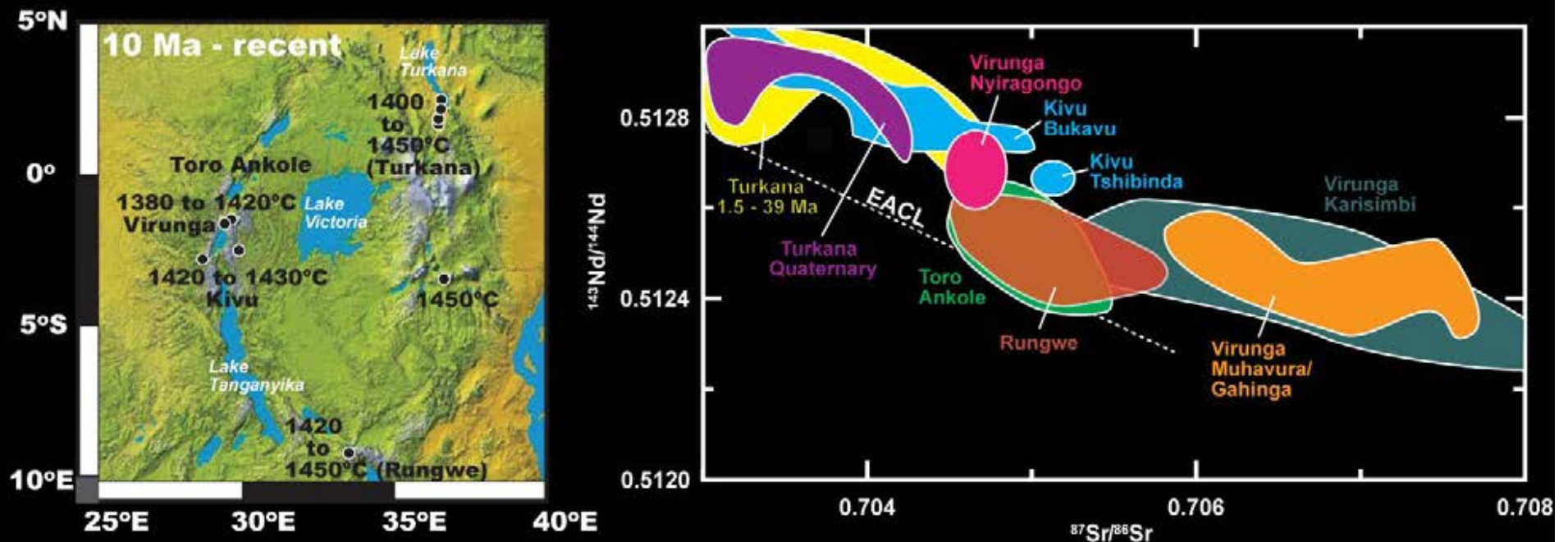
plume influence

There is evidence of a plume contribution in the south too!



