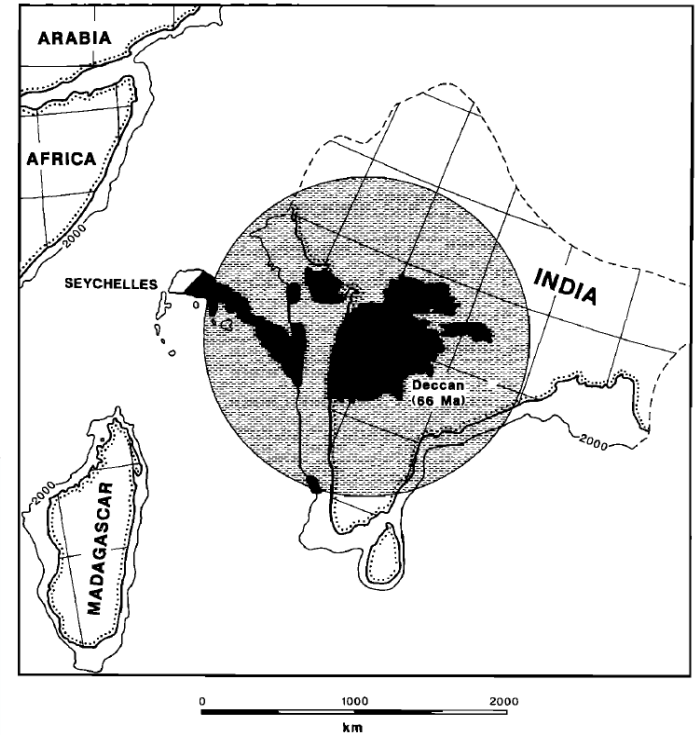
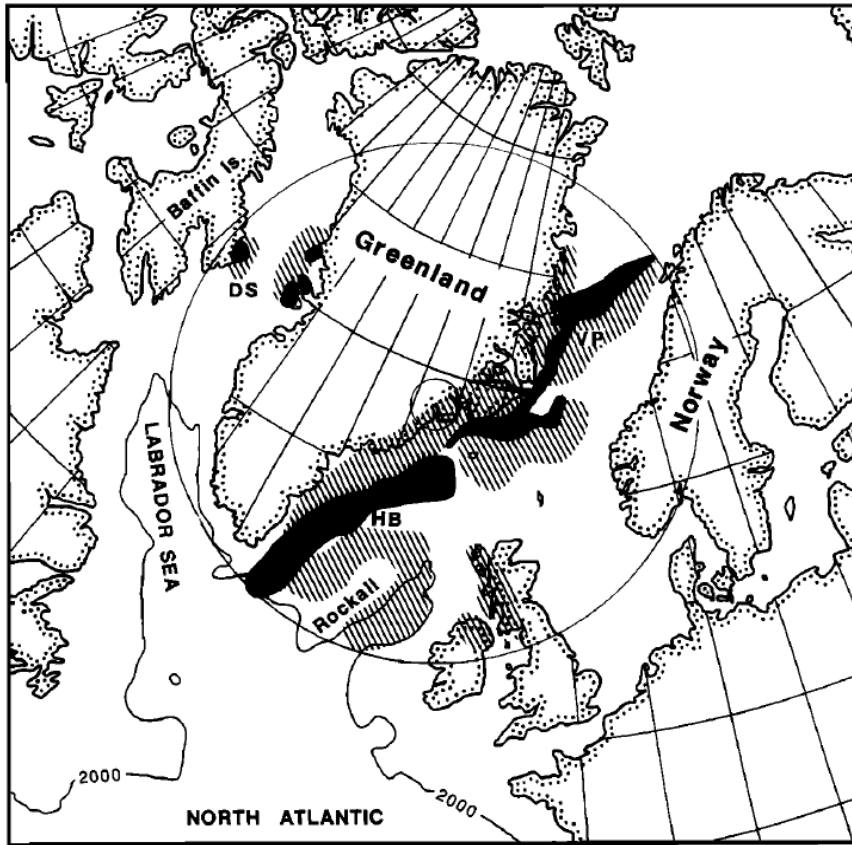