southern EAR

Southern portion of the EAR exhibits elevated temperature

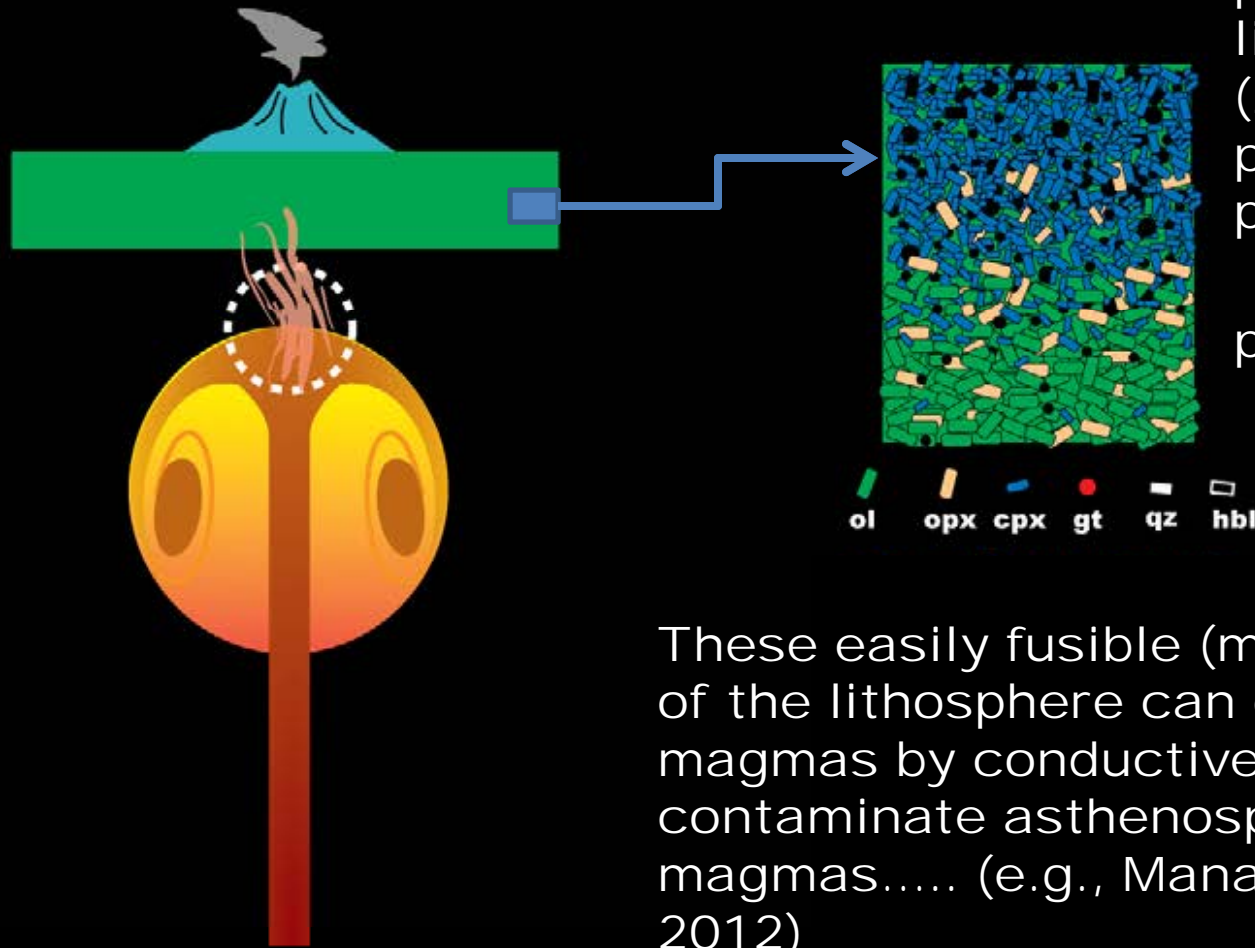


It is difficult to link magmas in the southern EAR with asthenospheric (plume) processes as the veil of the lithosphere obscures critical details.

composition

Hornblende- or
phlogopite-bearing
lithospheric metasome
(200°C lower melting
point than dry
peridotite)