Anologue and Numerical models that inform the rifting process

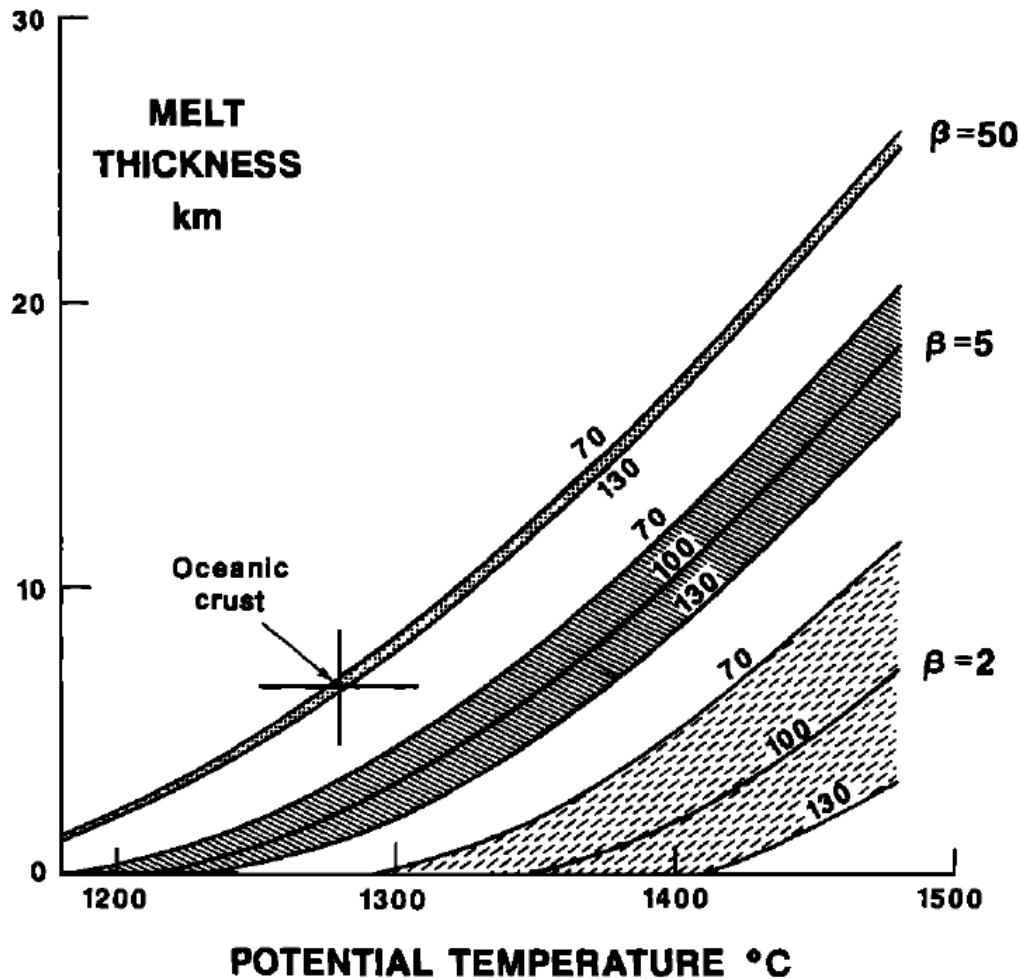
John Armitage



On shore flood basalts and volcanic margins

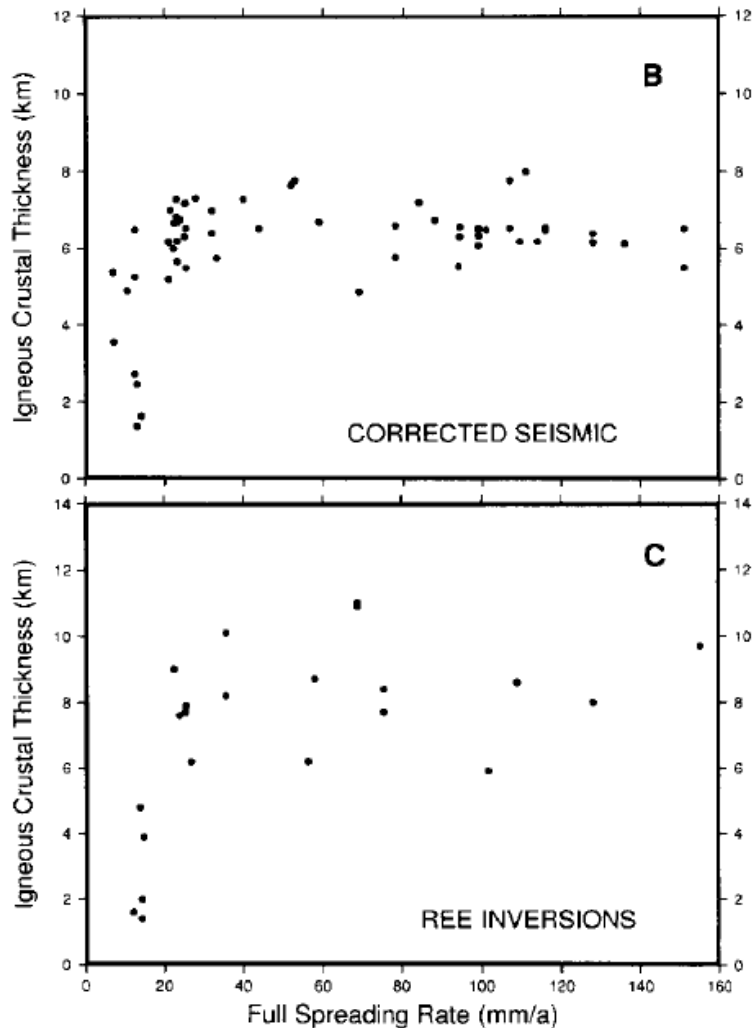


Mantle temperature, lithosphere and melt thickness



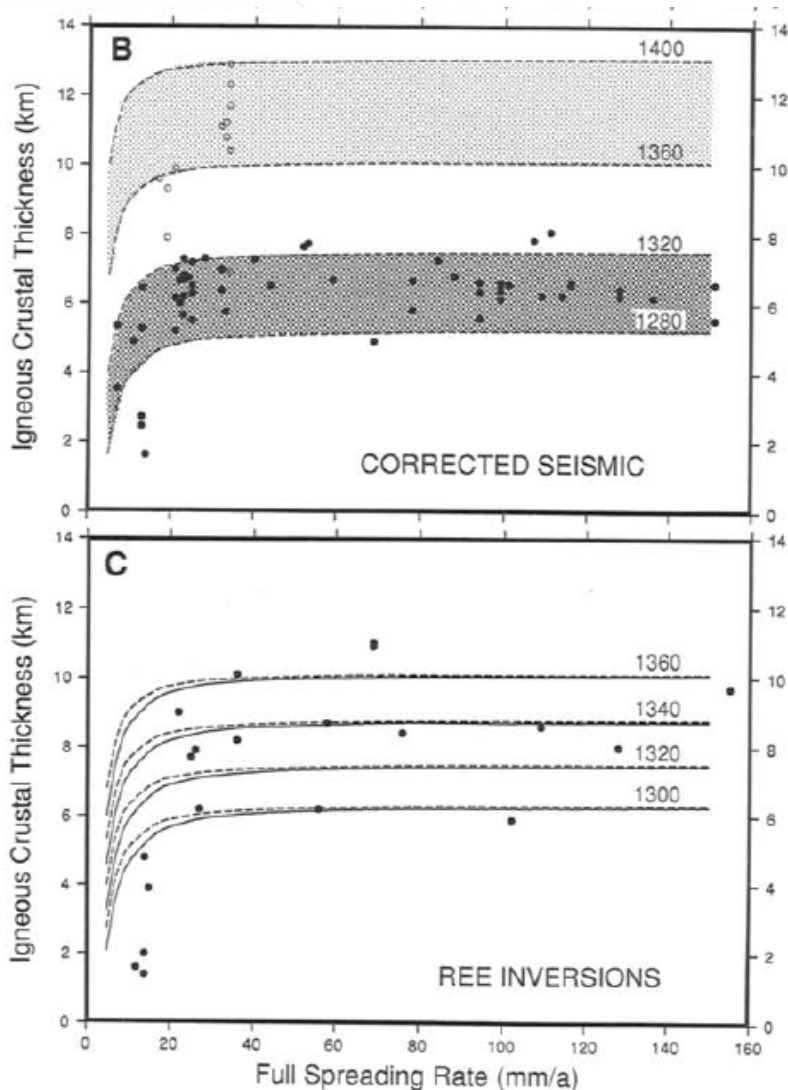
White & McKenzie, JGR, 1989

How do you make less melt?



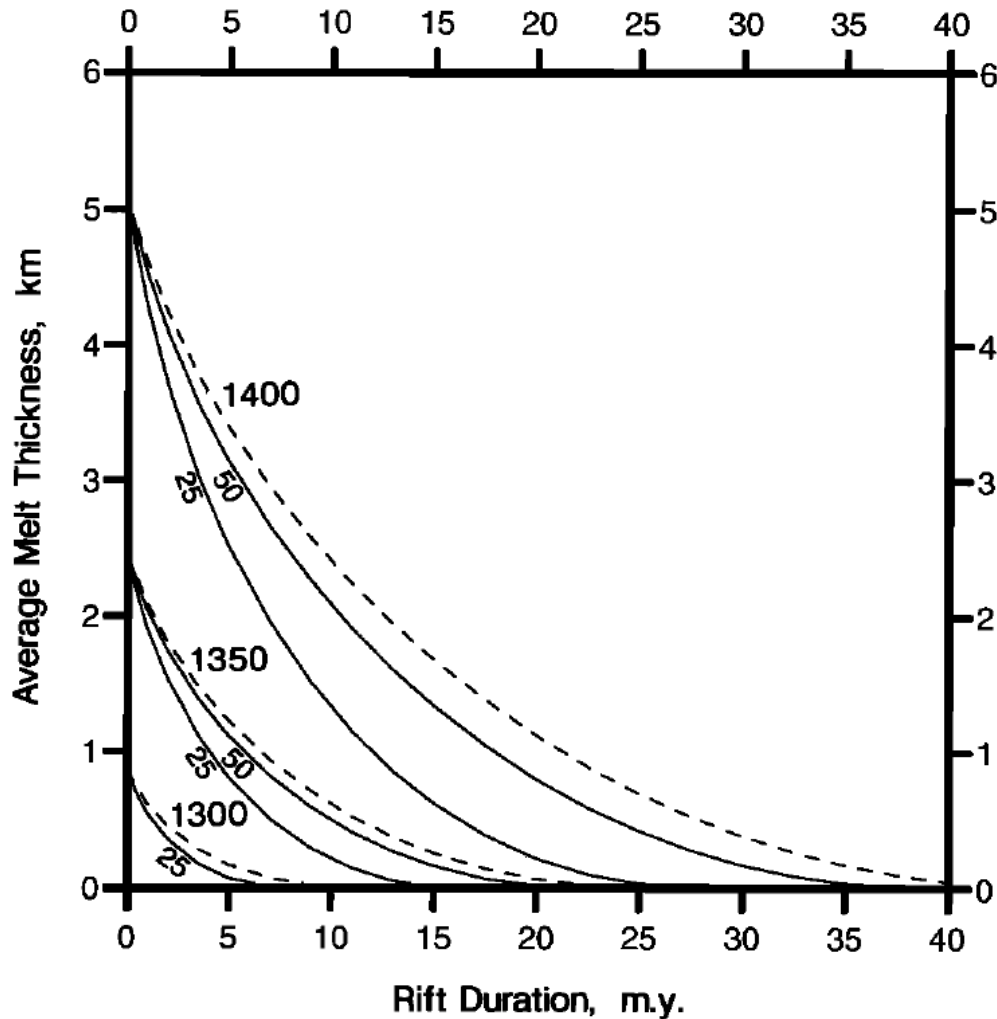
- At low rates of extension (< 20 mm/yr full rate) the crustal thickness can be low (< 5 km).
- Crustal thicknesses at mid-ocean ridges are typically between 5 and 8 km.
- Bown & White, EPSL, 1994.

How do you make less melt?