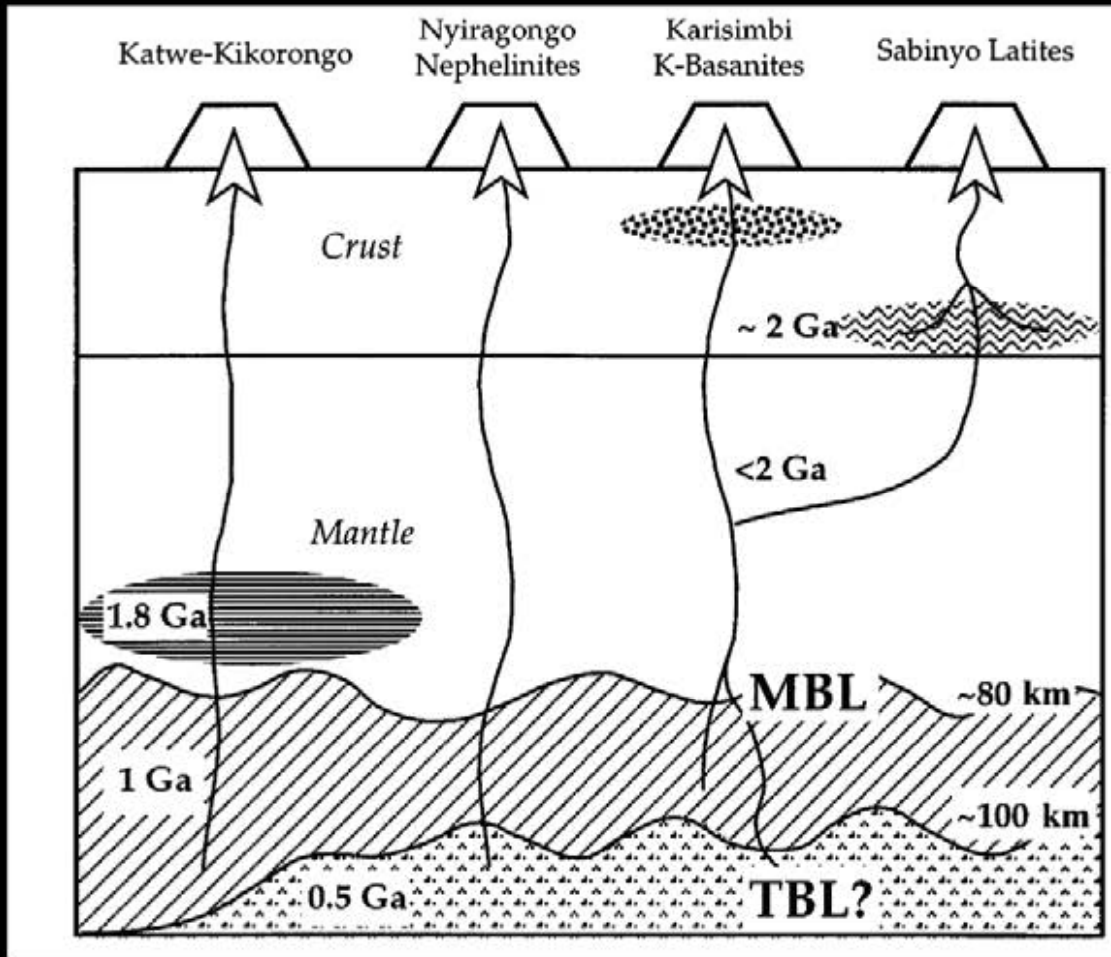
peridotite



These easily fusible (meltable) parts
of the lithosphere can create
magmas by conductive melting or
contaminate asthenospheric