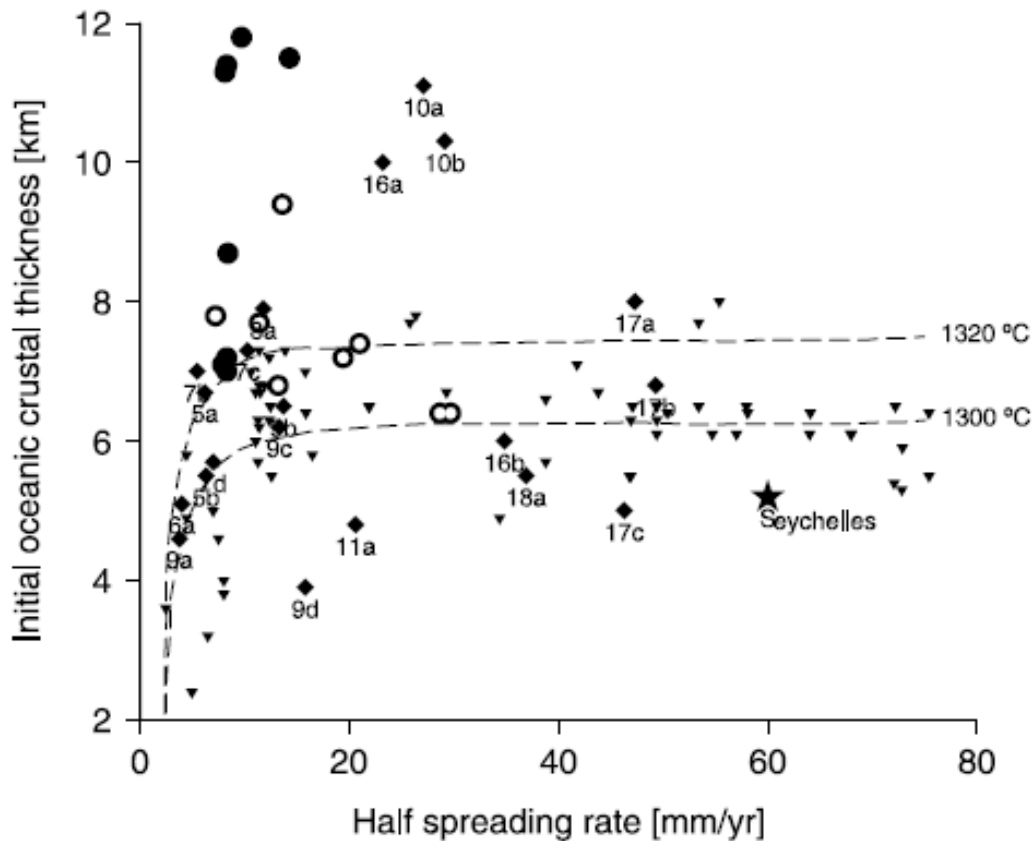
- For mantle of a temperature of 1280 to 1320 °C a corner-flow model can match thickness.
- The North Atlantic margins fit a 100 °C hotter mantle.

How do you make less melt?



- Extend slowly, forming a wide rift with low melt volumes
- Average melt thickness is the volume of melt generated divided by the width.
- Bown & White, JGR, 1995

Update the plot – add a few margins...



- Not all slow forming margins lack significant melt.
- Not all thin margins are slow...

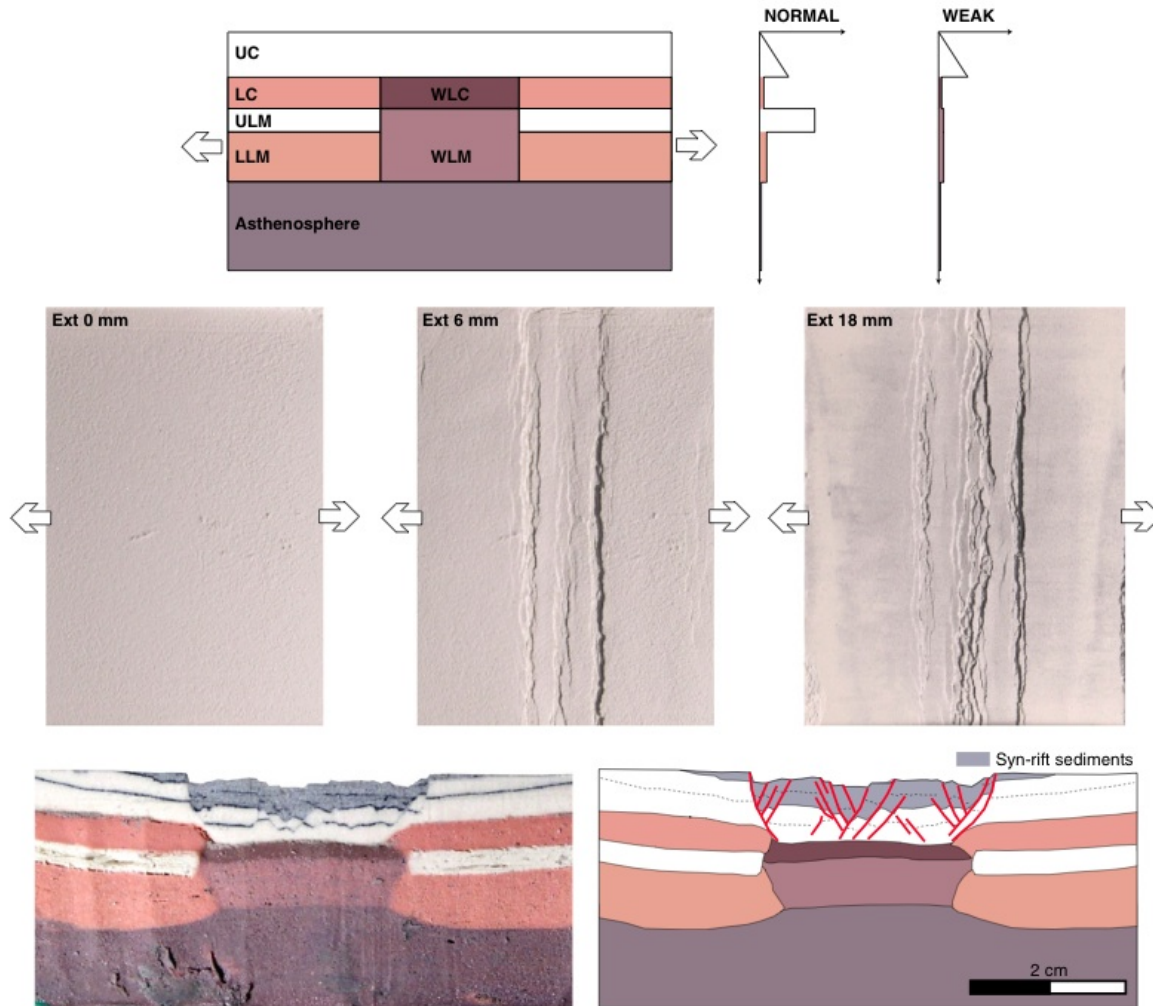
Furthermore, is break-up not by definition volcanic?

- Sills observed in the Newfoundland margin. Evidence for post-rift magmatism (Peron-Pinvidic et al., JGI, 2010).
- The magnetic J-anomaly is possibly associated with magmatic intrusions (Bronner et al., N. Geo., 2011).
- What comes first, melting or exhumation and serpentinitisation?

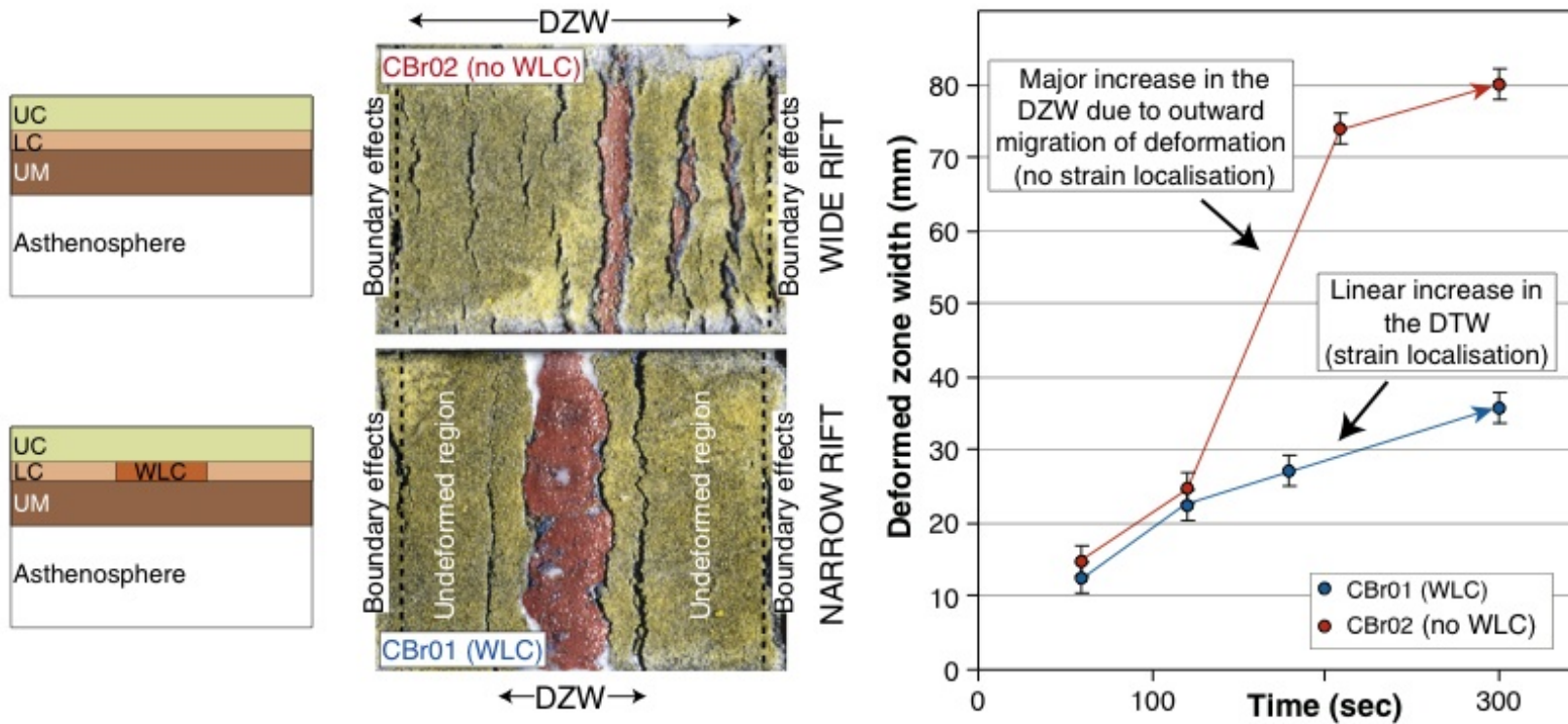
How do you thin lithosphere without melting (too much)?