magmas..... (e.g., Mana et al., - Lithos
2012)

lithospheric melt

Rogers et al., 1998

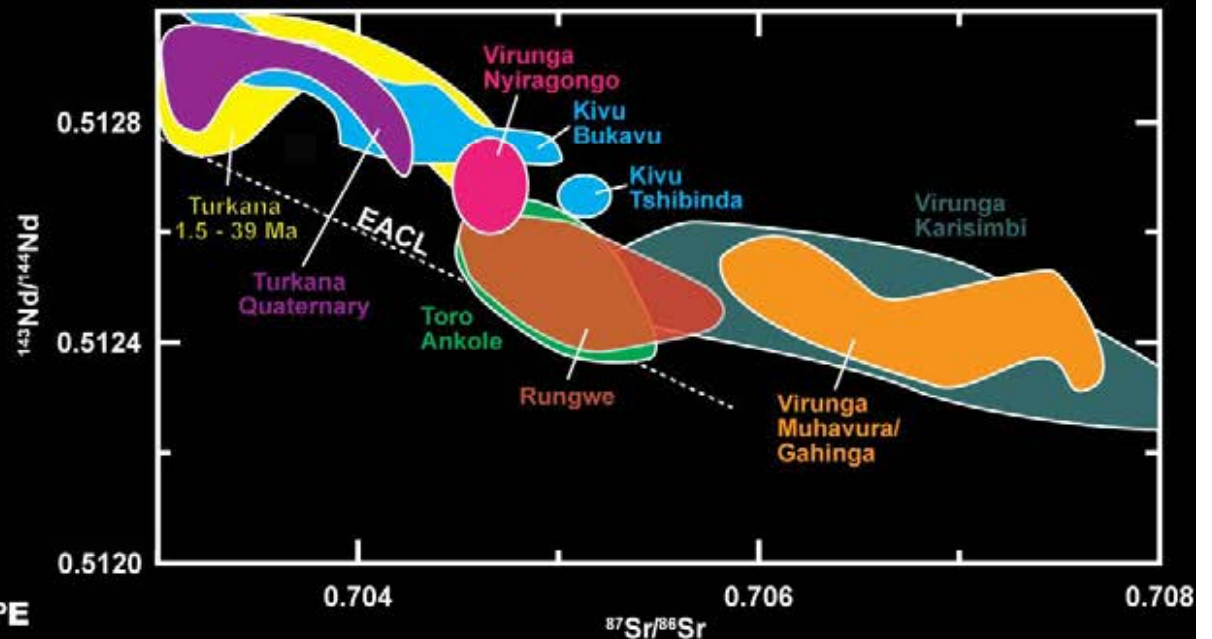
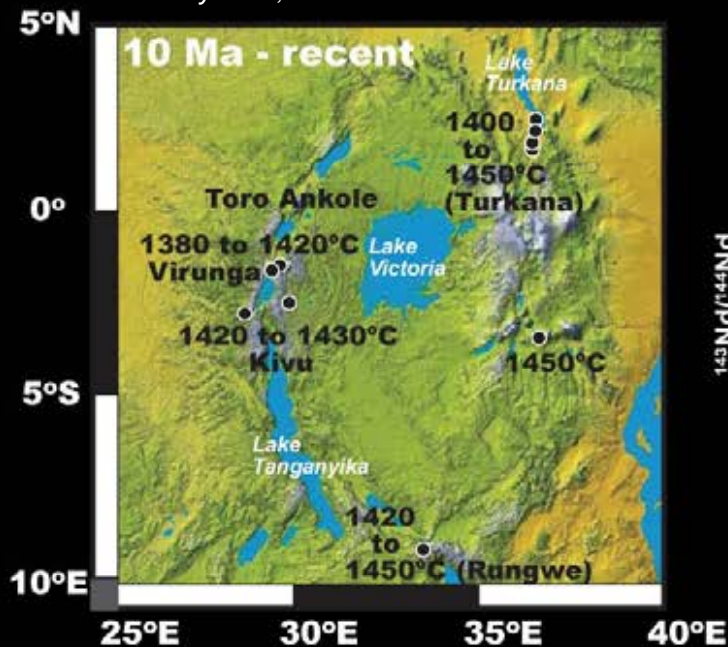