- In the right circumstance:
 - The lack or not of pre-existing weakness.
- With the right sort of lithosphere:
 - Serpentinisation of upper mantle.
 - Strain rate softening.
- *Or, is the mantle that upwells a spent force – depleted.*

A pre-existing weak zone?



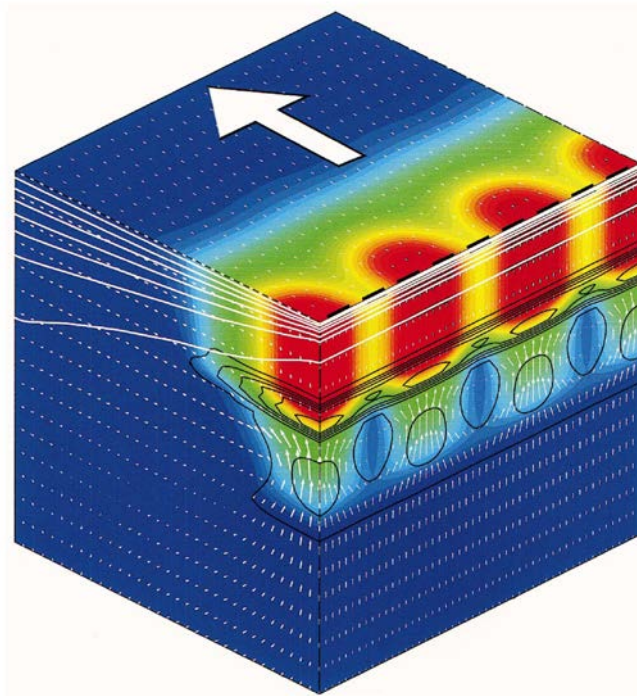
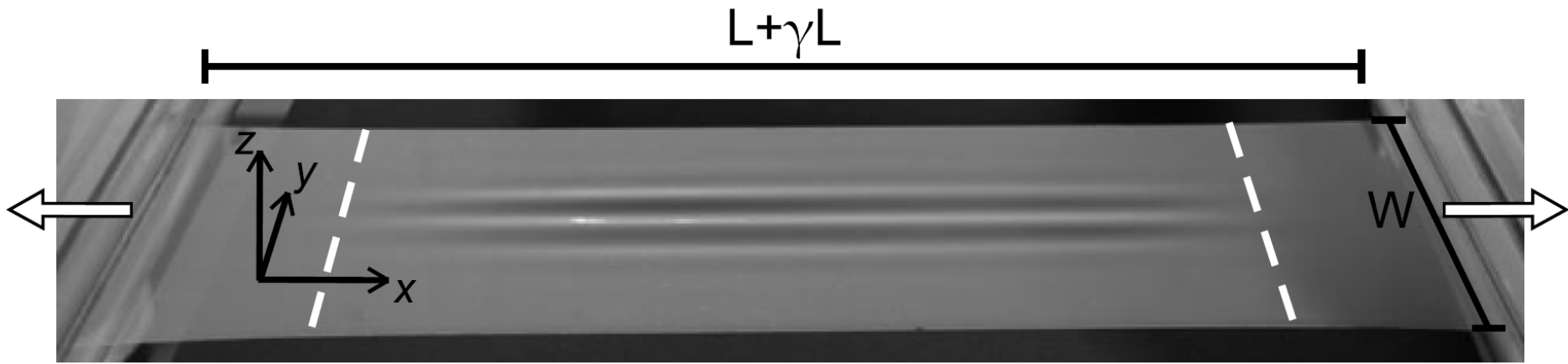
Or no pre-existing weakness