Layered lithosphere can produce a wide array of magmas at the surface

southern EAR

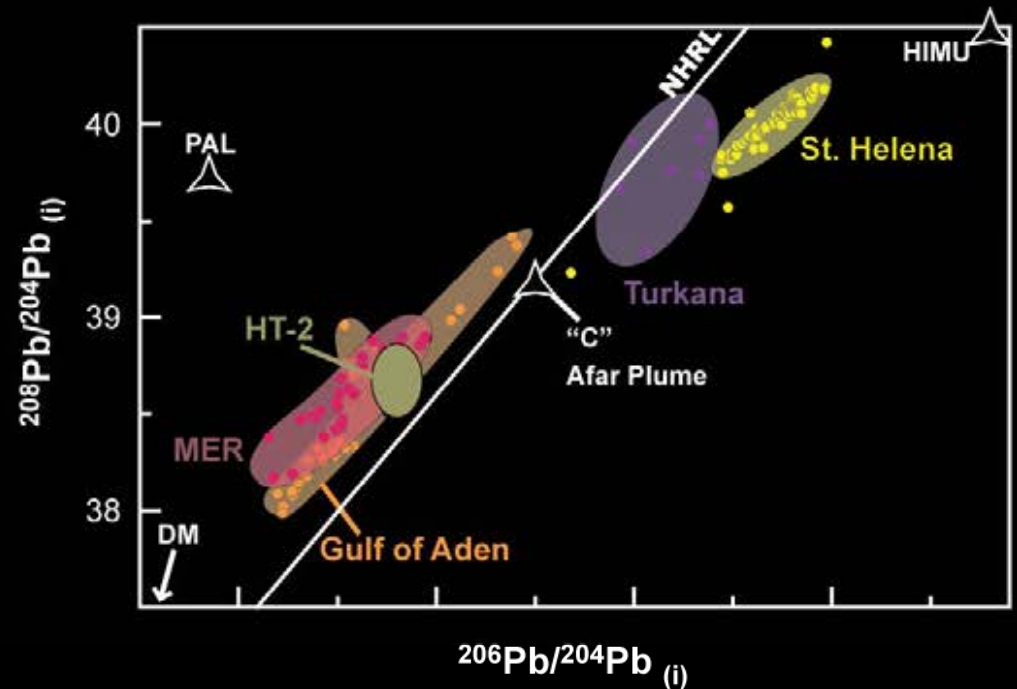
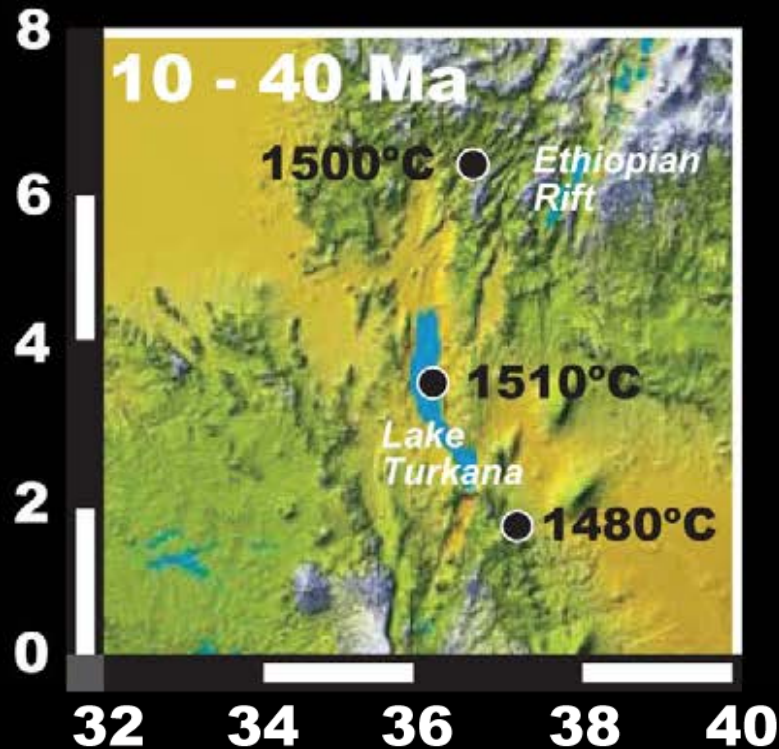
Data: Virunga- Karisimbi, Gahinga, Muhavura (Rogers et al., 1992; 1998); Virunga – Nyiragongo (Chakrabarti et al., 2009); Rungwe (Furman & Graham, 1999); Toro Ankole (Rosenthal et al., 2009); Kivu – Tshibinda, Bukavu (Furman & Graham 1999); Turkana – Quaternary (Furman et al., 2004); Turkana 1.5-39 Ma (Furman et al., 2006). The EAACL is the East Africa Carbonatite line (Bell & Blenkinsop 1987).

Rooney et al., 2012



While much of the southern EAR has a lithospheric overprint – The Turkana region may offer clues as to the asthenospheric processes active here

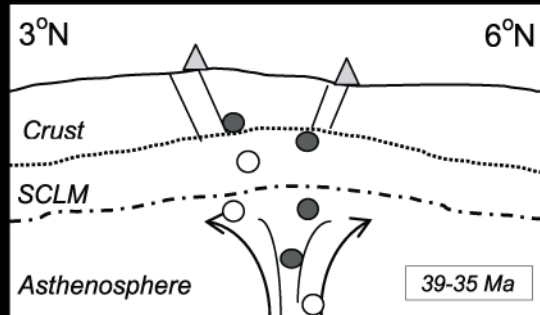
turkana



Gulf of Aden (Schilling et al., 1992; Rooney et al., 2012); St. Helena (GEOROC); Turkana (Furman et al., 2004;2006); 30 Ma HT-2 (Pik et al., 1999). MER – Main Ethiopian Rift (Furman et al., 2006a; Rooney et al., 2012). Endmembers PAL (Pan African Lithosphere); DM (Depleted mantle); "C" – Afar Plume are from Rooney et al., 2012.