DZW – deformed zone width

Corti et al., GRL, 2003

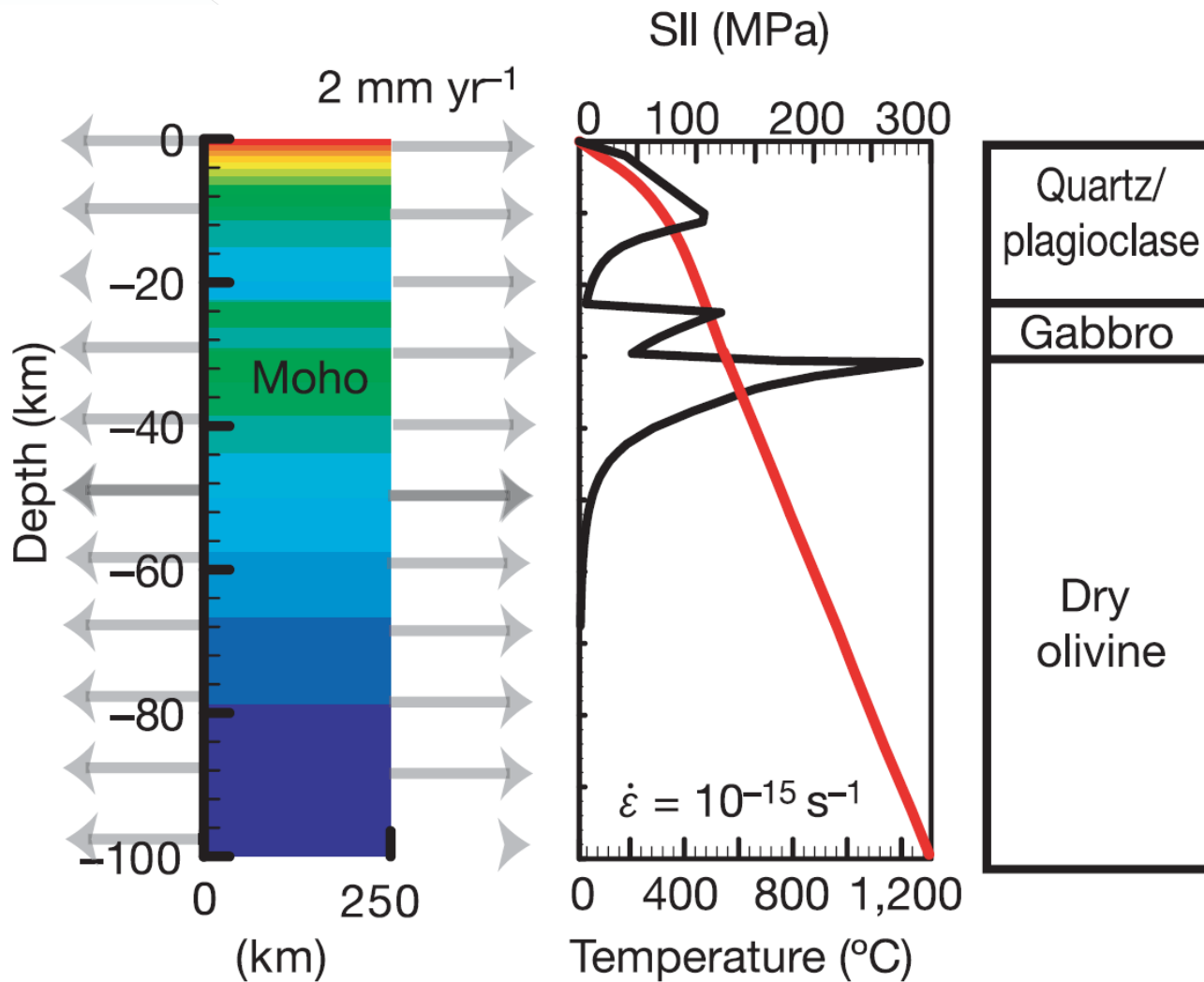
Wrinkles due to stretching



Levy & Jaupart, JGI, 2011

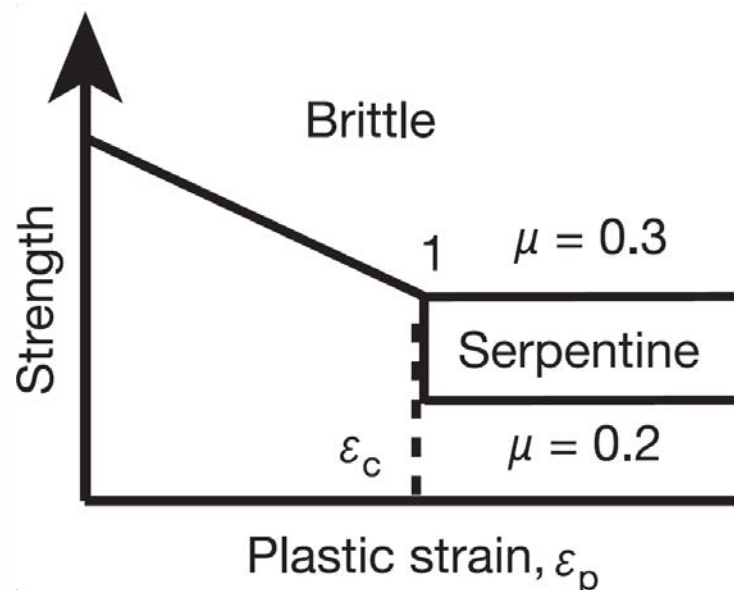
Choblet & Parmentier, EPSL, 2001

Weak lower crust

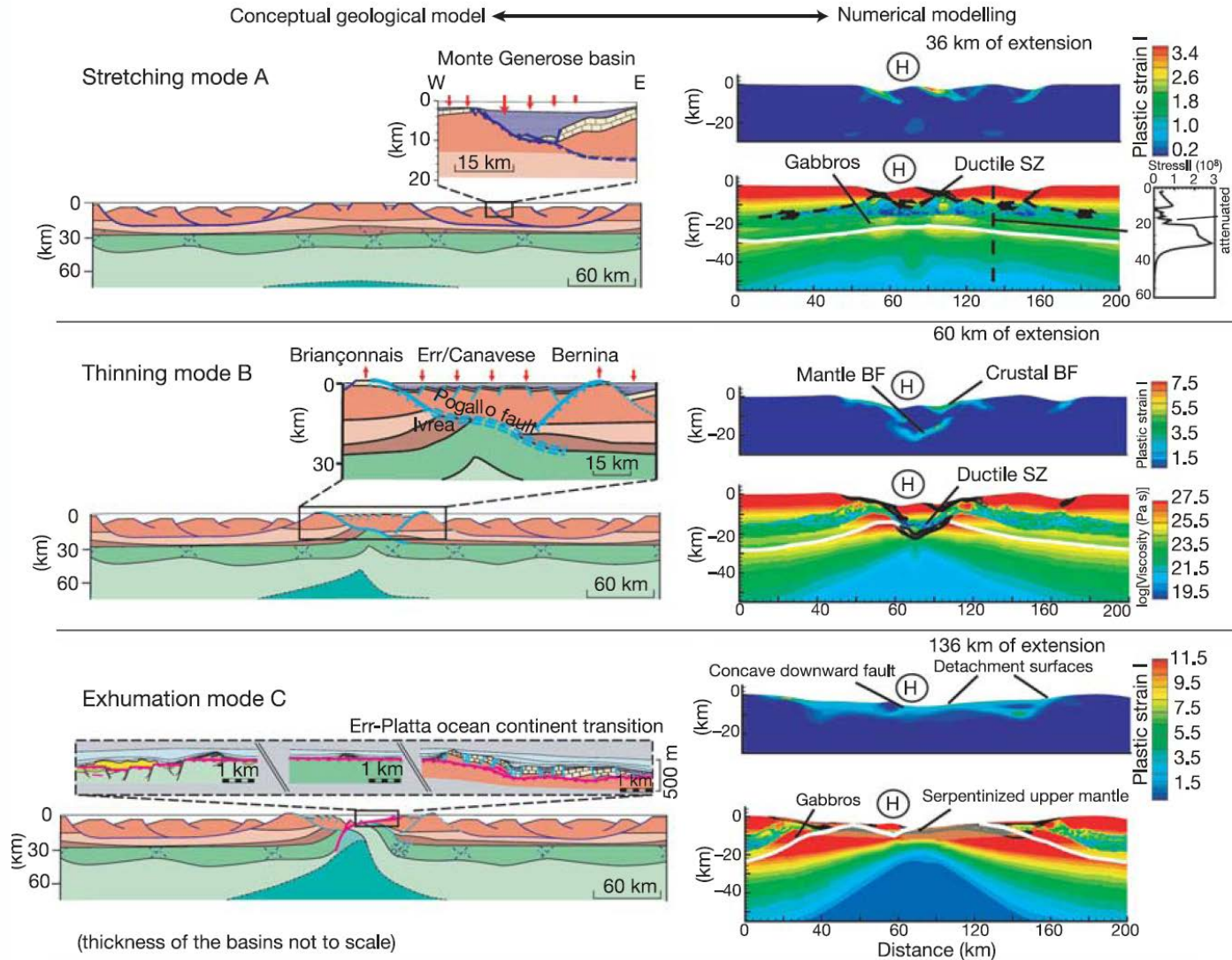


Serpentinite

- Weakens the upper mantle as it is formed as it lowers the friction coefficient within the Mohr Coulomb yield criterion.



Stretch, thin and exhume

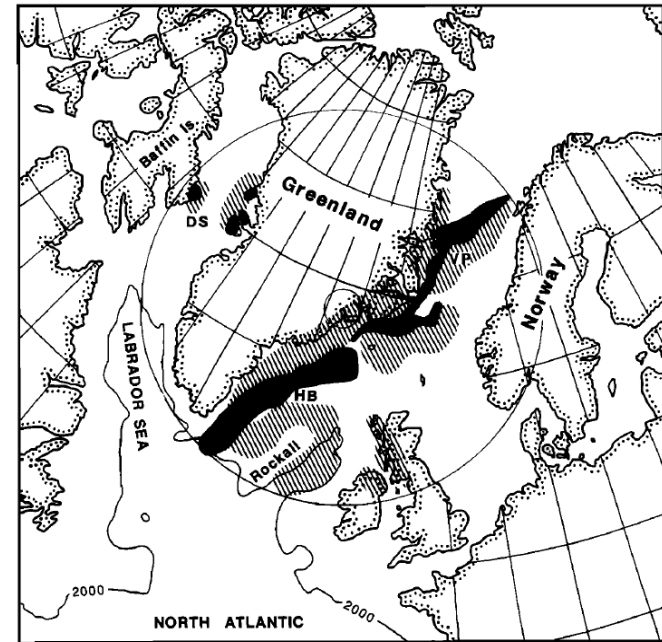


So what is it to be?

- ◉ Weak layers to facilitate thinning/shearing of the continental lithosphere.
- ◉ Pre-existing weakness in the crust and lithosphere, rifts form in old suture zones, don't they?
- ◉ Or is melt key?

After all, not all rifts succeed...

Paleocene - Early Eocene

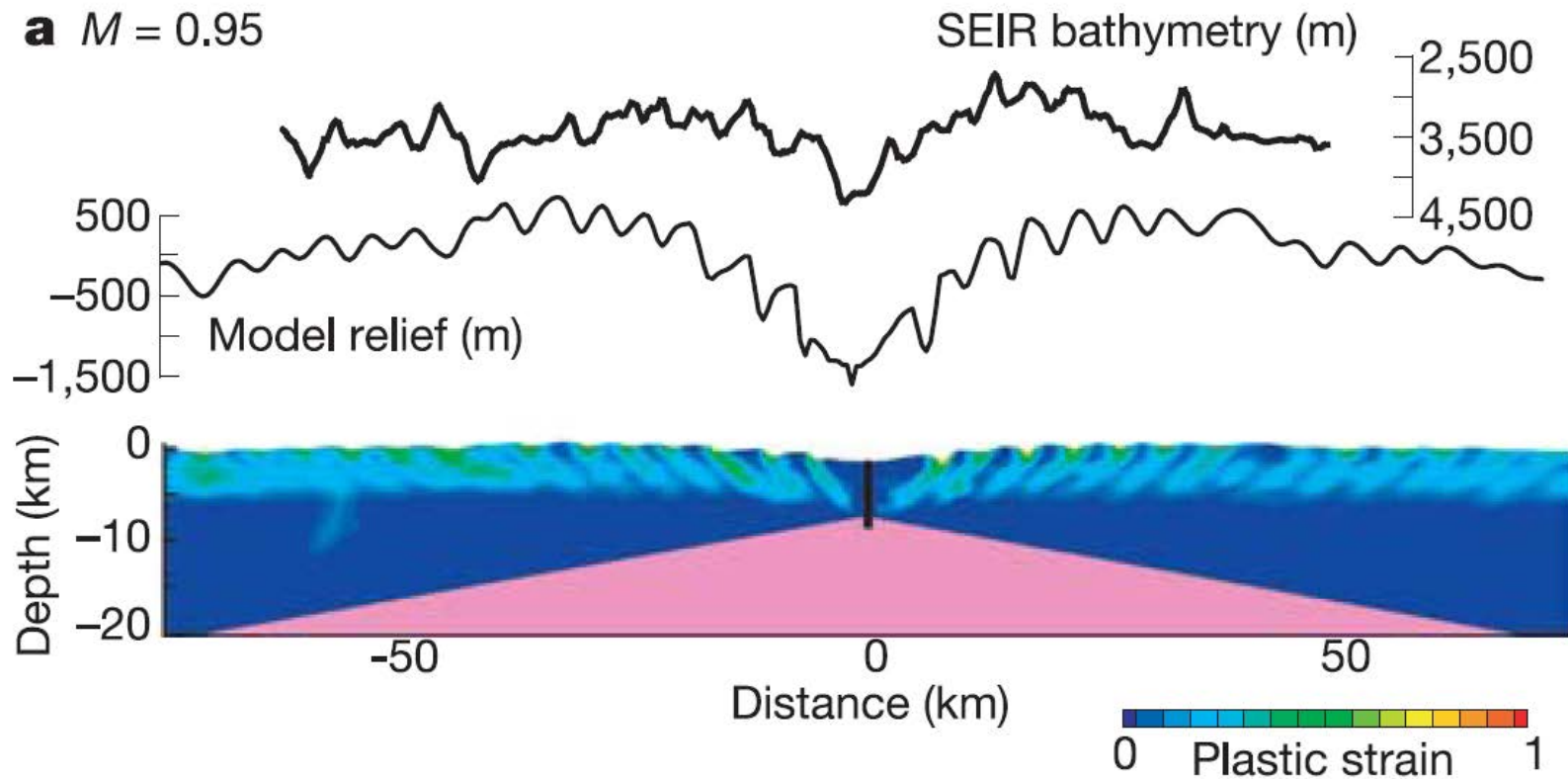


0 500 1000
km

Armitage *et al.*, *G-cubed* (2009)

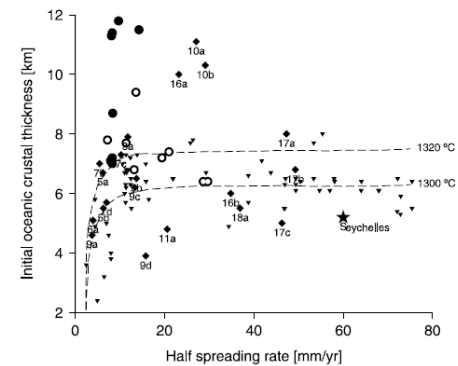
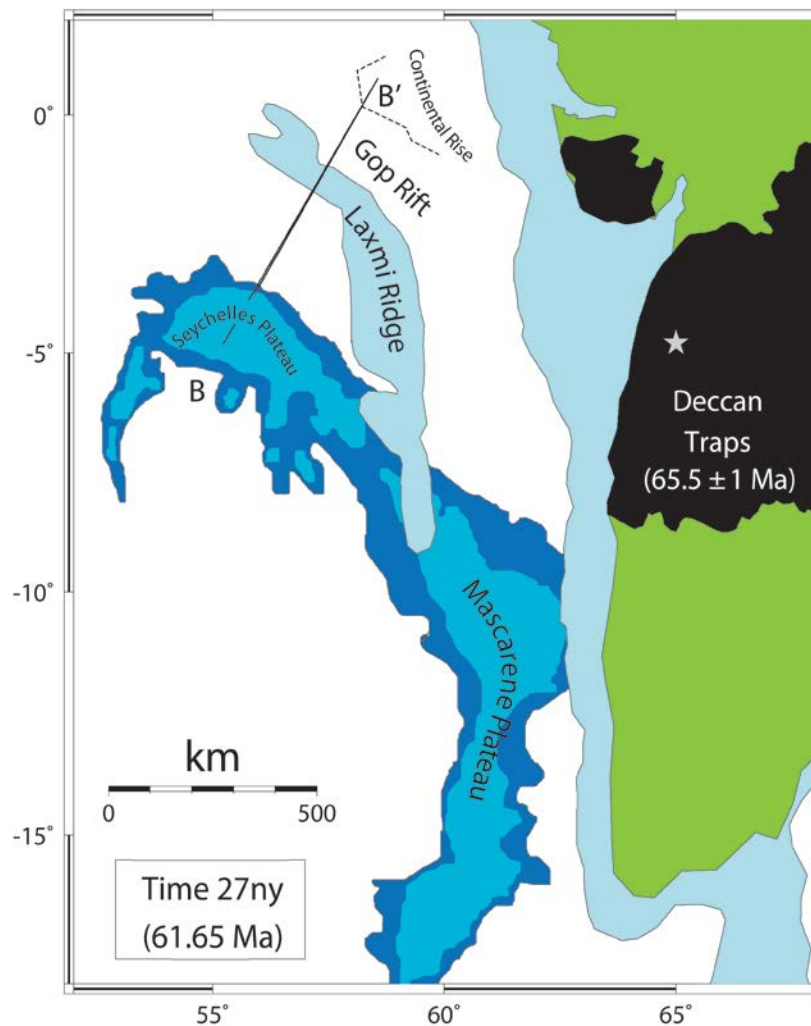
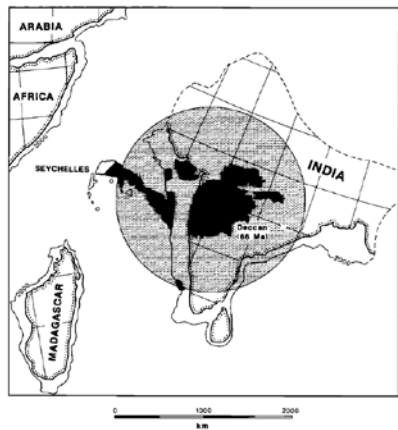
White & McKenzie, *JGR* (1989)

and melt will thermally weaken...



- At mid-ocean ridges, slow extension and melt intrusion match bathymetry.
- Buck et al., Nature, 2007

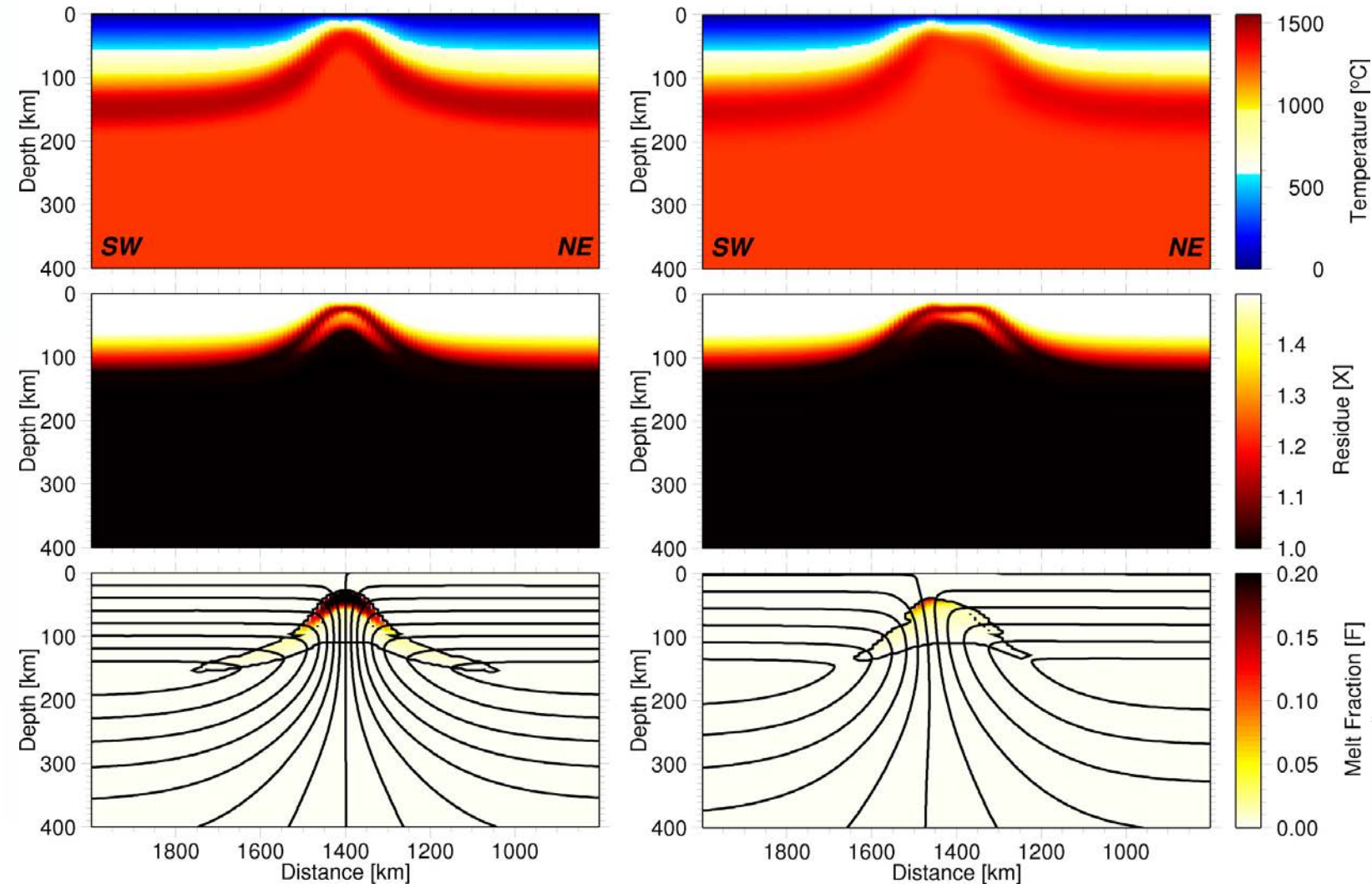
Could the rift history control magmatism?