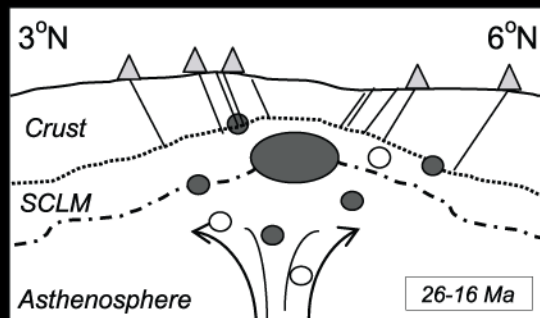
turkana model

Furman et al., 2006

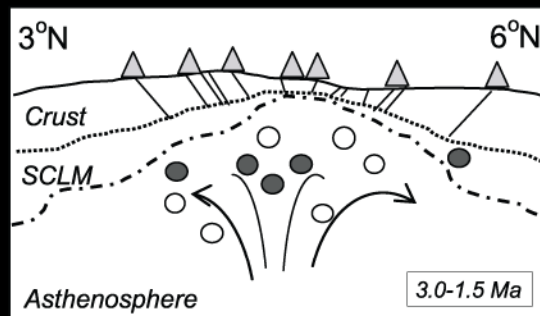


Model for Turkana

-Oligocene initial mixed plume component



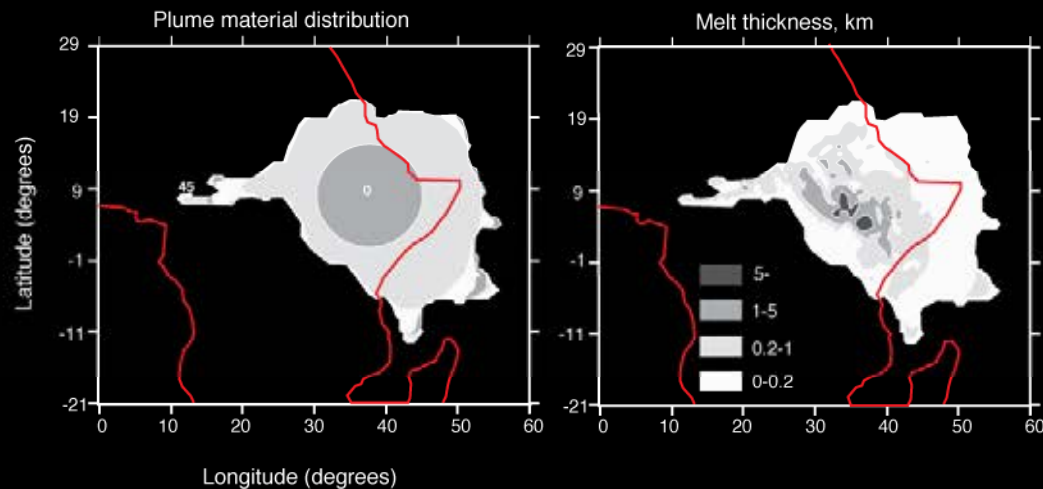
-Early Miocene, the HIMU flavor dominates



-Pliocene – Quaternary, the “C” signature dominates

single plume

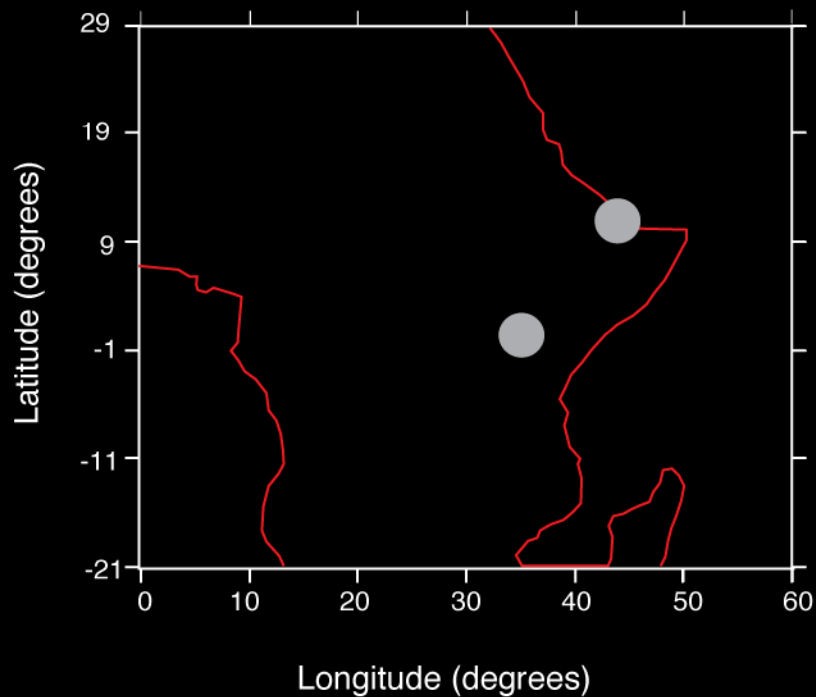
Ebinger & Sleep, 1998



A single plume impacted the lithosphere and radiated outwards, channeled along the base of the lithosphere.

two plume

George et al., 1998; George & Rogers, (2002)



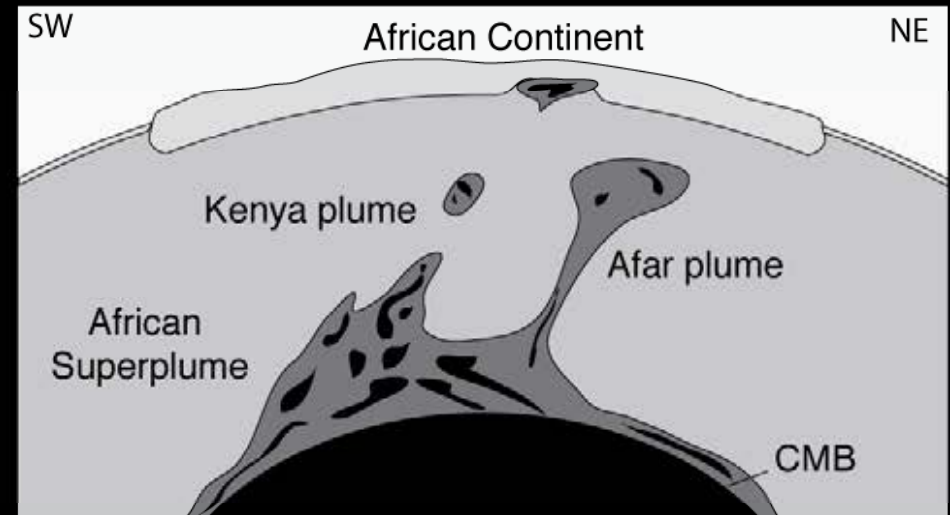
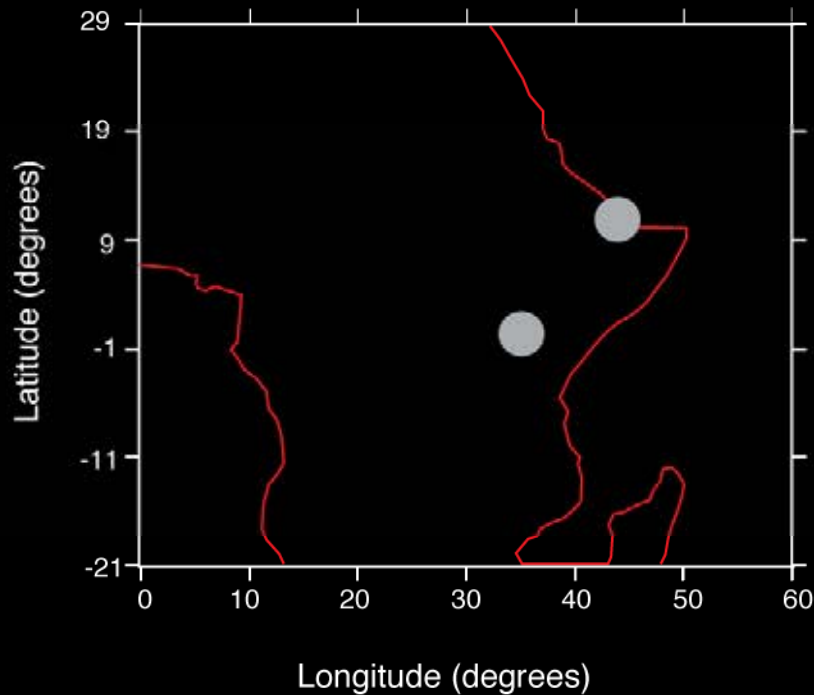
Geochemistry requires
two sources

-Afar plume

-Another plume
beneath Tanzania
Craton?

modified single plume

Furman et al., 2006; Nelson et al., 2012



challenges

What is the extent of the temporal and spatial contribution of plume materials to rift magmas?

Are the plume(s) spatially or temporally heterogeneous?

How does plume-lithosphere interaction control rift development?

How do we distinguish lithospheric enriched zones from plume signatures?

How deep could melt be generated and how can the impact the seismic images of the region?

Can we communicate these challenges and work with the other disciplines to come to some conclusions?