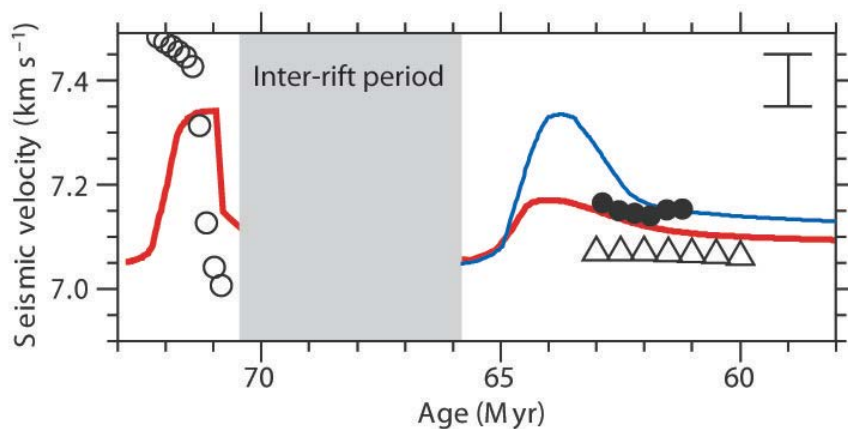
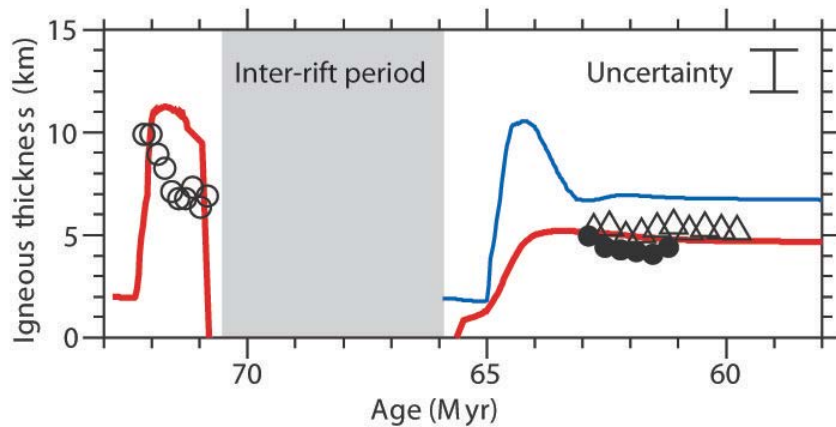
Rift history effects melt generation

(a) Lithosphere structure at 71 Ma

(b) Lithosphere structure at 65 Ma



Such that break-up goes with a bang or a whimper?



- Prior extension depletes the asthenosphere such that once break-up is achieved the mantle is depleted.
- The Seychelles margin is thin despite an association with flood basalts.
- Armitage et al., Nature, 2010.

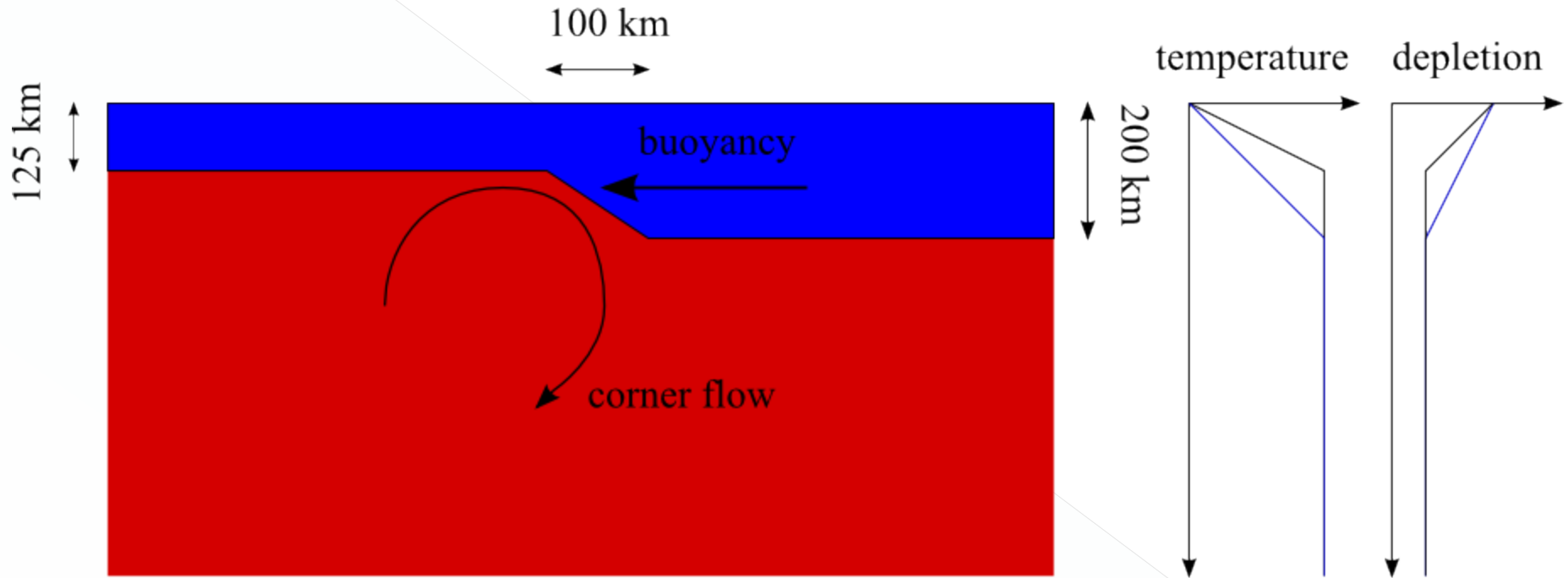
1. How do you localise extension – break-up?

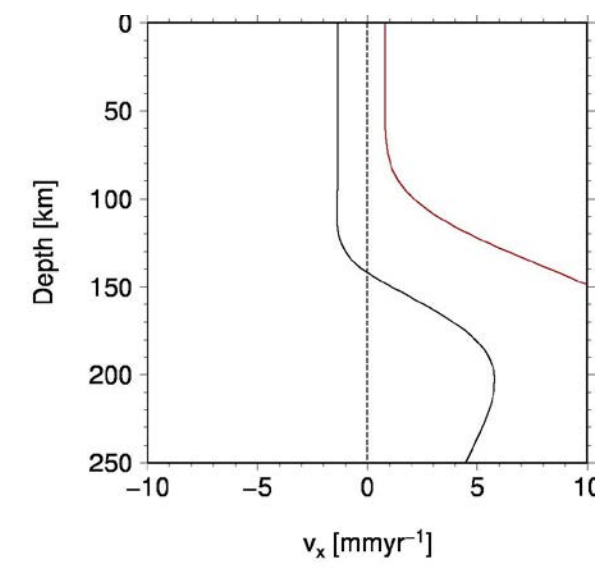
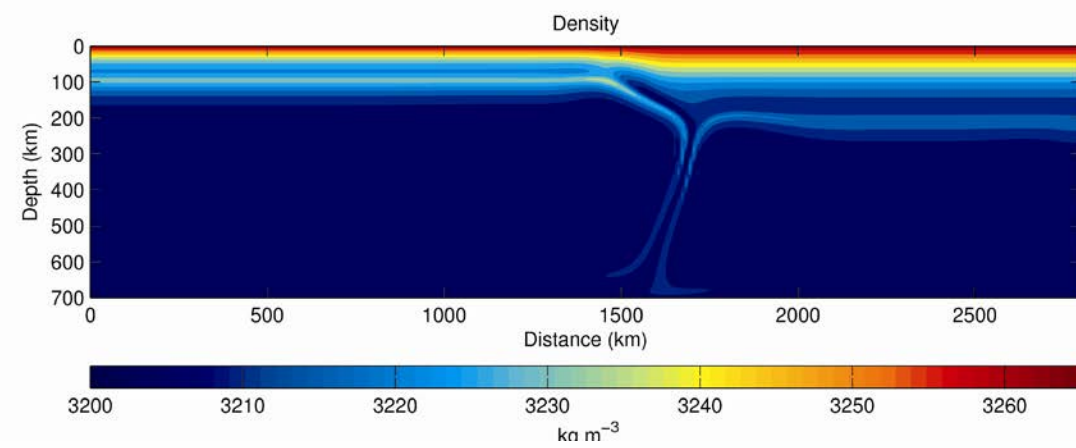
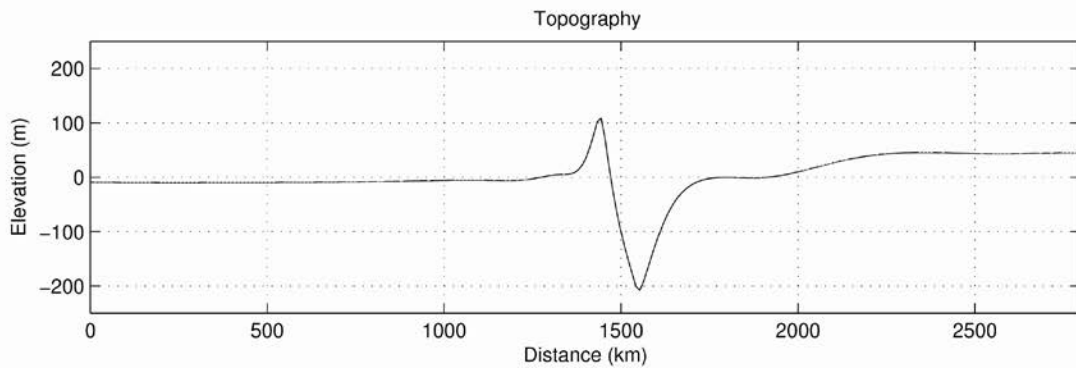
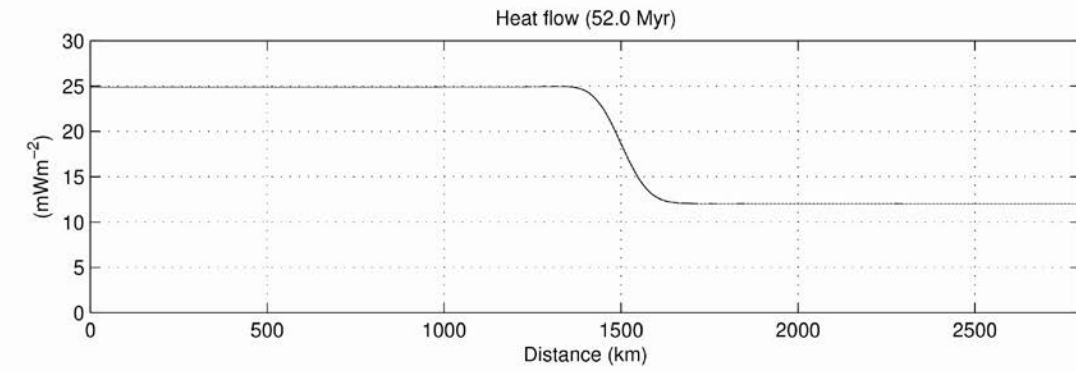
- a) Pre-existing local weakness in the upper lithosphere.
- b) Serpentinisation and a weak lower crust.
- c) Melt intrusion and so at least some melt generation.

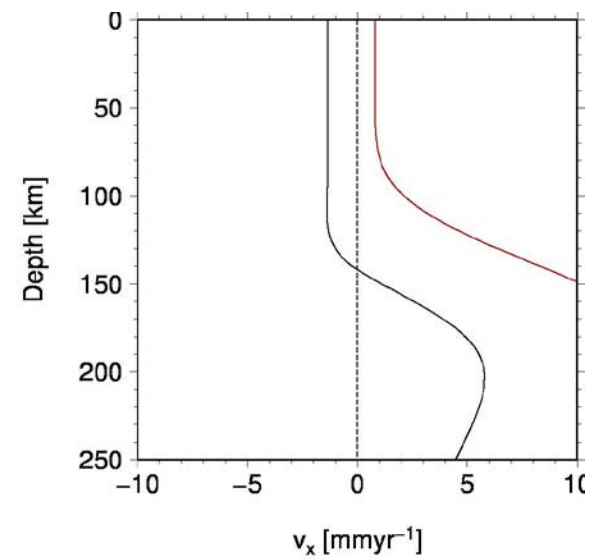
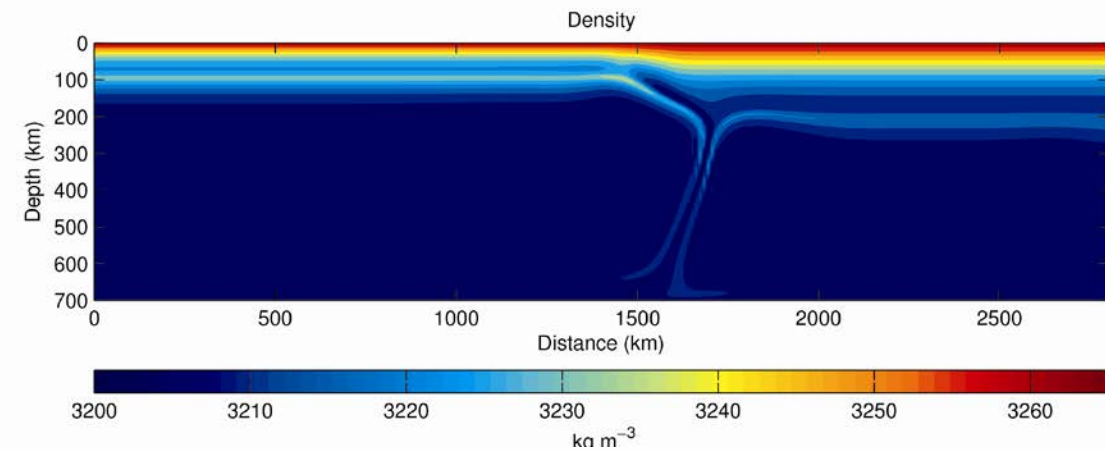
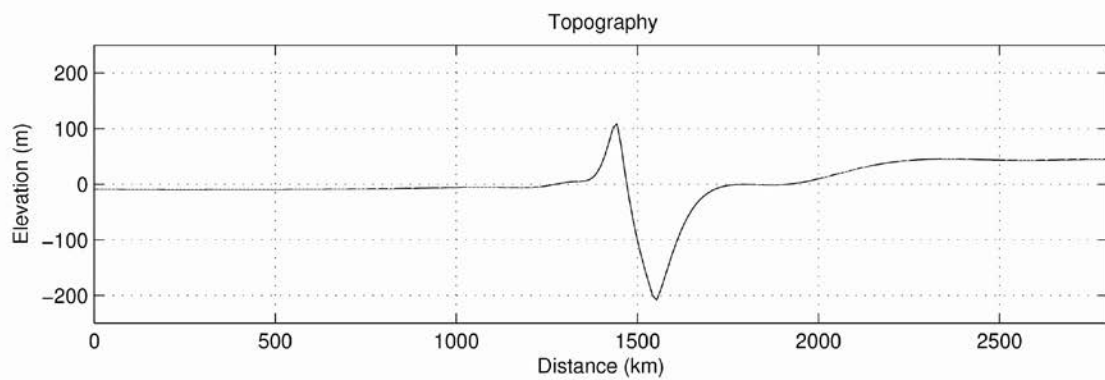
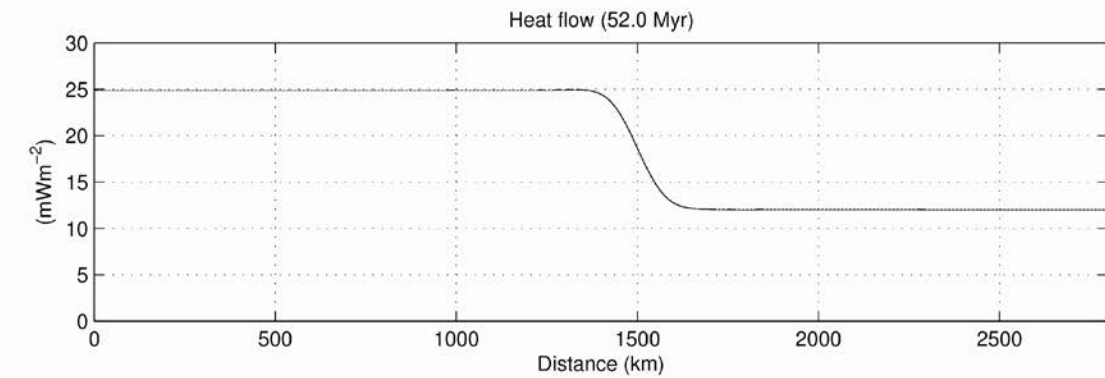
2. What controls melt generation?

- a) Temperature. But a-magmatic margins exist near flood basalts (Seychelles)
- b) Rate of extension. But not all a-magmatic margins extended slowly (Seychelles)
- c) Inherited structure of the lithosphere, but is this relevant for the slow forming North East American margins?

What defines the steady-state shape of a margin?







What is the fate of a margin?

time

